INTRODUCTION

In marine environments, benthic communities below the euphotic zone are supported by primary production from pelagic or benthic photic areas, with a few exceptions such as hot vents and cold seeps. The magnitude of this supply of organic matter (OM) is determined by temporal and spatial patterns. Phytoplankton bloom sedimentation events represent a major source of OM for the benthic system where it fuels benthic life (Graf 1992). Moreover, the source and nature of the organic material settling on the substratum are important for benthic dynamics. Also, the fate of the organic material that reaches the sea floor is dependent on the nature of the sediment (Graf 1992).

In fine-grained depositional sites, sharp vertical profiles of labile OM (measured as chlorophyll a concentrations) can emerge after the sedimentation of
phytoplankton blooms (e.g. Franco et al. 2007). Mineralisation of this newly-arrived carbon often creates oxygen stress (Graf 1992) and breakdown of this OM can be retarded (Boon & Duineveld 1998, P. Provoost et al. unpubl. data). In coarser permeable sediments, these sharp vertical gradients may be absent and subsurface peaks of chl a are regularly reported (e.g. Ehrenhauss & Huettel 2004, Vanaverbeke et al. 2004b). Degradation of OM within these sediments is often rapid (Vanaverbeke et al. 2004b, Bühring et al. 2006). The different biogeochemical conditions in such contrasting sediments affect the responses of the resident bacterial (Franco et al. 2007) and nematode communities (Steyaert 2003, Vanaverbeke et al. 2004a). After phytoplankton sedimentation, nematode communities have variable response times (delayed in fine-grained sediments versus fast in permeable sediments) and variable diversity patterns (no change in fine-grained sediments versus increased in permeable sediments). In addition, an aberrant nematode morphotype (the stout and short nematodes defined by Vanaverbeke et al. 2004a) react opportunistically to phytoplankton sedimentation in permeable sediments (Vanaverbeke et al. 2004a), while this has not been observed at depositional stations. Nematode density responses are depth- and species-specific in fine-grained sediment (Steyaert 2003), while this is not the case in permeable sediments. Our hypothesis was that all these patterns are related to the quality and quantity of available food sources and the ability of nematodes to feed on particular food items (Moens & Vinçx 1997).

In this paper we take a closer look at nematode community dynamics at 2 sites with contrasting sediments in the North Sea. Both stations have been studied in detail (Steyaert 2003, Vanaverbeke et al. 2004a,b). The earlier studies suggest that meiobenthic responses to phytoplankton sedimentation may vary by sediment biogeochemical characteristics. Our aim was to: (1) investigate the overall community response (densities and vertical distributions) to the spring phytoplankton bloom deposition in 2 sediment types, and (2) use stable isotope analysis to determine how newly deposited phytodetritus is exploited by nematodes in 2 sediment types. Carbon (13C) and nitrogen (15N) stable isotope signatures provide powerful tools for estimating carbon flows to consumers and for determining their respective trophic positions in food webs (see Post 2002). Dual stable isotopic signatures of nematodes have mostly been reported at the community level (e.g. Riera et al. 1996, Riera & Hubas 2003) and detailed studies determining the food source of different nematode taxa are still rare (Carman & Fry 2002, Moens et al. 2005). Studies including other meiobenthic taxa (e.g. meiobenthic polychaetes, Halacaroidea and harpacticoid copepods) are also rare (Couch 1989, Carman & Fry 2002). Therefore, little is known about possible competition for food sources within the meiobenthos. Furthermore, there is little known about how meiobenthic organisms are able to exploit the gradient (Graf 1992) from high quality fresh food at the sediment-water interface to older food sources of lesser quality deeper down in the sediment (Rudnick 1989, Widbom & Frithsen 1995). We compared δ13C and δ15N signatures in nematodes (genus or group level) and other meiobenthic taxa, and in sediment particulate organic matter (POM) and suspended particulate matter (SPM) to clarify nematode assemblage responses to the spring phytoplankton bloom deposition on different sediment types. The null hypotheses tested were (1) there would be no differences in nematode vertical distribution among sampling times within stations, (2) there would be no differences in isotopic signals of the food sources and meiobenthic organisms among sampling dates or among sediment depths, and (3) there would be no differences in isotopic signals among meiobenthic taxa or nematode groups within stations.

**MATERIALS AND METHODS**

**Study site and sampling.** Samples were taken from the Belgian Continental Shelf (BCS) Stns 115bis (close to the coast, 51°09.2’N, 02°37.2’E; 13 m depth) and 330 (further offshore, 51°26.0’N 02°48.5’E; 20 m depth) (Fig. 1).

Stn 115bis is a depositional station characterised by the presence of fine sediments (mean grain size: 185 µm) with a small fraction of mud (4%) (Steyaert 2003), while Stn 330 consists of medium sand (mean grain size: 329–361 µm) without mud (Vanaverbeke et al. 2004a,b) and highly permeable sediments.

Sampling at the stations was conducted monthly from October 2002 until October 2003 from the vessels RV ‘Zeeleeuw’ or ‘Belgica’. Stn 330 was not sampled in December 2002 due to bad weather and rough seas.

The water column was sampled 3 m below the air-sea interface and 1 m above the sea floor using 10 l Niskin bottles. Pigment samples were collected by filtering 500 ml of water from each depth onto GF/F glass microfibre filters (i.d. 4.7 cm) using a vacuum pump. This procedure was repeated 3 times. The samples were kept in the dark, preserved at –20°C on board and stored at –80°C in the laboratory.

Sediment was sampled using a Reineck box corer (surface area 180 cm²) or another box corer with a larger surface area (February, April and October 2003 cruises). The box corer was deployed three times at each sampling station. In February only two box corer samples were taken from Stn 330.
From each box corer, 2 acrylic plastic cores (i.d. 3.6 cm) were taken, 1 for meiobenthos and 1 for pigment analysis. These cores were sliced in 1 cm slices to a maximum depth of 10 cm. Samples for pigment analysis were preserved at –20°C on board and stored at –80°C in the laboratory. The meiobenthos samples were preserved in a hot (70°C), neutral formaldehyde tap-water solution.

During 3 detailed sampling cruises (February, April and October 2003), 3 extra cores were taken on each sampling occasion at each station for studying carbon and nitrogen stable isotope signatures in sediment and meiobenthos. Samples were kept frozen at –20°C until further processing. For the stable isotope signatures in the SPM, the water column 1 m above the sea floor was sampled and filtered as described above.

Laboratory treatment of samples. The sediment samples for pigment analysis were weighed and chl a concentrations in the sediment were determined by HPLC (Gilson) following Wright & Jeffrey (1997).

Meiobenthos (animals passing through a 1 mm sieve and retained on a 38 µm sieve) was extracted from the sediment by centrifugation in a LUDOX HS-40 solution (Heip et al. 1985). After staining with Rose Bengal, all organisms were counted and sorted into higher taxa under a binocular microscope.

Stable isotope analysis. Each meiobenthic sample (3 replicates per sampling occasion at both stations) was unfrozen and the specimens were picked out with a fine needle under a binocular microscope. The organisms were rinsed, initially in 0.2 µm-filtered sea-water, then in 0.2 µm-filtered Milli-Q water to remove adhering particles, and transferred to tin capsules. The capsules were oven dried, pinched closed and stored (~–20°C) until further analysis. The procedure was repeated with 10 empty capsules to obtain blank values (5 for carbon and 5 for nitrogen).

About 60 and 160 nematodes were picked out from each capsule for $^{13}$C and $^{15}$N analyses, respectively. The taxonomic resolution of the samples taken for stable isotope signatures depended on the number of nematodes present. When sufficient numbers were available from the samples at Stn 330, ‘stout nematodes’ (as defined by Vanaverbeke et al. 2004a) were picked out separately from the rest of the nematodes. Other meiobenthic taxa (copepods, Halacaroidea, polychaetes) were also picked out separately. Nematodes from the genera Sabatieria and Richtersia (also a stout nematode) were picked out separately from the other nematodes at Stn 115bis. When densities were insufficient to analyse a specific group of nematodes or a specific meiobenthic taxon separately, these groups were included in the bulk nematode or meiobenthic sample. Replicates for all groups were not always obtained due to low abundances of nematodes. For the same reason, it was not always possible to estimate the $\delta^{15}$N of stout nematodes at Stn 330.

Stable isotope ratios of sediment POM, SPM (filters) and meiobenthos were measured by elemental analyser–isotope ratio mass spectrometry (EA-IRMS) (Middelburg et al. 2000). Data are expressed in standard $\delta$-unit notation:

$$\delta X = \left(\frac{R_{\text{sample}}}{R_{\text{reference}}} - 1\right) \times 10^3$$

where $R_{\text{sample}}$ is either the $^{13}$C/$^{12}$C ratio or the $^{15}$N/$^{14}$N ratio in the sample and $R_{\text{reference}}$ is the isotope ratio of the reference material. These ratios are reported as ‰ deviations from standards, viz. the carbon isotope ratio of Vienna Pee Dee Belemnite ($R_{\text{VPDB}} = 0.0112372$) and the nitrogen isotope ratio of air N$_2$ ($R_{\text{AIR}} = 0.0036765$).

As trophic fractionation is low for carbon isotopes (mean = 0.4 ‰, SD = 1.3 ‰) and high for $\delta^{15}$N (mean = 3.4 ‰, SD = 1 ‰), a consumer will have a $\delta^{13}$C value similar to its food source, whilst the $\delta^{15}$N value will be enriched by, on average, 3.4 ‰ relative to the food source (Post 2002).

Data analysis. Spearman rank correlation tests were used to investigate the correlation between the nematode densities and the chl a concentration in the sediment. Changes in nematode densities over time, sediment depth and time $\times$ depth were tested by constructing a univariate split-plot ANOVA design following Steyaert et al. (2001). Replicates were nested within time, but not within depth. This analysis was performed on both the whole sediment column (0 to 10 cm) and on only the upper 5 cm, since changes in
quantity and quality of OM were most evident there (P. Provoost et al. unpubl. data; present study) as was the response of the nematodes communities at Stn 115bis in a previous study (Steyaert 2003). Variation in the $^{13}$C signal of sediment POM by time, depth and time × depth was analysed using a 2-way ANOVA. Since it was not possible to obtain homogeneity of variances for $\delta^{13}$C and $\delta^{15}$N values of meiobenthos, Kruskal-Wallis (ANOVA by ranks) tests were conducted to test for differences between the meiobenthic taxa, sediment depth and sampling dates. All tests were performed using the STATISTICA 6 software package.

RESULTS

Environmental variables

Chl a concentrations in the water column at both stations started rising in February, peaked in April (48 mg m$^{-3}$ at Stn 115bis and 32 mg m$^{-3}$ at Stn 330) and decreased afterwards (Fig. 2). Smaller peaks were also observed in July for both stations and in September only at Stn 330, reaching values no higher than 17 mg m$^{-3}$. The sediment chl a concentrations were about 5 times higher at Stn 115bis (Fig 3) than at Stn 330 (Fig. 4) throughout the sampling period and generally followed the patterns observed in the water column. At both stations, peak values were observed in April. At Stn 330, other peaks were observed in August 2003 and October 2003 reflecting the deposition of the late summer and autumn blooms (Fig. 4).

At Stn 115bis, chl a concentration in the sediment was highest near the sediment surface and decreased with sediment depth, especially after deposition of phytodetritus in April (Fig. 3). Except for October 2003, this gradient was not observed at Stn 330. In most cases, there was no clear vertical gradient in chl a vertical distribution (until March 2003), while in the second half of the study period, only minor differences between surface and deeper chl a concentrations occurred (ca. 100 to 200 ng cm$^{-3}$) (Fig. 4).

Temporal patterns in meiobenthic densities

In total, 12 meiobenthic taxa were found (Nematoda, harpacticoid Copepoda and nauplii, Cumacea, Gastrotricha, Halacaroidea, Kinorhyncha, Oligochaeta, Ostracoda, Polychaeta, Tardigrada and Turbellaria) among which only Cumacea was never observed at Stn 330.

At Stn 115bis, Nematoda were highly numerically dominant, with an average representation of 96%. Nematodes also dominated at Stn 330, but made up a considerably lower proportion (64%). Here, other taxa were present with relatively high representation; harpacticoid Copepoda and nauplii represented 15% and 11% of the meiobenthos, respectively.

At Stn 115bis, nematode densities were lowest in winter (1251 ind. 10 cm$^{-2}$ in December) and increased towards April (4841 ind. 10 cm$^{-2}$) (Fig. 5). No clear changes were observed in the following months until densities increased to highest values in September and October 2003 (7385 ind. 10 cm$^{-2}$ in October). Nematode densities were not correlated with chl a in the sediment ($r$ = 0.30, $p$ = 0.067). The bulk of the nematode community was found in the upper 4 cm of the sediment column for most of the year, except for October 2003 when relatively high densities (804 ind. 10 cm$^{-2}$) were found down to 7 cm depth (Fig. 5). From June until September, higher densities were also found in the 9 to 10 cm layer.

Nematode densities in the top 5 sediment layers at Stn 115bis were not significantly affected by time ($F_{2, 4} = 0.59$, $p$ = 0.59), by sediment depth ($F_{4, 8} = 1.56$, $p$ = 0.27) nor by the interaction of time and depth ($F_{8,16} = 0.97$, $p$ = 0.48). When including the upper 10 cm in the analysis, a significant effect for depth was observed ($F_{9,18} = 7.35$, $p$ = 0.002), while this was not the case for time ($F_{2, 4} = 0.87$, $p$ = 0.487) or the interaction term ($F_{18, 36} = 1.19$, $p$ = 0.319).

Nematode densities at Stn 330 were much lower than at Stn 115bis. Lowest values were recorded in November (349 ind. 10 cm$^{-2}$), followed by a spring peak in May (1310 ind. 10 cm$^{-2}$) (Fig. 6). Densities decreased towards July and increased again in August.
and October; highest values were observed in October 2003 (1989 ind. 10 cm–2). At Stn 330, nematode densities were correlated with chl \(a\) in the sediment \((r = 0.36, p = 0.033)\). Subsurface maximum values were common, especially in October 2003 when densities were highest (Fig. 6). The “split-plot” ANOVA revealed that time \((F_{2, 4} = 23.78, p = 0.006)\), sediment depth \((F_{4, 8} = 4.14, p = 0.041)\) and their interaction \((time \times depth, F_{8,16} = 2.64, p = 0.046)\) significantly affected the nematode densities. A second analysis including all sediment layers (0 to 10 cm) showed that densities per sediment layer were significantly different between sampling events \((F_{3, 4} = 10.29, p = 0.041)\) and sediment depths \((F_{9,18} = 3.36, p = 0.014)\). However, the effect of their interaction was not significant \((F_{18, 36} = 1.53, p = 0.136)\).

**Carbon and nitrogen stable isotope signatures**

At Stn 115bis, sediment POM \(\delta^{13}C\) values were similar to the SPM \(\delta^{13}C\) values. Sediment POM \(\delta^{13}C\) values differed by sediment depth \((F_{2,12} = 15.53, p = 0.0005)\) (Fig. 7). Most depleted values \((-22\%o)\) occurred in February in both sediment layers (0 to 1 and 4 to 5 cm). In April, values increased slightly and this increase was more prominent in the uppermost centimetre of sediment. Sediment POM \(\delta^{13}C\) values in October were similar in both layers and resembled the values for April in the surface layer.

Vertical differences in nematode isotopic signatures also occurred. Before sedimentation of the phytoplankton bloom, surface-dwelling nematodes were about 3\%o enriched compared to the corresponding OM signatures. *Sabatieria* spp. and *Richtersia* spp. had similar values when sampled in the 4 to 5 cm layer. Deeper-dwelling ‘other nematodes’ were more depleted (Fig. 7). In April, all surface-living nematodes and *Sabatieria* from the deeper sediment horizon had \(\delta^{13}C\) values slightly above or equal to the OM values. Deeper-dwelling ‘other nematodes’ were more depleted. A similar pattern occurred in October for *Sabatieria* and *Richtersia*, while ‘other nematodes’ in both sediment layers had more depleted values than the OM signal. Very depleted values for \(^{13}C\) (average
value –38.53 and –38.54‰) and 15N (0.30 and 1.27‰) occurred in surface-living harpacticoid copepods in October.

At Stn 330, δ13C values of sediment POM did not follow the δ13C isotope values of SPM (Fig. 8). The δ13C values of sediment OM differed significantly between sampling dates (F2,10 = 32.15, p = 0.00004) (Fig. 8). More enriched values occurred in both sediment layers in October than in February and April. During February and April, all benthic groups from both sediment layers had average δ13C values 3‰ more enriched than the corresponding OM values in the sediment [February: –20.6 ± 0.3‰ for POM and –17.8 ± 0.2‰ for meiobenthos; April: –21.8 ± 0.3‰ for POM and –18.7 ± 0.2‰ for meiobenthos [means ± SE]]. In October, both OM and faunal average δ13C values were similar (–18.2 ± 0.4‰ for POM and –18.4 ± 0.2‰ for meiobenthos [means ± SE]).

At both stations, δ13C values were affected by time (Kruskal-Wallis, p = 0.000 and p = 0.020 for Stns 115bis and 330, respectively) and differed among meiobenthic taxa (Kruskal-Wallis, p = 0.010 and p = 0.021 for Stns 115bis and 330, respectively), while there were no significant differences between sediment layers (Kruskal-Wallis, p = 0.071 and p = 0.990 for Stns 115bis and 330, respectively). When the extremely depleted δ13C values for the copepods were excluded from the analysis for Stn 115bis, δ13C values of the different meiobenthic taxa were no longer significantly different (Kruskal-Wallis, p = 0.081), though there were significant differences among depths (Kruskal-Wallis, p = 0.027).

δ15N values for the meiobenthos were similar for the 2 stations (Fig. 7 and 8) and among the different meiobenthic taxa, except for the copepods at Stn 115bis in October (Fig. 7) (Kruskal-Wallis, p = 0.203 for Stn 115bis [not including copepods] and p = 0.450 for Stn 330). For both stations, there were no significant differences in δ15N between the 2 sediment layers (Kruskal-Wallis, p = 0.055 and p = 0.556 for Stns 115bis and 330, respectively). δ15N values were not significantly affected by time at Stn 330 (Kruskal-Wallis, p = 0.053) but did differ significantly over time at Stn 115bis (Kruskal-Wallis, p = 0.011).

**DISCUSSION**

**Environmental variables**

As described previously (Steyaert 2003, Vanaverbeke et al. 2004a,b), the concentrations in the water column indicated a strong phytoplankton bloom in
spring at both stations. The higher chl a concentrations at Stn 115bis can be explained by its position closer to the coast (Joint & Pomroy 1993).

Although chl a concentrations in sediment also indicated a seasonal signal, considerable differences between the stations were evident. Pigment concentrations at Stn 115bis were about 5 times higher than at Stn 330 (Fig. 3 & 4).

At the fine sandy Stn 115bis, phytodetritus accumulated near the sediment surface during spring, until mineralisation in late summer (P. Provoost et al. unpubl. data), as previously reported for other fine-sandy North Sea stations (Boon & Duineveld 1998).

In contrast, there were no clear vertical gradients of chl a in the coarser sediments of Stn 330. Relatively strong bottom water currents can prevent the deposition of sedimenting phytodetritus at the sediment surface of permeable sediments (Huettel & Rusch 2000). Through advective water flow, sedimenting phytoplankton cells can penetrate deeper in the sediment, inducing subsurface peaks (Huettel & Rusch 2000, Ehrenhauss & Huettel 2004). Such peaks occurred at Stn 330. Advective transport of oxygen into the sediment and fast removal of decomposition products (Huettel et al. 1998) accelerate POM degradation, resulting in a fast mineralisation of organic carbon and recycling of nutrients (Huettel & Rusch 2000, Bühring et al. 2006). These processes prevent a build up of labile OM and the establishment of clear vertical gradients as observed in the finer sediments at Stn 115bis.

**Temporal patterns in meiobenthic densities**

The stations differed in nematode responses to phytoplankton deposition. A fast increase in density at Stn 330 coincided with sedimentation of labile OM from the water column. Highest densities of ca. 600 ind. 10 cm⁻² were also reported in May 1999 at the same station (Vanaverbeke et al. 2004b); however, sampling only lasted until summer and there are no data in this earlier study (op. cit.) on the further responses of nematodes to late summer blooms (if present). This increase in nematode abundance was largely attributable to a rapid increase in densities of short and stout nematodes (Vanaverbeke et al. 2004a). These nematodes belong to the selective deposit feeder category sensu Wieser (1953), which has bacteria as an important food source (Moens & Vincx 1997). The rapid incorporation of algal biomass into bacterial biomass in these sediments (Bühring et al. 2006) allows this opportunistic response in the nematode communities, whose densities were correlated with chl a concentrations in the sediment. Moreover, there was an increase in nematode densities after each deposition event in our study, indicating rapid and independent

**Fig. 7. Stn 115bis. **$\delta^{13}$C and $\delta^{15}$N signatures for February, April and October 2003 from suspended particulate matter (SPM) in the water (1 m above the floor), $\bullet$, in organic matter (OM) in the sediment, and in different meiobenthic groups at 0–1 cm and at 4–5 cm depth. $\cdots \cdots$: mean of animals from station 330; $\cdots \cdots$: SD on either side of mean of animals from station 330. Cop: copepods; O.Nem: other nematodes; Richt: Richtersia; Sabat: Sabatieria
responses by the nematode community after each event. At Stn 330, the significant differences in the vertical distribution of nematodes in the upper 5 cm were probably the result of an upward migration of the animals towards their food source (Vanaverbeke et al. 2004b). There were no significant differences in nematode densities when we included all sediment layers in the analysis. There is no obvious explanation for this observation, but we suggest that it may be attributable to the continually low densities of nematodes at greater depths in this station.

In contrast, there were no significant differences in total nematode densities or vertical distributions in the upper 5 cm the of finer sediments at Stn 115bis, although sedimentation of phytodetritus clearly occurred. The lower densities in the sediment layers below 5 cm depth are the main reason why significant differences in densities among depths were found when we included the entire sediment column in the split-plot ANOVA. However, time of sampling and the interaction term were not significant in this analysis either. In these finer grained sediments, peak mineralisation of newly arrived OM is delayed until summer (Boon et al. 1998, P. Provoost et al. unpubl. data) allowing a more gradual but extended increase in nematode densities. This may explain why there was no correlation between total nematode densities and chl a at this station. There was also a gradual increase in nematode densities following the phytoplankton bloom of 1999 at the same station, with a first peak in April (ca. 2500 ind. 10 cm–2) and highest densities in July (slightly above 4000 ind. 10 cm–2) (Steyaert 2003). At both stations, nematode densities were lower in 1999 than in 2003; however, seasonal changes were similar.

The absence of clear temporal trends in the vertical distribution of the nematode community as a whole was previously reported for Stn 115bis (Steyaert 2003). However, vertical distribution patterns of dominant taxa did show species-specific temporal patterns that were related to food source partitioning among the dominant nematode species (Steyaert 2003). *Sabatiera celtica* and *S. punctata*, which were the dominant species during periods of low food availability, increased rapidly in density shortly after the arrival of fresh food. Both species concentrated at the sediment surface during that period. In May and June, higher densities of *Daptonema riemannii* and *D. fistulatum*, respectively, were encountered at greater depths. This coincided with the decomposition and burial of the fresh OM, triggering a seasonally-timed reproduction (Steyaert 2003).
Carbon and nitrogen stable isotope signatures

The $\delta^{13}C$ values of sediment POM were within the range reported in other studies in the Schelde estuary (The Netherlands) and in Tokachi (Japan) (Moens et al. 2002, Moens et al. 2005, Usui et al. 2006). Marine phytoplankton $\delta^{13}C$ values can vary from $-30$ to $-18\%$, but they are typically near $-22\%$ (Boutton 1991). Middelburg & Nieuwenhuize (1998) reported $\delta^{13}C \approx -18\%$, and $\delta^{15}N = 9\%$ for OM of marine origin in the Schelde estuary (The Netherlands), which is located less than 100 km from our study area. Megens et al. (2001) observed that $\delta^{13}C$ values in POM from the North Sea were more enriched in $^{13}C$ in spring and summer when the samples corresponded almost exclusively to fresh phytoplankton; most depleted $\delta^{13}C$ values were observed in the spring and summer when the samples were dominated by primary consumers, with $\delta^{13}C$ values in water SPM. Excluding copepods and deep-dwelling ‘other nematodes’ from Stn 115bis (discussed below), the meiobenthic $\delta^{13}C$ signals at both stations varied little over time (when compared to water SPM and sediment POM) and differed little between stations ($-18.81 \pm 0.18\%$, $-19.48 \pm 0.34\%$ and $-20.82 \pm 0.26\%$).

At Stn 115bis in April, more enriched POM $\delta^{13}C$ values were observed in the sediment surface than in the 4 to 5 cm layer, indicating sedimentation of phytodetritus. In October, POM $\delta^{13}C$ values in both sediment layers became more similar, probably as a result of downward transport and mineralisation during summer and late summer (P. Provoost et al. unpubl. data). At the coarser grained Stn 330, vertical differences in $\delta^{13}C$ values were less evident, probably as a result of the permeability of the sediment, which hindered the vertical separation of fresh and older OM (Huettel et al. 1998).

Since data on isotopic signals of meiobenthic organisms from shallow subtidal habitats are presently unavailable, we compared our results with those obtained from intertidal areas and from the Antarctic deep sea (Table 1). Our $\delta^{13}C$ signals in the meiobenthos were generally more depleted than previously reported for intertidal sediments (with the exception of Carman & Fry 2002 and Riera & Hubas 2003). However, nematodes ranged from only $-16$ to $-13\%$ in the study by Carman & Fry (2002). All these studies were conducted in intertidal habitats where microphytoplankton is common and $\delta^{13}C$ values are more enriched than those of organisms dependent solely on water column-derived OM or its degradation products. Moens et al. (2007) reported that nematode $\delta^{13}C$ values were generally more depleted in the Antarctic deep sea (130 to 2021 m) than those we observed (Table 1). However in the Antarctic deep sea, POM $\delta^{13}C$ values ($-24$ to $-23\%$) were also generally more depleted than those we observed.

Even though $\delta^{13}C$ values of meiobenthos at both stations changed significantly with time, these changes did not reflect the changes in sediment POM or water SPM. Excluding copepods and deep-dwelling ‘other nematodes’ from Stn 115bis (discussed below), the meiobenthic $\delta^{13}C$ signals at both stations varied little over time (when compared to water SPM and sediment POM) and differed little between stations ($-18.81 \pm 0.18\%$, $-19.48 \pm 0.34\%$ and $-20.82 \pm 0.26\%$).

The $\delta^{15}N$ values of sediment POM were within the range reported in other studies in the Schelde estuary (The Netherlands) and in Tokachi (Japan) (Moens et al. 2002, Moens et al. 2005, Usui et al. 2006). Marine phytoplankton $\delta^{15}N$ values can vary from $-34$ to $-31.8\%$, but they are typically near $-32\%$ (Boutton 1991). Middelburg & Nieuwenhuize (1998) reported $\delta^{15}N \approx -32.5\%$, and $\delta^{15}N = 9\%$ for OM of marine origin in the Schelde estuary (The Netherlands), which is located less than 100 km from our study area. Megens et al. (2001) observed that $\delta^{15}N$ values in POM from the North Sea were more depleted in $^{15}N$ in spring and summer when the samples corresponded almost exclusively to fresh phytoplankton; most depleted $\delta^{15}N$ values were observed in the spring and summer when the samples were dominated by primary consumers, with $\delta^{15}N$ values in water SPM. Excluding copepods and deep-dwelling ‘other nematodes’ from Stn 115bis (discussed below), the meiobenthic $\delta^{15}N$ values at both stations varied little over time (when compared to water SPM and sediment POM) and differed little between stations ($-27.9 \pm 0.18\%$, $-38.54 \pm 0.17\%$ and $-42.7 \pm 0.34\%$).

### Table 1. Natural stable carbon isotope signatures of meiobenthos

<table>
<thead>
<tr>
<th>Location</th>
<th>Habitat</th>
<th>Meiobenthic taxa</th>
<th>$\delta^{13}C$ (%)</th>
<th>Source</th>
</tr>
</thead>
<tbody>
<tr>
<td>North Inlet Estuary, Georgetown, USA</td>
<td>Intertidal mudflat</td>
<td>Nem, Cop</td>
<td>$-15.83$ to $-13.87$</td>
<td>a</td>
</tr>
<tr>
<td>Bay of Marennes-Oléron, France</td>
<td>Intertidal mudflat</td>
<td>Nem, Cop</td>
<td>$-16.4$ to $-15.3$</td>
<td>b</td>
</tr>
<tr>
<td>Gulf of Mexico, USA</td>
<td>Salt marsh mudflat</td>
<td>Nem, Cop, Ost</td>
<td>$-25$ to $-13$</td>
<td>c</td>
</tr>
<tr>
<td>Schelde estuary, Netherlands</td>
<td>Saltmarsh, tidal flats</td>
<td>Nem</td>
<td>$-18.48$ to $-14.03$</td>
<td>d</td>
</tr>
<tr>
<td>Roscoff Aber Bay, France</td>
<td>Intertidal sediments</td>
<td>Nem</td>
<td>$-27.0$ to $-14.8$</td>
<td>e</td>
</tr>
<tr>
<td>Schelde estuary, The Netherlands</td>
<td>Intertidal flat</td>
<td>Nem</td>
<td>$-16.5$ to $-12.5$</td>
<td>f</td>
</tr>
<tr>
<td>Schelde estuary, Netherlands</td>
<td>Intertidal flat</td>
<td>5 Nem spp.</td>
<td>$-17.9$ to $-12.8$</td>
<td>g</td>
</tr>
<tr>
<td>Weddell Sea, Antarctic</td>
<td>Deep-sea</td>
<td>Nem</td>
<td>$-19.7$ to $-19.3$</td>
<td>h</td>
</tr>
<tr>
<td>North Sea, BCS</td>
<td>Shallow subtidal</td>
<td>Nem, Cop, Hal, Pol, O.Meio</td>
<td>$-38.54$ to $16.02$</td>
<td>i</td>
</tr>
</tbody>
</table>

- Source: a: Couch (1989); b: Riera et al. (1996); c: Carman & Fry (2002); d: Moens et al. (2002); e: Riera & Hubas (2003); f: Moens et al. (2005); g: Steyaert et al. (2007); h: Moens et al. (2007); i: Present study.
in February, April and October, respectively, at Stn 115bis; −17.76 ± 0.21‰, −18.67 ± 0.21‰ and −18.45 ± 0.18‰ in February, April and October, respectively, at Stn 330; means ± SE). If there were significant changes in their C-sources throughout the year, these were not reflected in the δ13C signatures.

In mesocosm experiments lasting for 5 or 6 mo, meiobientic taxa showed preference for fresh detritus, even when not responding quantitatively to the OM input (Rudnick 1989, Widbom & Frithsen 1995). For the period we studied in the North Sea, meiobentic feeding preferences also remained relatively constant. Meiobentic organisms were feeding selectively, as previously proposed Moens et al. (2007) for the Arctic deep sea.

At Stn 115bis in October, the copepods had δ13C values of −38.53 and −38.54‰ that deviated considerably from other meiobentic signatures. Similar δ13C signatures were found in macrobenthos carrying symbiotic sulphur-oxidising chemosynthetic bacteria or having chemosynthesis-based nutrition (e.g. Tsutsumi et al. 2001, Levin & Michener 2002). Chemoautotrophic bacteria can obtain required energy by oxidation of reduced inorganic substrates (Fenical & Jensen 1993). The δ13C signal of the copepods from Stn 115bis indicates exploitation of a chemosynthetically-derived food resource.

At Stn 330, there were no vertical differences within the meiobentic δ13C values or in POM. This probably resulted from the permeable nature of the sediment, which inhibits the build up of vertical gradients in POM quality and quantity. However the stout nema-
todes, which showed the strongest response to phyto-
detritus deposition at this station (Vanaverbeke et al. 2004a), had δ13C values that were similar to other nematodes, copepods and polychaetes. Hence, our results do not indicate differences in carbon sources and/or feeding strategy. Stout and short nematodes can become dominant during certain periods of time and/or areas with permeable sediments and almost absent in others (Vanaverbeke et al. 2004a; Urban-Malinga et al. 2006). What life cycle feature makes these stout and short nematodes such successful colonizers is not yet known.

Except for the copepods at Stn 115bis in October 2003, there were no significant δ15N differences between the different meiobentic groups at either station. δ15N values similar to ours for nematodes have been reported in the Bay of Marennes-Oléron and in Roscoff Aber Bay, France, ranging from 8.9 to 9.3‰, and from 9 to 14.3‰, respectively (Riera et al. 1996; Riera & Hubas 2003). Slightly more depleted values than those we observed (ranging from ca. 3 to 8‰) occurred in the North Inlet Estuary, Georgetown and in the Gulf of Mexico, USA, respectively (Couch 1989;
Carman & Fry 2002). The more enriched δ15N values reported by Moens et al. (2005) for the Schelde estuary (reaching values of 20‰) correspond the predatory feeding of the respective species. These species, if present in our samples, would be found within the ‘other nematodes’. Even though the relative importance of their signal depends on the relative biomass of such consumers (for which we have no information), predators are generally not dominant (in terms of relative densities) within the nematodes at the sites studied (Steyaert 2003, Vanaverbeke et al. 2004b), so it is likely that their signature would be diluted in the ‘other nematodes’ signature.

In conclusion, the nematode community responses to the deposition of the phytoplankton bloom differed by the biogeochemical nature of the sediment; in fine sediments the prolonged presence of OM resulted in a gradual response of the nematode community, with highest densities long after the deposition of phyto- detritus. Depth-related differences in food-webs and feeding strategies occurred in these sediments. In permeable sediments, OM reaching the sediment was quickly mineralised and the whole system responded quickly. The genus Sabatieria’s ability to migrate towards a food source may be the reason for its success, being able to respond opportunistically to an OM input. On the other hand, stout nematodes showed no differences from other nematodes as initially expected. In general, the meiobiometric 13C signatures remained similar over time and were not coupled with changes in the water and sediment OM, indicating selective feeding.

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