Modeling transport of mixed sediments on the Belgian Continental Shelf

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The effects of anthropogenic impacts, such as harbour extension, dredging and dumping works, deepening of navigation channels and aggregate extraction, is becoming increasingly important. A good knowledge of the natural sediment balance to allow an appropriate evaluation of these anthropogenic impacts and to develop marine policies for the development of socio-economic activities in equilibrium with marine environmental protection. In parallel with the execution of high quality measurement, in situ and by remote sensing, also the application of numerical models can be a very powerful tool to get information on the sediment balance.

Two sediment transport models are applied at the Belgian Continental Shelf (BCS). The mu-STM model is an advection-diffusion model, suited for the simulation of fine material in suspension. The model accounts for erosion and sedimentation and for consolidation of the sea bed. The model was already applied to model the turbidity maximum area in the Belgian coastal waters, between Oostende and the mouth of the Westerschelde (Fettweis & Van den Eynde, 2003) or for the simulation of the dispersion of tracers on the BCS (Van den Eynde, 2004). The mu-SEDIM model on the other hand simulates the transport of sand, by applying a local sand transport formulae. The model was already applied for the modeling of the kink in the Westhinder sand bank (Deleu et al., 2004), for the simulation of different scenarios of sand extraction on the Kwinte Bank (Van den Eynde, et al., 2008) or for the assessment of the wave influence on the sand transport on the Kwinte Bank (Giardino et al., 2008).

However, it is becoming clear that the division between non-cohesive sand and cohesive mud is artificially and that in the bed, a mixture of sediments usually is found. The cohesive sediments can contain a varying amount of sand, while the sandy sediments usually contain a varying amount of mud. The mud and sand content of a sea bed influences the transition between cohesive and non-cohesive sediments, which has a major influence on the erosion behaviour and possibly on the benthic ecological properties (Flemming & Delafontaine, 2000).

For sandy, non-cohesive sediments it is possible to calculate reasonably well the critical erosion stress and the erosion rate when only grain-size and sorting is known. For cohesive sediments the erodability depends mainly upon the degree of cohesion, consolidation, aggregation and compaction, the biological community structure, the sand content of the sediment and the layering. Mixed sediments may behave as cohesive or non-cohesive sediments, depending on the mud content, e.g. the addition of mud increases the sediment shear strength and thus the erosion threshold of a sandy bed (Williamson & Torfs, 1996).

Furthermore, these mixed sediments do not necessarily occur as well mixed. Frequently alternation of sand and mud layers are observed. These alternations are interpreted as tidal or storm/calm weather influences. Fan et al. (2004) describe storm waves as random destructive factors of the sediment bed and highlight them as effective agents of sediment transport and deposition of the sand-dominated layers. This type of segregation can only occur if cohesive suspended sediment concentration is low.
Recently, some studies have been published that take into account mutual interactions between cohesive and non-cohesive sediments (Van Ledden, 2002; van Ledden et al., 2004; Waeles, 2005; Waeles et al., 2007; Sanford, 2008). It is clear that more accurate predictions of the sediment transport are possible if the mutual influence of sand-mud mixtures is incorporated in the model.

For the BCS, a combined model is set up, to model both the transport of the material in suspension and the transport of the non-cohesive sand. A detailed bed model will be set up, which can represent the erosion, sedimentation, consolidation and bioturbation of the mixed sediments. The bed model will take into account the detailed distribution of the sediment on the BCS, the clay content and the critical erosion stress of the bed. Newly collected measurements of grain size distribution, bulk density, critical erosion stress and erosion rate (Westrich and Jancke, 2007, 2008a, 2008b) will be incorporated.

Some first results of the model will be presented.

References


