

Stratigraphic ranges of mosasaurs in Belgium and the Netherlands (Late Cretaceous) and cephalopod-based correlations with North America

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Abstract

Mosasaur taxa currently known from Campanian-Maastrichtian strata in the Liège-Limburg (SE Netherlands, NE Belgium) and Mons (southern Belgium) basins are listed and briefly discussed and their stratigraphic ranges indicated. Recently published and/or ongoing work on coleoid and ammonoid cephalopods in these areas allows tie points between NW Europe and the United States (Western Interior, Gulf Coast, Atlantic Seaboard) to be established. Future studies need to refine the resultant, rather crude, scheme. The ultimate aim is a more robust picture of mosasaur taxonomy, biostratigraphy and palaeobiogeography, and a detailed evaluation of migratory patterns across the Atlantic.

Keywords: Mosasauridae, Late Cretaceous, Campanian, Maastrichtian, biostratigraphy, palaeobiogeography, correlation, cephalopods

Introduction

To date, at least eight species of mosasaur have been recorded from strata of Campanian-Maastrichtian age in the extended type area of the Maastrichtian Stage (Liège-Limburg and Aachen area, Germany). A comparable number is known from the Mons Basin (southern Belgium) where, with the exception of *Globidens dakotensis* and possibly also of *Prognathodon giganteus*, all species stem from a narrow interval within the lower Maastrichtian (*obtusa* Zone).

The main aim of the present contribution is to put these mosasaur taxa in a biostratigraphic context, using cephalopod molluscs. In NW Europe, belemnite coleoids constitute the most reliable correlation tools, despite the fact that several species appear to be confined to particular areas (see e.g., Christensen, 1975, 1986, 1990, 1991, 1993, 1994, 1995, 1996, 1997a, b, 1998, 1999, 2000a, b; Christensen et al., 1975, 2000; Christensen & Schmid, 1987; Christensen & Schulz, 1976; Schulz, 1979, 1982; Schulz & Schmid, 1983). Unfortunately, none of the taxa recognised in Europe is known from the United States, where merely two (or three) species of *Belemnitella* are listed (Christensen in Kennedy et al., 1998). In the Western

Interior, *Belemnitella bulbosa* ranges from the *Hoploscaphites birkelundae* to *Jeletzkytes nebrascensis* zones (Fox Hills Formation, South Dakota), and possibly also occurs in the *Baculites baculus* and *B. clinolobatus* zones (Pierre Shale Formation, South Dakota) (see Table 1). The other species, *Bt. americana*, is known from the Mount Laurel Formation and basal Navesink Formation of New Jersey, Delaware and Maryland (upper upper Campanian-lower Maastrichtian); in South Carolina, Christopher & Prowell (2002) showed it to range through the lower and middle Peedee Formation (upper lower Maastrichtian). Samples of *Bt. americana* from the upper Mount Laurel Formation of New Jersey are currently being revised by N. Keutgen (pers. comm., October 2004). So long as both taxonomy and stratigraphic ranges remain poorly understood, coleoids cannot be used for intercontinental correlations.

Although the resultant scheme is still a crude one, ammonites are more promising in this respect. Below, ammonite-based correlations between the Liège-Limburg and Mons basins on the one hand, and the Gulf Coast, Atlantic Seaboard and Western Interior on the other, are given (Table 2). Added are distribution data of ammonite taxa in Europe and beyond. It is hoped that future work will lead to a refinement of this scheme, and

a tighter control on the stratigraphic ranges of mosasaur taxa on both sides of the Atlantic. Only in this way can comparisons between areas become more meaningful (compare Gallagher, 1993; Bardet & Pereda Suberbiola, 1996; Mulder, 1999, 2003a), nomenclatorial issues be resolved, and habitat preferences (compare Kiernan, 2002) be assessed in more detail. Such cephalopod-based correlations have great potential, as the recent suggestion of an intercontinental marine extinction event

Table 1. Correlation, based on ranges of selected ammonoid cephalopods, between the Campanian-Maastrichtian of the Western Interior, Gulf Coast and Atlantic Seaboard (US) (after Cobban & Kennedy, 1991a, 1992b; Kennedy & Cobban, 1993b) and Europe (Kaplan et al., in press; pers. obs.), with standard white chalk section of NW Germany as reference. * added by Landman et al. (2004) as highest ammonite zone in the uppermost Maastrichtian of New Jersey, Alabama, Mississippi and Missouri.

	US	NW Europe (NW Germany)
MAASTRICHTIAN	<i>Discoscaphites iris</i> *	
	<i>Jeletzkytes nebrascensis</i>	
	<i>Hoploscaphites nicolletii</i>	
	<i>Hoploscaphites birkelundae</i>	<i>fastigata, cimbrica</i> + <i>tegulatus/junior</i>
	<i>Baculites clinolobatus</i>	
	<i>Baculites grandis</i>	
	<i>Baculites baculus</i>	<i>obtusa</i>
	<i>Baculites eliasi</i>	
CAMPANIAN	<i>Baculites jenseni</i>	<i>grimmensis/</i> <i>granulosus (= hyatti)</i>
	<i>Baculites reesidei</i>	
	<i>Baculites cuneatus</i>	
	<i>Baculites compressus</i>	
	<i>Didymoceras cheyennense</i>	
	<i>Exiteloceras jenneyi</i>	
	<i>Didymoceras stevensoni</i>	
	<i>Didymoceras nebrascense</i>	
	<i>Baculites scotti</i>	
	<i>Baculites reduncus</i>	
	<i>Baculites gregoryensis</i>	
	<i>Baculites perplexus</i>	
	<i>Baculites</i> sp. (smooth)	<i>roemeri, vulgaris/</i> <i>stolleyi</i>
	<i>Baculites asperiformis</i>	
	<i>Baculites mclearni</i>	
	<i>Baculites obtusus</i>	<i>basiplana/spiniger</i>
	<i>Baculites</i> sp. (weak flank ribs)	
	<i>Baculites</i> sp. (smooth)	
	<i>Scaphites hippocrepis</i> III	<i>pilula to conica/</i> <i>gracilis</i>
	<i>Scaphites hippocrepis</i> II	<i>lingua/quadrata</i>
	<i>Scaphites hippocrepis</i> I	
	<i>Scaphites leei</i> III	<i>granulataquadrata</i>

across the lower/upper Campanian boundary shows (Lindgren, 2004a, b; Lindgren & Siverson, 2002, 2004). This is not only reflected in mosasaur distribution, but also in a turnover of ammonite taxa, in particular pachydiscids and scaphitids.

Abbreviations

To denote the repositories of material referred to in the text, the following abbreviations are used: BMNH - The Natural History Museum, London; MNHN - Muséum National d'Histoire Naturelle, Paris; NHMM - Natuurhistorisch Museum Maastricht, Maastricht; IRScNB - Institut royal des Sciences naturelles de Belgique, Brussels.

1 – Type area of the Maastrichtian Stage (SE Netherlands, NE Belgium) (Fig. 1)

In recent years, two working quarries in the Eben Emael area (Bassenge, province of Liège, NE Belgium) have yielded a number of (semi-)articulated mosasaur skeletons, in particular from the upper Gulpen and lower/middle Maastricht formations. Both lower and higher in the section, mosasaur material usually comprises disarticulated and isolated remains only, especially tooth crowns and vertebrae.

Mosasaur taxa of Campanian age

To date, only a handful of specimens (tooth crowns and poorly preserved post-cranial elements) are known from strata of Campanian age.

A – Vaals Formation

A single tooth crown (NHMM VG 2184), collected autumn 1965 from outcrop 62G-10 along the Lüttigerstraße at Kelmis (La Calamine, NE Belgium) and subsequently identified as a plioplatacarpine (*Platecarpus* sp.) by Kuypers et al. (1998), is perhaps better referred to the tylosaurine genus *Hainosaurus*, thus adding another tylosaurine record for NW Europe (Bardet, 1990; Jagt et al., 2005). Macrofossil assemblages observed in the section exposed there suggest it to have been correlatable with either the Gemmenich Member or Vaalsbroek Member, of early (though not earliest) Campanian age.

The Vaals Formation in its type area ranges from the *granulataquadrata* to at least the *pilula* zones of the NW German scheme (Table 3), and possibly even higher (see below). The base of the *granulataquadrata* Zone in Germany corresponds to the *Scaphites leei* III Zone in the Western Interior, dated at 83.5 ± 0.5 Ma (McArthur et al., 1993). Ammonite taxa particularly suited for trans-Atlantic correlation include *Scaphites hippocrepis* II-III (sensu Cobban, 1969) and *Baculites vaalsensis*. In the Western Interior, the *S. hippocrepis* II Zone has been dated at 81 Ma (McArthur et al., 1994).

Table 2. Occurrences of selected ammonoid taxa in North America (Atlantic Seaboard, Gulf Coast and Western Interior Seaway) and the Maastrichtian type area, plus Mons Basin. Names are formations and/or members (compare Figs. 1, 2). Abbreviations: AL – Alabama; CA – California; CO – Colorado; DE – Delaware; MO – Missouri; MD – Maryland; MS – Mississippi; NC – North Carolina; ND – North Dakota; NJ – New Jersey; SD – South Dakota; TN – Tennessee; TX – Texas; WY – Wyoming.

Taxa		US	Mons	Liège-Limburg
Desmoceratidae	<i>Hauericeras rembda</i>	Prairie Bluff (AL)	-	Vijlen 6
Kossmaticeratidae	<i>Brahmaites brahma</i>	Arkansas (B. sp.)	-	Meerssen
Pachydiscidae	<i>Pachydiscus (P.) gollevillensis</i>	Prairie Bluff (MS)	-	pre-Emael
	<i>Pachydiscus (P.) j. jacquoti</i>	Prairie Bluff (MS)	-	pre-Emael, Nekum
		Corsicana (TX)		
	<i>Pachydiscus (P.) neubergicus</i>	Navesink (NJ)	Ciply-Malogne	Vijlen 1-6
	<i>Pachydiscus (P.) haldemisi</i>	Pecan Gap (TX)	-	Zeven Wegen
Placenticeratidae	<i>Hoplitoplacenticeras (H.) marroti</i>	Anacacho (TX)	-	Vaals (top)
Sphenodiscidae	<i>Sphenodiscus lobatus/binckhorsti</i>	Fox Hills (SD, ND)	-	Nekum-Meerssen
		Pierre Shale (SD, ND)		
		Escondido (TX)		
		Corsicana (TX)		
		Ripley (MS)		
		Providence (AL)		
		Peedee (NC)		
		Severn (MD)		
		Red Bank/Tinton (NJ)		
Diplomoceratidae	<i>Glyptoxoceras aquisgranense</i>	Merchantville (NJ)	-	Vaals
	<i>Glyptoxoceras rugatum</i>	Severn (MD)	-	Vijlen 6-Meerssen
		Corsicana (TX)		
Nostoceratidae	<i>Nostoceras spp.</i>	Kemp Clay (TX)?	-	Meerssen (base)
Scaphitidae	<i>Jeletzkytes dorfi</i>	Fox Hills (CO, WY)	-	Vijlen 6
		Pierre Shale (CO)		
	<i>Scaphites (S.) hippocrepis</i> II-III	Merchantville (NJ, MD, DE)	-	Vaals
		Gober (TX)		
		Western Interior		
	<i>Trachyscaphites s. spiniger</i>	Mooreville (AL)	Nouvelles	Vaals (top)
		Ozan (TX)		Zeven Wegen
		Annona (TX)		
Baculitidae	<i>Baculites aquilaensis</i>	Lower Campanian, Western Interior	Obourg (base)	-
	<i>Baculites baculus</i>	Maastrichtian, Western Interior, Canada,		
		US Gulf Coast	Ciply-Malogne	-
	<i>Baculites sp. (smooth)</i>	mid-Campanian, Western Interior	Obourg (base)	Zeven Wegen (top)
	<i>Baculites vaalsensis</i>	Merchantville (NJ)	-	Vaals
	<i>Baculites vertebralis</i>	Severn (MD)	-	Emael-Meerssen
	<i>Eubaculites carinatus</i>	Prairie Bluff (MS, AL)	-	Meerssen
		Owl Creek (MO, MS, TN)		
		Corsicana (TX)		
		Hornerstown (NJ)		
	<i>Trachybaculites columna</i>	Fox Hills (CO, SD)	-	Vijlen 6
		Prairie Bluff (AL, MS)		
		Corsicana (TX)		
		Garzas (CA)		

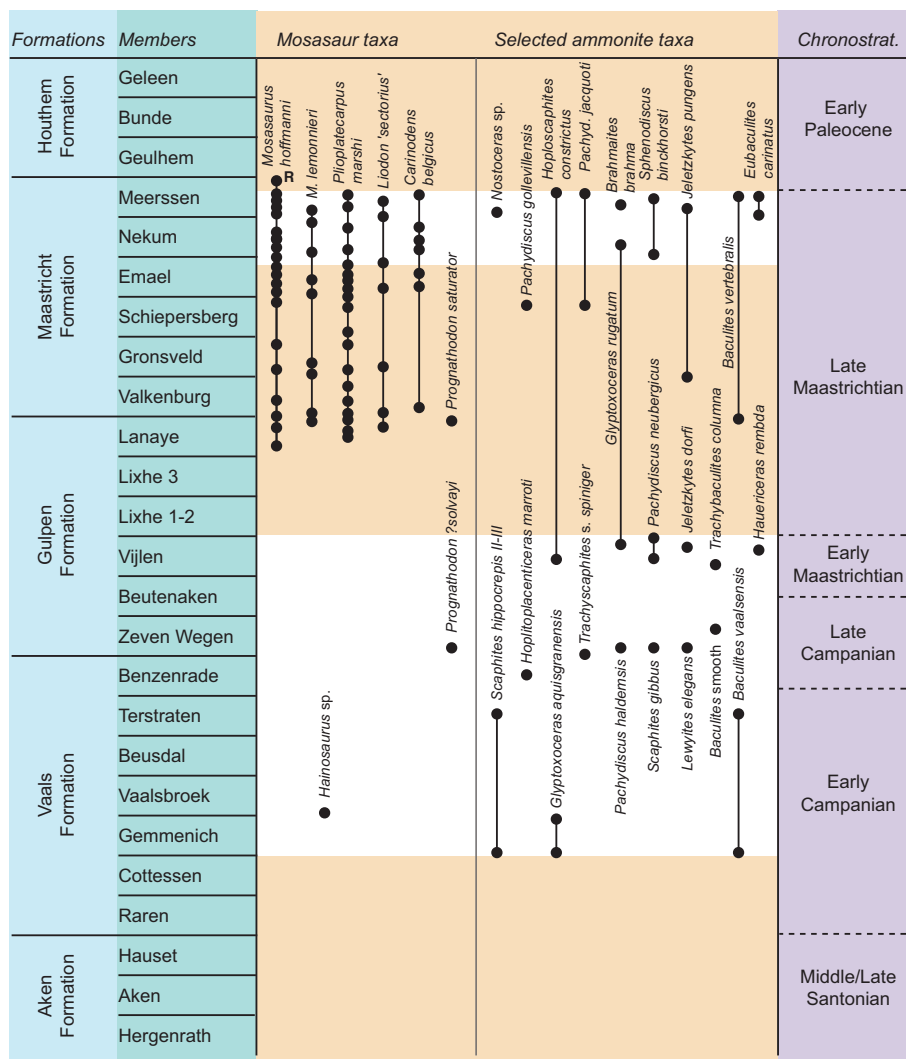


Fig. 1. Lithostratigraphy of Santonian-lower Paleocene strata in the Maastrichtian type area (after Felder & Bosch, 2001) and stratigraphic ranges of mosasaur taxa known to date, plus those of selected ammonoid cephalopods. White are portions of the sequence which correlate well with North America, using particular ammonite taxa.

Scaphites (*S.*) *hippocrepis* II is known from the Mooreville Formation in Alabama (Kennedy et al., 1997a), while *hippocrepis* III has been recorded from the Merchantville Formation of New Jersey, Maryland and Delaware (Kennedy & Cobban, 1993d; Kennedy et al., 1997c). In addition, it is widely distributed in the lower Campanian of the Western Interior (Montana, Wyoming, South Dakota, Utah, Colorado, New Mexico), and questionably occurs in the Blufftown Formation (Alabama) and Woodbury Formation (New Jersey). Passage forms between *hippocrepis* II and III occur in NE Belgium, Aquitaine and Provence (France; see Kennedy, 1986b), Hampshire and Sussex (England; see Gale, 1980), Münsterland (Wippich, 1995; Hauschke et al., 1999), the Lehrter Westmulde (Hannover area, Germany; see Niebuhr, 1996), southern Sweden (Kennedy & Christensen, 1997), southern Poland (Jagt et al., 2004) and Israel. Form III has been described from the Gober Chalk of Texas (Cobban & Kennedy, 1992b) and from Saskatchewan (Canada; see Cobban, 1993) as well. The species complex may range into the 'middle' Campanian of NE Texas (Cobban & Kennedy, 1992a) and the Barranca (northern Spain; Küchler, 2000).

Baculites vaalsensis occurs in the lower Campanian of NE Belgium, the Vaals-Vaalseberg area (the Netherlands), northern Aquitaine (France), Münsterland (Kaplan et al., in press) and the Merchantville Formation of New Jersey (Kennedy & Jagt, 1995; Kennedy et al., 1997c; Klinger & Kennedy, 2001).

Glyptoxoceras aquisgranense, well known from the Lower Campanian of NE Belgium, Ariège (France), the Aachen area (Germany) and the Münsterland Basin (Kennedy & Kaplan, 1995), has recently been described from the Merchantville Formation of New Jersey (Kennedy et al., 1997c).

Other ammonite species from the Vaals Formation have correlative value mostly within Europe. *Placenticeras bidorsatum* (see Kennedy & Jagt, 1995) appears to be confined to the *lingua/quadata* Zone in Münsterland and the Braunschweig area (Kennedy & Kaplan, 1995; Kaplan et al., in press), and a similar range has been described for northern Aquitaine (Kennedy, 1986b). *Pachydiscus* (*P.*) *duelmensis* has a comparable range in Münsterland (Kennedy & Kaplan, 1995; Wippich, 1995) and northern Aquitaine (Kennedy, 1986b). *Pachydiscus* (*P.*) *launayi* also occurs in the latter area (Kennedy, 1986b) and possibly in the upper lower Campanian of Münsterland

(Wippich, 1995) as well, while *Eupachydiscus lewyi* is also known from the mid-lower Campanian of Navarra (Küchler, 2000), the upper lower Campanian of Münsterland (Wippich, 1995; Hauschke et al., 1999), southern France (Alpes Maritimes), northern Spain, Poland, Kopet Dagh and Madagascar.

The upper part of the Vaals Formation near Cottessen (Jagt et al., 1995a), and the Benzenrade Member in the Benzenrade area (the Netherlands), have yielded different ammonite assemblages. The single taxon in this assemblage to have been recorded from both sides of the Atlantic is *Hoplitoplacentceras* (*H.*) *marroti*, which is now known from Ariège (Kennedy et al., 1992b), Portugal, Israel, central Asia, Madagascar and south-central Texas (Anacacho Limestone; Kennedy & Cobban, 2001). In northern Spain (Küchler, 2000), *H. (H.) marroti* is typical of the basal upper Campanian, as it is in northern Aquitaine (Kennedy, 1986b).

Pachydiscus (*P.*) *colligatus* occurs in southern Sweden (uppermost lower Campanian to upper Campanian), Brabant (Belgium), northern Aquitaine, possibly Tercis and ?Madagascar

(Kennedy, 1986b, 1987; Kennedy & Christensen, 1997; Kennedy & Jagt, 1998), while *Pachydiscus* (*P.*) *subrobustus* has been recorded from southern Sweden (Kennedy & Christensen, 1997), Münsterland, Tercis, Ariège (Kennedy & Bilotte, 1995), Austria, Pontus (Turkey), Ukraine and southern Poland (Kaplan et al., 1996; Machalski et al., 2004).

Patagiosites stobaei ranges from the upper lower Campanian to the ?lower roemeri Zone, being best known from the basal *stobaei/basiplana* Zone (Table 3), with records from northern Germany (Kaplan et al., 1996; Kennedy & Kaplan, 1997), southern Sweden, Donbass, southern Belgium, province of Liège and southern Limburg (the Netherlands).

B – Gulpen Formation (lower portion)

From the lower upper Campanian Zeven Wegen Member, there is a questionable record of *Prognathodon solvayi*, based on a single tooth crown from Heure-le-Romain (Liège, NE Belgium) (Kuypers et al., 1998). From the basal portion of this unit in

Table 3. Biozonation of the Campanian-Maastrichtian of Lägerdorf-Kronsmoor-Hemmoor (LKH), the Münsterland Basin and the Lehrter Westmulde (Hannover area, Germany; after Kaplan et al., in press; Niebuhr, 2003, 2004; Niebuhr et al., 1997); 1 – Dülmen Formation, lower portion; 2 – Dülmen Formation, upper portion; 3 – Holtwick Formation, Legden Member, lower part; 4 – Holtwick Formation, Legden/Asbeck members, in part; 5 – Holtwick Formation, Asbeck Member, middle part; 6 – Holtwick Formation, Asbeck Member, upper part; 7 – Holtwick Formation, Darfeld Member; 8 – Coesfeld Formation, lower part; 9 – Coesfeld Formation, middle part; 10 – Coesfeld Formation, upper part.

	LKH	Münsterland	Lehrter Westmulde
MAASTRICHTIAN	(kazimiroviensis)	-	-
	baltica/danica	-	-
	danica/argentea	-	-
	argentea/junior	-	-
	tegulatus/junior	-	-
	fastigata	-	-
	cimbrica	-	-
	sumensis	-	-
	obtusa	-	-
	pseudobtusa	-	-
	lanceolata	-	-
CAMPANIAN	grimmensis/granulosus	-	-
	langei	-	bipunctatum/roemeri
	polyplocum	polyplocum ¹⁰	minor/polyplocum
	roemeri	roemeri ⁹	vulgaris/stolleyi
	basiplana/spiniger	vulgaris/basiplana ⁹	vulgaris/basiplana
	basiplana/stobaei	basiplana/stobaei	stobaei/basiplana
	conica/mucronata	conica/mucronata ⁸	conica/mucronata
	gracilis/mucronata	gracilis/mucronata ⁷	gracilis/mucronata
	conica/papillosa	conica/papillosa ⁶	conica/papillosa
	papillosa	papillosa ⁵	papillosa
	senonensis	senonensis ⁴	senonensis
	pilula/senonensis	pilula/senonensis ³	pilula/senonensis
	pilula	pilula ²	pilula
	lingua/quadrata	lingua/quadrata ¹	lingua/quadrata
	granulataquadrata	granulataquadrata	-

the Haccourt area nearby, a number of poorly preserved mosasaur vertebrae are known, which must remain indeterminate. From strata ('Tuffeau jaune', not 'jaunâtre') at Folxles-Caves (province of Brabant, Belgium) which, on calcareous nannoplankton evidence, Mulder & Mai (1999) considered to be of early late Campanian age, these authors recorded the tylosaurine *Hainosaurus* cf. *bernardi*. However, Robaszynski et al. (2002) noted that that age assignment must be based on reworked material, since the 'tuffeau jaune', or Jauche Member, contains the typically late Maastrichtian brachiopod *Thecidea papillata* in great numbers.

Coleoid cephalopods, which are much commoner than ammonites in this member, demonstrate this unit to be equivalent to the *conica/mucronata*, *stobaei/basiplana*, *vulgaris/basiplana* and *vulgaris/stolleyi* zones (see Keutgen & Jagt, 1999) (Table 3).

Ammonite taxa used for trans-Atlantic correlation include *Trachyscaphites* s. *spiniger* and *Baculites* sp. (smooth; sensu Cobban, 1962). Typical members of the former taxon are known from the Arcola Limestone Member (Mooreville Formation) in Alabama (Kennedy et al., 1997a), the Ozan Formation and Annona Chalk in Texas (Cobban & Kennedy, 1992a). European records include northern Germany (Niebuhr, 1996), southern Limburg, province of Liège, Mons Basin, Aquitaine, northern Spain, southern Sweden, central and southern Poland, Russia, Ukraine, Armenia and Turkmenistan (Błaszkiwicz, 1980; Kennedy, 1986b; Kennedy & Kaplan, 1997; Küchler, 2000; Jagt et al., 2004; Machalski et al., 2004). In northern Germany, it is common in the lower *basiplana/spiniger* Zone (Table 3), but extends up into the *roemeri* Zone (Kaplan et al., in press).

Of other species and subspecies of *Trachyscaphites* recorded from North America, just a single one, *T. pulcherrimus* is known from Europe as well, ranging into the *polyplacum* Zone, with records from central Poland, Aquitaine, northern Spain, Germany, Austria, Russia and Armenia (Kennedy & Summesberger, 1984; Kennedy, 1986b; Niebuhr, 1996; Kennedy & Kaplan, 1997; Niebuhr et al., 1997; Küchler, 2000). American records include the Wenonah Formation in New Jersey (Kennedy & Cobban, 1994a) and the Bergstrom Formation of central Texas (Kennedy & Cobban, 1999). *Trachyscaphites* s. *porchi* has been described from the Wolfe City Sand (NE Texas; Cobban & Kennedy, 1993), the Pecan Gap Chalk (central and NE Texas), Anacacho Limestone (south-central Texas; Kennedy & Cobban, 2001), Demopolis Formation (Texas), from the *Baculites mclearni* and *B. asperiformis* zones (Montana, Colorado) (Table 1), and possibly from the upper Mishash Formation in Israel (Cobban & Kennedy, 1994b). *Trachyscaphites redbirdensis* is known from the *Baculites perplexus* and *B. gregoryensis* zones in Wyoming, South Dakota, Colorado and Arkansas (Kennedy & Cobban, 1993b), while *T. densicostatus* seems to be confined to the Middle Campanian Ozan Formation of NE Texas (Cobban & Kennedy, 1992a). From Maastrichtian strata, *T. alabamensis*

has been described from the Prairie Bluff Formation of Alabama and Mississippi (Cobban & Kennedy, 1995) and *T. yorkensis* from the Prairie Bluff of Alabama and the Corsicana Marl of Texas. In Europe, this lineage appears to have become extinct during the late Campanian, *T. pulcherrimus* being the last representative. Following that, species of *Jeletzkytes* (uppermost Campanian), *Acanthoscaphites* (lower Maastrichtian) and *Hoploscaphites* (lower-upper Maastrichtian) dominated scaphitid faunas there.

The second species of correlative value is *Baculites* sp. (smooth, sensu Cobban, 1962), which occurs in the uppermost Zeven Wegen Member in the province of Liège, east of Haccourt. As noted by Kennedy (1993), in the northern Western Interior this species defines a zone between the *Baculites asperiformis* Zone below and the *B. perplexus* Zone above, in the 'middle Campanian' (see Table 1).

Nostoceras hyatti (= *N. pozaryskii*), which is typical of the upper upper Campanian in North America, is also known from various European localities (see Kennedy et al., 1992a; Küchler & Odin, 2001; Küchler et al., 2001). Unfortunately, no material is yet known from Liège-Limburg, although the Beutenaken Member is the age-correlative of strata yielding this ammonite elsewhere (Keutgen & Jagt, 1999). Kennedy & Cobban (1993c), Cobban & Kennedy (1994a), Küchler (2000) and Kennedy et al. (2000a) and Odin et al. (2001) recorded *N. hyatti* from the Ripley Formation of Tennessee, the Saratoga Chalk of Arkansas, the Nacatoch Sand in NE Texas, the basal Navesink Formation of New Jersey, the Pierre Shale (*Baculites jenseni* Zone) of southern Colorado, plus the upper Campanian of Angola, ?Israel, Tercis and Aquitaine (France), northern Spain and central Poland.

Recently, a fourth species, *Pachydiscus* (*P.*) *haldemsi*, has been documented from the Pecan Gap Chalk of central and northeast Texas (Cobban & Kennedy, 1994b). This is also known from Austria, Aquitaine, northern Spain, southern Sweden, northern Germany, northern Ireland, Norfolk, central and southern Poland, Ukraine and Turkmenistan (Kennedy & Summesberger, 1984; Kennedy, 1986b; Kennedy & Christensen, 1997; Kennedy & Kaplan, 1997; Küchler, 2000; Machalski et al., 2004).

Other ammonite taxa known from the Zeven Wegen Member are predominantly European in distribution, e.g., *Patagiosites stobaei* (see above), *Patagiosites* aff. *griffithi* (northern Ireland), *Hoplitoplacenticeras* (*H.*) cf. *coesfeldiense* (Wyoming, Austria, northern Germany, Aquitaine, Russia; Cobban, 1963; Kaplan et al., 1996, in press; Kennedy & Kaplan, 1997; Kennedy & Summesberger, 2001), *Scaphites* (*S.*) *gibbus* (northern Germany, southern Belgium, province of Liège, southern Limburg, Aquitaine, southern Poland, Donbass, Ukraine and Kazakhstan; Wippich, 1995; Kaplan et al., 1996; Kennedy & Kaplan, 1997; Jagt et al., 2004; Machalski et al., 2004), *Lewyites elegans* (northern Germany, southern Sweden, Aquitaine, southern Poland, Austria; Kennedy & Christensen,

1997; Kennedy & Kaplan, 1997; Jagt et al., 2004; Machalski et al., 2004; Summesberger & Kennedy, 2004), and *Neancyloceras? phaleratum* (northern Spain, ?Austria, ?Lehrter Westmulde, central Poland; K  chler, 2000). Close to the last-named species is *N.? bipunctatum* from the upper Campanian of Aquitaine, northern Germany, central Poland, Austria, and possibly northern Spain and the Ozan Formation of Arkansas (Kennedy & Cobban, 1993b; Kennedy & Kaplan, 1997; Niebuhr et al., 1997; Kennedy & Summesberger, 1999, 2001; K  chler, 2000).

The Beutenaken I and II members, of late Campanian (upper *polyplacum* and lower *langei* zones) to early Maastrichtian (upper *grimmensis/granulosus* to *pseudobtusa* zones) age (Table 4), have so far not yielded any ammonites, nor mosasaur remains. Biozonation and correlation with Norfolk (England) and the Mons Basin rely primarily on coleoids (Keutgen & Jagt, in prep.). Current interpretations of the Zeven Wegen and Beutenaken members assume large sedimentary gaps in the middle to upper upper Campanian, which, together with possible facies-related differences, may explain the absence in the Li  ge-Limburg Basin of such key index ammonite taxa as *Hoploscaphites vistulensis*, *H. pumilus*, *Pseudokosmaticeras galicianum* and *Jeletzkytes compressus*. The first-named is known from central Poland (*Nostoceras pozaryskii* Zone, upper upper Campanian) and the Mount Laurel Sand of Delaware (B  laskiewicz, 1980; Kennedy & Cobban, 1994b), while *H. pumilus* has been recorded from Haute Garonne, Tercis (Landes) and Petites Pyr  n  es (France), the Saratoga Chalk of Arkansas, the Nacatoch Sand in Texas and the basal Navesink Formation of New Jersey (Kennedy et al., 1986, 2000b; Kennedy & Cobban, 1993c; Machalski & Odin, 2001). The third species, *P. galicianum*, is known from the Saratoga Chalk of Arkansas, as well as from the uppermost Campanian of central Poland and the lower Maastrichtian of Austria, Italy, Bulgaria, the Ukraine, Crimea, Armenia and the Bithynian Peninsula (Turkey) (Kennedy & Cobban, 1993c). Finally, *J. compressus* has been recorded from the upper Campanian of northern Ireland, northern Germany, central Poland and from the Mount Laurel Sand of Delaware (Kennedy & Cobban, 1994b); Kennedy & Kaplan (1997) listed it from the upper Haldem Schichten (*polyplacum* Zone), while Niebuhr (1996) and Niebuhr et al. (1997) collected specimens at Hannover-Ahlten (upper *bipunctatum/roemerii* Zone). Its congener, *J. nodosus*, is known from the Ripley Formation (Coon Creek Tongue) of Tennessee, the Nacatoch Sand in NE Texas, the basal Saratoga Chalk of southwest Arkansas, the basal Navesink Formation of New Jersey and the Pierre Shale (*Baculites compressus*, *B. cuneatus*, *B. reesidei* and *B. jenseni* zones) of the Western Interior, as well as from the upper Campanian of central Poland. From the basal Navesink Formation of New Jersey and the *B. compressus* Zone of the Pierre Shale in Colorado, respectively, Kennedy et al. (2000b) and Cobban et al. (1992) mentioned *J. cf. nodosus*.

An additional typically North American taxon, *Didymoceras stevensoni*, index of the *stevensoni* Zone in the Western Interior,

has been recorded to co-occur with European species such as *Pseudokosmaticeras brandti* and *Pachydiscus (P.) subrobustus* by Kennedy & Bilotte (1995) in Ari  ge (sub-Pyrenees, southern France).

Mosasaur taxa of Maastrichtian age

C – Gulpen Formation (upper portion)

The Vijlen Member (Table 4), subdivided into seven intervals (numbered 0 to 6) and ranging from the *obtusa* Zone to the lower part of the *tegulatus/junior* Zone, has yielded merely a single tooth and two tooth crowns from the CPL SA-Haccourt and ENCI-Maastricht bv quarries. For the time being, these are best referred to the genus *Mosasaurus*, and seem close (?ancestral) to *M. hoffmanni* (compare also Machalski et al., 2003). From the *sumensis* Zone at Altembroeck (Voer, NE Belgium), there is a fragmentary dentary, which may prove conspecific with one of the species recorded from the Mons Basin (see below); so far, no detailed description of this jaw has appeared (compare Jagt et al., 1995b).

In the extended type area of the Maastrichtian Stage, coleoids again constitute the best correlation tools (Keutgen, 1996, 1997; Keutgen & Van der Tuuk, 1991). Ammonites are much rarer, but the taxa that have been recorded do include such that can be used for trans-Atlantic correlation (Table 2). Vijlen Member interval 6 (= Vijlen 6) has yielded (Jagt & Kennedy, 1994; Kennedy & Jagt, 1998; Jagt, 2002):

- *Jeletzkytes dorfi*, occurring in the *Hoploscaphites birkelundae* Zone (upper Pierre Shale and Fox Hills Formation of Colorado and Fox Hills in Wyoming; Landman & Waage, 1993; Landman & Cobban, 2003);
- *Trachybaculites columna*, known from the Prairie Bluff Chalk of Alabama and Mississippi, the Fox Hills Formation of South Dakota and ?Colorado, the Corsicana Formation of Texas as well as the Garzas Formation in California. All other Western Interior records are from the zones of *Hoploscaphites nicolletii* and *Jeletzkytes nebrascensis* in the Fox Hills Formation (Kennedy & Cobban, 1993a; Cobban & Kennedy, 1995; Klinger & Kennedy, 2001; Landman & Cobban, 2003);
- *Pachydiscus (P.) neubergicus*, recorded from the Western Interior and Atlantic Coastal Plain (Navesink Formation, New Jersey), as well as from northern Germany, southern Limburg and Li  ge, Denmark, southern Belgium, Landes (France), northeast and ?southeast Spain, Austria, the ?Czech Republic, central Poland, Bulgaria, Ukraine, European Russia, Balochistan, South India, Nigeria, Zululand, Madagascar, United Arab Emirates/Oman border region and ?Sakhalin (Kennedy & Summesberger, 1986; Kennedy & Henderson, 1992a; Kennedy et al., 1995, 2000b; Jagt, 2002; Jagt & Felder, 2003; Niebuhr, 2003; W  greich et al., 2003); and

Mons Basin	Maastricht-Aachen-Liège	Lägerdorf-Kronsmoor-Hemmoor	
	Interval 6	<i>tegulatus/junior</i> Zone (pars)	MAASTRICHTIAN
	Interval 5	<i>fastigata</i> Zone	
		<i>cimbrica</i> Zone	
	Intervals 2-4	upper <i>sumensis</i> Zone	
	Interval 1	middle <i>sumensis</i> Zone	
		lower <i>sumensis</i> Zone	CAMPANIAN
Ciply-Malogne Fm.	Interval 0	<i>obtusa</i> Zone	
		<i>pseudobtusa</i> Zone	
	Beutenaken II Member	<i>lanceolata</i> Zone	
		<i>grimmensis/granulosus</i> Zone	
Spiennes Fm.		<i>langei</i> Zone	
	Beutenaken I Member	<i>polyplocum</i> Zone	

Table 4. Correlation between the Spiennes and Ciply-Malogne formations in the Mons Basin (southern Belgium), the Beutenaken and Vijlen members (Gulpen Formation) in the extended type area of the Maastrichtian Stage (SE Netherlands-NE Belgium; Aachen area, Germany), and the combined Lägerdorf-Kronsmoor-Hemmoor sections (northern Germany) (after Keutgen & Jagt, in prep.).

- *Hauericeras* cf. *rembda*, also documented from the Bay of Biscay sections, South India, Madagascar, Zululand, NE Mexico and the Prairie Bluff Chalk of Alabama (Kennedy & Henderson, 1992a; Ward & Kennedy, 1993; Cobban & Kennedy, 1995; Ifrim et al., 2004).

Other ammonites from the Vijlen Member, and intervals 1, 2 and 6 in particular, are either cosmopolitan or strictly European in distribution, and include:

- *Pachydiscus* (*P.*) aff. *armenicus* (Bay of Biscay sections, Tercis (Landes), Pyrénées Atlantiques and Armenia; Kennedy & Hancock, 1993; Ward & Kennedy, 1993; Courville & Odin, 2001);
- *Diplomoceras cylindraceum* (occurring worldwide; Klinger & Kennedy, 2003a, b);
- *Hoploscaphites constrictus* (ranging from the base to the very top of the Maastrichtian; see e.g., Machalski, 1996; Niebuhr, 2003);
- *Hoploscaphites tenuistriatus* (narrow range in northern and eastern Europe);
- *Acanthoscaphites* (*A.*) *tridens* (*lanceolata* and *sumensis* zones in northern Germany, Poland, Ukraine, NE Belgium and unspecified Maastrichtian of Kangerlussuaq (East Greenland); Kennedy & Summesberger, 1987; Jagt et al., 1999; Niebuhr, 2003; Kelly et al., 2004);
- *Acanthoscaphites* (*Euroscaphites*) *varians blaszkiewiczii* (confined to Vijlen 5 and 6 in the Maastrichtian type area, and ranging from the *cimbrica* Zone to low in the *junior* Zone in Denmark; Jagt et al., 1999); and
- *Baculites knorriani* (northern Germany, Denmark, Poland, Czech Republic, Ukraine and NE Belgium (Birkelund, 1993; Kennedy & Summesberger, 1987; Kennedy & Christensen, 1997).

In recent years, inoceramid bivalves have been shown to constitute valuable tools in European trans-Atlantic correlations (Walaszczyk, 2004; Walaszczyk et al., 1996, 2001, 2002a, b). From all intervals of the Vijlen Member, but the *sumensis* and lower *tegulatus/junior* zones (= Vijlen 1 - 4, 6) in particular, inoceramids have been collected. These are currently being studied (Walaszczyk & Jagt, in prep.).

The Lixhe 1 - 3 members have so far only yielded coleoids (*Belemnitella junior* group), of early late Maastrichtian age. The overlying Lanaye Member is also very poor in ammonites, with the exception of phosphatised baculitids, apparently of European stock and thus not suitable for correlation with the North American biozones. However, as far as mosasaurs are concerned, this unit documents an important shift in faunal distribution (Fig. 1), closely linked to a change from an open oceanic to a marginal marine setting. This member has yielded at least five species (*Mosasaurus hoffmanni*, *M. lemonnieri*, *Plioplatecarpus marshi*, *Prognathodon saturator* and *Liodon 'sectorius'*; see Dortangs et al., 2002; Jagt et al., 2002). The holotype of *P. saturator* (NHMM 1998141) is the sole specimen referable to this taxon, at least for the time being. It may turn out that some isolated tooth crowns referred to other species (NHMM and private collections) actually belong to this species.

The holotype of *P. marshi* is IRScNB R38 (Dollo, 1882) from the 'craie grisâtre à silex gris', near Zichen (NE Belgium), close to Eben Emael, which in all probability would correspond to the Emael Member. Recent fieldwork in the Eben Emael area has shown that the first appearance datum (FAD) of this species is within the Lanaye Member (Gulpen Formation), and that the Emael Member (Maastricht Formation) yields the best-preserved material, inclusive of partially articulate skull and post-cranial material, confirming earlier records (Lingham-Soliar, 1994). It should be noted that it is possible that

Plioplatecarpus depressus from the 'Maastrichtian greensands of New Jersey', the type of which is lost (Russell, 1967; Gallagher, 1993) is conspecific with *P. marshi*. Confirmation or rejection of this observation must await the discovery of additional material from New Jersey; of *P. marshi* the remains of at least ten individuals are now known, allowing proper documentation of range of skeletal variation.

D – Maastricht Formation

This unit, of late Maastrichtian age (zones of *Belemnitella junior* and *Belemnella (Neobelemnella) kazimiroviensis*), has produced the best-preserved mosasaur material, especially from the Eben Emael area (Bassenge, Liège, NE Belgium), and from the Emael and Nekum members in particular. At least five species of mosasaur are now known from the Maastricht Formation, although, as a consequence of limited material available, a few of these remain poorly known. In addition, relationships between taxa referred to in previous literature as *M. hoffmanni*, *M. lemonnieri* and *L. 'sectorius'* (sensu Lingham-Soliar, 1993) have yet to be determined. Recent finds of partial skulls preserving dentary, maxillary and pterygoid teeth are currently being assessed; these should provide definitive clues for identification of isolated tooth crowns.

Although *Mosasaurus hoffmanni* (holotype: MNHN AC 9648; see Lingham-Soliar, 1995) first occurs in the Lanaye Member (Gulpen Formation), and possibly even lower in the section (see above), the best material has been collected from the Emael and Nekum members near Eben Emael, Maastricht (Sint Pietersberg) and Bemelen, from which at least ten fragmentary, yet articulated, individuals have been collected in the past three decades.

To *Mosasaurus lemonnieri* have been assigned isolated tooth crowns and a single partial dentary (Kuypers et al., 1998). It is still not clear whether these remains are truly conspecific with material from the lower Maastrichtian (*obtusa* Zone; see below) of the Mons Basin or not. In addition, in view of their rarity in comparison to tooth crowns of *M. hoffmanni*, it may be that these remains represent stray animals and/or floating carcasses. Even more tantalising is the recent suggestion by Mulder et al. (2004) that *M. lemonnieri* may in fact be nothing more than the juvenile of *M. hoffmanni* (see below).

Of *Liodon 'sectorius'*, collections available comprise isolated tooth crowns and fragmentary dentaries of at least two individuals (Jagt et al., 2002). The type material (Russell, 1967) is probably from the Navesink Formation of New Jersey, but Gallagher (1993) noted that other material may stem from the Hornerstown Formation, of late Maastrichtian to early Paleocene age. Since not much remains of the type material, this taxon cannot be properly evaluated. Furthermore, there seem to be differences in dental morphology between the two dentaries referred to this taxon.

A previous record (Lingham-Soliar, 1996) of halisaurine mosasaurs from the Maastrichtian type area has now been shown to be based exclusively on *Plioplatecarpus marshi* (Mulder, 2003b), a common species in the area (see previous).

Carinodens fraasi (holotype: IRScNB R43), based on a jaw of unknown stratigraphic provenance (Dollo, 1913), has been formally synonymised with *Carinodens belgicus* by Schulp et al. (2004), a species first described from the Mons Basin (see following), and ranges throughout the Maastricht Formation (Fig. 1). Not much is known of this species, with the exception of isolated tooth crowns and two dentaries, one of them heavily broken and abraded.

Mosasaur populations appear to have been well established at the time of deposition of the lower/middle Maastricht Formation, suggesting relatively little (if any) exchange with populations elsewhere, and with a decline towards the K/T boundary. The highest unit of the Maastricht Formation, the Meerssen Member, yields comparatively little mosasaur material, isolated pterygoid teeth of *P. marshi* being amongst the highest finds known to date. This may be related to the extremely shallow nature of the Meerssen depositional setting, with numerous hardgrounds and omission surfaces (?subaerial exposure). Albeit, this proves that mosasaurs did extend up to within less than a metre from the K/T boundary, at least in the study area, contrary to claims made by Sullivan (1987, table 1) that mosasaurs became extinct in the earliest Maastrichtian (around 74.5 Ma). For additional records of mosasaurs from very close to the K/T boundary, both in Alabama (Gulf Coast) and in Denmark, reference is made to Kiernan (2002) and Lindgren & Jagt (2005), respectively.

As far as ammonite distribution within the Maastricht Formation is concerned, it should be noted that assemblages from the 'tuffaceous chalk facies' in the Maastricht-Geulhem area on the one hand and those from the 'Kunrade limestone facies' in the Kunrade-Benzenrade area (southern Limburg, the Netherlands) on the other, have certain elements in common, but are otherwise quite different (Kennedy, 1987; Jagt, 2002). Species useful in trans-Atlantic correlations include:

- *Eubaculites carinatus*, known from the Prairie Bluff Formation in Mississippi and Alabama, the Owl Creek Formation of Missouri, Mississippi and Tennessee, the Corsicana Formation of NE Texas, and remanié in the Hornerstown Sand of New Jersey, plus Bay of Biscay sections, southwest France, north-west Spain, Austria, Mozambique, Zululand, South India, Western Australia, Argentina, Chile and California (Kennedy et al., 1986; Henderson et al., 1992; Klinger & Kennedy, 1993; Ward & Kennedy, 1993; Cobban & Kennedy, 1995; Kennedy & Cobban, 1996, 2000);
- *Baculites vertebralis*, which is common in the Emael, Nekum and Meerssen members, and occurs widely in Manche, Petites Pyrénées (France), Denmark, northern Germany, Poland, southern Russia and ?Tunisia. It has recently been recorded

- from the Severn Formation of Maryland (Kennedy et al., 1997b);
- *Glyptoxoceras rugatum*, recorded from the Severn Formation of Maryland and the Corsicana Formation of Texas (Kennedy et al., 1997b, 2001), South India, Brazil, ?Chile, Western Australia, northern Spain and Pyrénées-Atlantiques, France (Kennedy & Hancock, 1993). In the study area, the first representatives of this species are known from Vijlen Member, interval 6;
 - *Sphenodiscus binckhorsti*, which is closely related to, if not conspecific with *S. lobatus*. The latter has been documented from the Severn Formation of Maryland, the Escondido Formation of Trans-Pecos Texas and northern Mexico, the Corsicana Formation of northeast Texas, the upper Ripley Formation of Mississippi, the Prairie Bluff Formation of Alabama and Mississippi, the Providence Sand in Georgia, the upper Peedee Formation in North Carolina, the Red Bank and Tinton sands in New Jersey, as well as the *Hoploscaphites nicolletii* and *Jeletzkytes nebrascensis* zones (perhaps underlying *H. birkelundae* Zone as well) in the Western Interior, plus Israel and Nigeria (Cobban & Kennedy, 1995; Kennedy et al., 1997b);
 - *Pachydiscus (P.) gollevillensis*, well known from Manche (France), Lleida (Spain), Bay of Biscay sections, Poland, northern Germany, Austria, Armenia, northern Caucasus, Bulgaria, the Bithynian Peninsula (Turkey) and Madagascar (Kennedy, 1986a; Ward & Kennedy, 1993). There is a recent record from the Prairie Bluff Chalk of Mississippi (Cobban & Kennedy, 1995);
 - *Pachydiscus (P.) j. jacquoti*, recorded from Manche, Pyrénées Atlantiques (France), Bay of Biscay sections, Armenia, Madagascar, central Poland (Machalski, in prep.), ?British Columbia and ?South India, and also known from the Prairie Bluff Chalk of Mississippi and the Corsicana Formation of Texas (Kennedy, 1986a; Kennedy & Hancock, 1993; Ward & Kennedy, 1993; Cobban & Kennedy, 1995; Kennedy et al., 2001);
 - *Nostoceras* sp. (spp.); the single specimen of a nostoceratid known from the Maastrichtian type area is from the basal Meerssen Member at Berg en Terblijt (former Blom quarry; see Van der Tuuk & Zijlstra, 1979) – it shows affinities with *N. colubriformis* from the Saratoga Chalk (upper upper Campanian) of Arkansas (Kennedy & Cobban, 1993c) and the Nacatoch Sand in Texas, as well as with *Turrilites' saundersorum* from the basal Paleocene of Texas, presumed to be derived from the underlying Kemp Clay of mid-Maastrichtian age (Kennedy, 1987);
 - The scaphitid genus *Jeletzkytes* – Maastrichtian strata in North America have yielded *J. criptonodosus* from the *Baculites baculus* Zone (Canadian Western Interior), the *B. grandis* Zone of Colorado, Wyoming, Montana and South Dakota, as well as the middle Navesink Formation of New Jersey (Kennedy et al., 2000b), the Prairie Bluff Chalk in Alabama (Cobban & Kennedy, 1995) and the Ripley Formation of Mississippi. Occurring, above the *Baculites clinolobatus* Zone, in the Fox Hills Formation in North and South Dakota and Nebraska, as well as in the Severn Formation of Maryland is *J. nebrascensis* (Kennedy et al. 1997b). It appears that *Hoploscaphites pungens* from the Gronsveld, Emael, Nekum and basal Meerssen members in the Maastrichtian type area represents the NW European counterpart of this lineage. Ornament of phragmocones and body chambers of macroconchs of this species indicate close affinities, and suggests that it is best referred to *Jeletzkytes* (Jagt, Kennedy & Yazykova, in prep.);
 - *Brahmaites brahma*; of note is the record of *Brahmaites* sp. from the *Nostoceras alternatum* Zone (Maastrichtian) of SW Arkansas by Cobban & Kennedy (1991b), associated with *Baculites claviformis* which ranges from the uppermost Campanian (Coon Creek Tongue, base of Ripley Formation, with *Nostoceras hyatti*) into the Owl Creek Formation, co-occurring with *Eubaculites carinatus*. In Europe, *B. brahma* has been recorded from the Bay of Biscay sections and Dordogne (France), plus material from the Meerssen Member of the Maastrichtian type area (Jagt, Kennedy & Yazykova, in prep.). Elsewhere, it occurs in the upper Maastrichtian of South India and Tunisia (Kennedy & Henderson, 1992a; Goolaerts et al., 2004), and similar forms are known from Madagascar.
- Other ammonite taxa from the Maastricht Formation either show a mainly European distribution or are known from across the globe, with the exception of North America. These include:
- *'Menuites' fresvillensis* (Manche, Pyrénées Atlantiques, Petites Pyrénées (France), Denmark, Yugoslavia, Armenia, South India, Madagascar, Western Australia, Chile, ?Brazil, South Africa (Kwa Zulu); Kennedy, 1986a; Kennedy et al., 1986; Klinger et al., 2001);
 - *'Menuites' terminus* (Bay of Biscay sections, central Poland, Bulgaria, Denmark and Crimea; Ward & Kennedy, 1993; Machalski & Jagt, 1998; Yazykova, Jagt & Naidin, in prep.);
 - *Diplomoceras cylindraceum* (see previous);
 - *Baculites anceps* (Manche, Russia, ?Spain, Bay of Biscay sections; Kennedy, 1986a; Ward & Kennedy, 1993);
 - *Acanthoscaphites (?Euroscaphites) verneuilianus* (Manche; Kennedy, 1986a);
 - *Hoploscaphites constrictus* (Manche, Petites Pyrénées, southern Belgium, northern Germany, Denmark, northern Spain, Bay of Biscay sections, southern Sweden, Poland, Austria, Bulgaria, Carpathians, Donbass, Crimea, Transcaspia and Kopet Dagh (Kennedy, 1986a; Ward & Kennedy, 1993; Machalski, 1996));
 - *Phylloptychoceras* sp., of which two specimens have been collected from the Meerssen Member at Maastricht and Geulhem (Jagt, Cremers, Goolaerts & Verhesen, in prep.); *P. siphon* has been recorded from the Upper Maastrichtian of

the Bay of Biscay sections, ?California and South India, and perhaps also from Denmark (Kennedy & Henderson, 1992b; Birkelund, 1993; Ward & Kennedy, 1993);

- *Pachydiscus (P.) noetlingi*, which to date has been recorded from the Kunrade Limestone facies and the lower Nekum Member in the Maastrichtian type area and the Upper Maastrichtian Korara Shale of Balochistan, Pakistan (Jagt & Kennedy, 2003);
- *Hoploscaphites* gr. *waagei/angmartussutensis*, known from a single specimen of the upper Meerssen Member at Geulhem (Kennedy & Jagt, 1998); this species is best known from Greenland (Birkelund, 1965; see also Kennedy et al., 1999; Kelly et al., 2004).

2 – Ciply area, Mons Basin (southern Belgium) (Fig. 2)

For a general overview of strata exposed in the Ciply area, reference is made to Robaszynski & Martin (1988), Robaszynski & Christensen (1989), Robaszynski et al. (1988) and Dupuis & Vandycke (1990). Only two units, currently referred to as the Obourg Chalk Formation and the Ciply-Malogne Phosphatic Chalk Formation (Robaszynski et al., 2002) appear to have produced mosasaur material.

The ‘Craie (brune) phosphatée de Ciply’, of Early Maastrichtian age (*obtusa* Zone), as based on coleoid faunas (Christensen,

1999), has yielded numerous articulated mosasaurs, since this unit was intensively quarried mostly in subterranean galleries (La Malogne), for the production of agricultural fertilisers during the second half of nineteenth century. Close co-operation between the engineers A. Lemonnier and L. Bernard and L. Dollo (Muséum d’Histoire naturelle de Belgique, Brussels) in the period between 1880 and 1895 resulted in the recovery of fifty-two relatively complete skeletons from these works. As Lingham-Soliar & Nolf (1990) noted, the bulk of these collections is referable to *Mosasaurus lemonnieri*. Interesting is their observation that there are virtually no shark teeth associated with these mosasaur finds; large-scale scavenging can thus be ruled out, which may also explain the comparatively large number of (semi-)articulated skeletons.

From an unspecified level within the Obourg Chalk Formation near Mons comes an isolated tooth crown (IRScNB R41), assigned to *Globidens alabamaensis* by Dollo (1924; see also Lingham-Soliar, 1999, pl. 5, fig. 7a-c), which Russell (1975) thought would better be placed in *G. dakotensis*. *Globidens alabamaensis* is well known from the Selma Chalk of Alabama and Mississippi, while *G. dakotensis* occurs in the Wolfe City or Lower Marl members (Taylor Formation) in Texas, the ?Marlbrook Marl in Arkansas and the lower Pierre Shale of South Dakota. Kiernan (2002, fig. 2) showed *G. alabamaensis* to occur in the upper Mooreville Chalk and Arcola Limestone.

The base of the Obourg Chalk Formation has remanié components, such as *Baculites aquilaensis*, which in the Western Interior occurs in the upper lower Campanian (Table 2), associated with *Scaphites hippocrepis* III. The overlying Nouvelles Chalk Formation has yielded *Trachyscaphites* cf. *spiniger*, while *Scaphites gibbus* ranges through the Trivières and Obourg formations (Fig. 2), confirming records from northern Germany. This unit correlates well with the Zeven Wegen Member in the Haccourt area, as based on coleoid cephalopod (Christensen, 1999) and brachiopod evidence (Simon, 2000; Simon & Owen, 2001).

As far as the ‘Craie phosphatée’ is concerned, the following can be stated. In the literature, seven mosasaur species have been described from this unit, as follows:

- *Hainosaurus bernardi* (holotype: IRScNB R23; see Dollo, 1885a, b, 1889c, 1904), type species of the genus, collected from a former complex of abandoned and partly filled-in quarries near Mesvin (Ciply) (see Lingham-Soliar, 1992);
- *Bottosaurus belgicus* (holotype: BMNH R.1802; see Woodward, 1891), originally considered to be crocodilian in nature, but later recognised as mosasaur (Dollo, 1924). The exact provenance of the type specimen is unclear; Woodward (1891, p. 114) noted that, ‘Among the vertebrate remains discovered by M. Houzeau de Lehaie in the ‘Craie brune phosphatée’ of Ciply, (...)’. In the caption to pl. 3, fig. 18, he stated, ‘Lower Danian (Craie brune), Ciply’. Here, there seems to be confusion between the ‘Craie phosphatée’ (= Ciply-Malogne Phosphatic Chalk Formation in current terminology;

Lithostratigraphy	Mosasaur taxa	Ammonite taxa	Chronostrat.
Saint-Symphorien	<i>Mosasaurus lemonnieri</i> <i>Platycarpus houzeaui</i> <i>Halsaurus ortlebi</i> <i>Prognathodon solvayi</i> <i>Hainosaurus bernardi</i>	<i>Pachydiscus neubergicus</i> <i>Baculites baculus</i> <i>Hoploscaphites constrictus</i>	Late Maastrichtian
Ciply-Malogne	● ● ● ● ●	● ● ●	Early Maastrichtian
Spiennes	●	●	
Nouvelles	●	●	
Obourg	<i>Prognathodon giganteus</i> <i>Globidens dakotensis</i>	<i>Baculites aquilaensis</i> <i>Baculites smooth</i>	Late Campanian
Trivières	●	● ●	Early Campanian

Fig. 2. Lithostratigraphy of Campanian-Maastrichtian strata in the Mons Basin (southern Belgium; modified from Dupuis & Vandycke, 1990; Robaszynski et al., 2002) and stratigraphic ranges of mosasaur taxa described from the area. Added are records of selected ammonoid taxa (after Kennedy, 1993; pers. obs.), and white are portions of the sequence which correlate well with North America using particular ammonite taxa.

Robaszynski et al., 2002) of early Maastrichtian age (*obtusa* Zone), and the Ciply Formation of middle and late Danian age (Laga et al., 2002). If the type was collected from the uppermost portion of the 'Craie phosphatée', i.e. directly below the 'Poudingue de Malogne' hardground, matters are even more complicated (see below). For the time being, it is assumed that the type of *B. belgicus* stems from the 'Craie phosphatée' proper and is thus of early Maastrichtian age, and older than occurrences in the Liège-Limburg Basin, until new stratigraphically well localised material from the Mons Basin becomes available;

- *Plioplatecarpus houzeau* (holotype: IRScNB R35 (ex 3111); see Dollo, 1889c; Lingham-Soliar, 1994), which is closely related to *P. marshi* from the upper Gulpen Formation and Maastricht Formation (see above). Newly collected material from the Eben Emael area suggests that distinctions between these taxa may not be as clear cut as indicated by Lingham-Soliar (1994);
- *Prognathodon solvayi* (holotype: IRScNB R33 (ex 4672); see Dollo, 1889a - c), from Mesvin (Solvay quarry; acquired 1889). Other material in the IRScNB collections is both from the same quarry and from the Houzeau quarry (Spiennes, acquired 1894);
- *Prognathodon giganteus* (holotype: IRScNB R106 (ex 3103), from the Solvay quarry, Spiennes; see Dollo, 1904; Lingham-Soliar & Nolf, 1990). Dollo (1904, p. 213) merely indicated that his new species (no diagnosis was given) came from Spiennes. Thus, it is not clear whether it was collected from the 'Craie phosphatée' or from the underlying 'Craie de Spiennes' (Spiennes Chalk Formation in current terms), of late Campanian age. It seems this possibility has not yet been considered (Fig. 2); if so, it may link the type to a recent record of *P. giganteus* from the upper Campanian of Champagne (France) by Bardet et al. (1997);
- *Halisaurus ortliebi* (holotype: IRScNB R34; see Dollo, 1889c; Lingham-Soliar, 1996), from near Mesvin; and
- *Mosasaurus lemonnieri* (holotype: IRScNB R28 (ex 1470); see Dollo, 1889c, 1892; Lingham-Soliar, 2000), from Mesvin. Mulder et al. (2004) have recently suggested that this 'species' might in fact be nothing more than the juveniles of *M. hoffmanni*. In addition, Caldwell et al. (2004) observed that one of the specimens referred to *M. lemonnieri* (IRScNB 3211) showed close similarities to members of the New Zealand genus *Moanasaurus*. These data clearly demonstrate that the status of *M. lemonnieri* and its relationship to other mosasaurines is far from clear and in need of further studies.

As far as ammonite faunas from this unit are concerned, it is interesting to note their 'mixed' character (Kennedy, 1993). Mostly preserved as phosphatic internal moulds, in old collections material is merely labelled 'Poudingue de la Malogne' (= top of 'Craie phosphatée') and comprises both late Campanian

species such as *Nostoceras* cf. *hyatti* and *Solenoceras* sp., and early Maastrichtian ones (*Hauericeras* cf. *sulcatum*). The genus *Solenoceras* is an important element since it is restricted to the 'middle' and upper Campanian of the Western Interior, occurring in the zones of *Baculites gregoryensis* and *B. scotti* (North and South Dakota), *Didymoceras stevensoni* (Colorado, Wyoming, Montana), *Didymoceras nebrascense* (Montana, South Dakota, Colorado), *Exiteloceras jenneyi* (Montana, South Dakota, Colorado), *B. compressus* (Colorado), as well as *Baculites cuneatus* and *B. reesidei* (Colorado, Wyoming, Montana). *Solenoceras* has also been recorded from Arkansas, Texas, Tennessee, Delaware, Mississippi, Georgia, California, ?Spain, Italy, northern Germany (Jagt & Neumann, in prep.), Angola, ?Nigeria, ?Israel, ?Egypt and Japan (Cobban & Kennedy, 1991a, 1994a; Cobban et al., 1992; Kennedy & Cobban, 1993c, 1994b, 1997; Kennedy et al., 2000a, c). However, Ifrim et al. (2004) have recently described species of *Solenoceras* from the Maastrichtian of Mexico.

The 'Craie phosphatée' proper yielded *Pachydiscus* cf. *neubergicus*, *Diplomoceras cylindraceum*, *Baculites knorrrianus*, *B. baculus* and *Hoploscaphites constrictus*. Of these, *B. baculus* is of particular importance in allowing direct correlation with the *Baculites baculus* Zone of the Western Interior (Kennedy, 1993) (see Tables 1, 2).

On the basis of brachiopod assemblages, Simon (1998) was able to provide the most detailed picture of the stratigraphy of the 'Craie phosphatée' yet. The bulk of the unit is correlatable with the *acutirostris-spinosa* Zone (see Table 5), while sediment fill of basal cavities within the 'Poudingue de Malogne' hardground has yielded taxa typical of the younger *subtilis-pulchellus* Zone, and those that occur at a higher level within that same hardground are markers of the late Maastrichtian *stevensis-chitoniformis* Zone. Material contained in old collections thus needs to be interpreted with great caution.

Table 5. Micromorphic brachiopod zonation of the Upper Campanian-Maastrichtian of NW Europe (after Surlyk, 1984; Johansen & Surlyk, 1990). On brachiopod evidence, Simon (1998) favoured the following age assignments; 1 - 'Craie phosphatée de Ciply' proper; 2 - sediment fill from basal cavities of overlying 'Poudingue de Malogne' hardground; 3 - sediment fill of cavities higher up-section within that hardground.

Upper Maastrichtian	<i>stevensis-chitoniformis</i> Zone³ <i>humboldtii-stevensis</i> Zone <i>semiglobularis-humboldtii</i> Zone <i>tenuicostata-semiglobularis</i> Zone <i>pulchellus-tenuicostata</i> Zone <i>pulchellus-pulchellus</i> Zone
Lower Maastrichtian	<i>subtilis-pulchellus</i> Zone² <i>spinosa-subtilis</i> Zone <i>acutirostris-spinosa</i> Zone¹ <i>jasmundi-acutirostris</i> Zone
Upper Campanian (<i>pars</i>)	<i>longicollis-jasmundi</i> Zone <i>tenuicostata-longicollis</i> Zone

Conclusions

Newly collected mosasaur material from the Maastrichtian type area has the advantage that it is stratigraphically well documented, and thus can be used to refine ranges of taxa and consider taxonomic issues in much more depth. Unfortunately, new mosasaur discoveries from the Ciply-Mesvin area cannot be expected, since quarrying activities have come to a halt. The crude cephalopod-based correlation between the Liège-Limburg and Mons basins on the one hand, and North America on the other, compiled primarily from literature data, needs to be refined as well. Perhaps the renewed interest in belemnite fossils from the Atlantic Seaboard may yield additional data to tie these in with the highly detailed zonation applied to NW European sections. There appears to be consensus over environmental perturbations, whatever triggered them, across the lower/upper Campanian boundary, both in North America and Europe, seen in turnover in both mosasaur and ammonoid faunas. Campanian-Maastrichtian ammonites from the Western Interior, often held to be highly endemic, have been yielding elements that occur on both sides of the Atlantic at some levels within these stages.

It is not surprising to see that fully nektonic animals, such as mosasaurs, could cross the Late Cretaceous Atlantic Ocean with ease, and that many taxa are in fact conspecific. Detailed biozonations and age assignments are needed for type and additional referred material to allow migratory patterns to be reconstructed more faithfully. The present paper can be no more than a first step in that direction.

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