

El Niño effects on rainfall in South America: comparison with rainfalls in india and other parts of the world

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Abstract. As a finer classification of El Niños, ENSOW were defined as years when El Niño (EN) existed on the Peru coast, Southern Oscillation Index SOI (Tahiti minus Darwin pressure) was negative (SO), and Pacific SST anomalies were positive (W). Further, Unambiguous ENSOW were defined as years when SO and W occurred in the middle of the calendar year, while Ambiguous ENSOW were defined as years when SO and W occurred in the earlier or later part of the calendar year (not in the middle). In contrast with India and some other regions where Unambiguous ENSOW were associated predominantly with droughts, in the case of South America, the association was mixed. In Chile on the western coast and Uruguay etc. on the eastern coast, the major effect was of excessive rains. In Argentina and central Brazil, the effects were unclear. In Amazon, the effects were not at all uniform, and were different (droughts or excess rains) or even absent in regions only a few hundred kilometers away from each other. Even in Peru-Ecuador, the effects were clear only in the coastal regions. In the interior and in the Andes, the effects were obscure. In NE Brazil, El Niños have been popularly known to be causing severe droughts. The fact is that during 1871–1998, there were 52 El Niño events, out of which 31 were associated with droughts in NE Brazil, while 21 had no association. The reason is that besides El Niños, another major factor affecting NE Brazil is the influx of moisture from the Atlantic. In some years, warmer Atlantic in conjunction with westward winds can bring moisture to NE Brazil, nullifying the drought effects of El Niños. A curious feature at almost all locations is the occurrence of extreme events (high floods or severe droughts) in some years, apparently without any El Niño or La Niña events. This possibility should always be borne in mind.

1 Introduction

Along the coast of Peru-Ecuador in South America (Niño 1–2 region), the average seasonal pattern (climatology) of sea surface temperature (SST) is annual, with a maximum temperature of $\sim 19^{\circ}\text{C}$ in February and a minimum temperature of $\sim 15.5^{\circ}\text{C}$ in October (Deser and Wallace, 1987). There is also an ocean current called Peru or Humboldt current. El Niño (means The Child, in Spanish) is defined as a warming of this ocean current, so called because it generally develops near Christmas, the birth of Jesus Christ. In some years, the current may extend southward along the coast of Peru to latitude 12°S , killing plankton and fish in the coastal waters. Quinn et al. (1978, 1987) determined the occurrence of El Niño events on the basis of the disruption of fishery, hydrological data, sea-surface temperature and rainfall along and near the Peru-Ecuador coast, defining El Niño intensities based on the positive sea-surface temperature anomalies along the coast, namely, Strong, in excess of 3°C ; Moderate, $2.0\text{--}3.5^{\circ}\text{C}$; Weak, $1.0\text{--}2.5^{\circ}\text{C}$. (This is only the local feature, but the term El Niño is now used to refer a much larger event, hence the term ENSO).

It is popularly believed that El Niños are associated with droughts in many parts of the globe (Ropelewski and Halpert, 1987). However, not all El Niños are equally effective. For relationship with El Niños, the listing given by Quinn et al. (1978, 1987) is often used. Kane (1997a, b; 1998a, b, c) noticed that not all El Niños in this list were associated with droughts in India and other locations. From the 25 El Niño events selected from the Quinn et al. (1978, 1987) list by Rasmusson and Carpenter (1983), only 9 were associated with severe droughts in India, 11 with only slightly below normal rainfall but 5 with excess rainfall. Trenberth (1993) refers to different “flavours” of El Niños. However, these have not been identified so far for every individual El Niño event in the past. As a hindsight study, Kane (1997a, b; 1998a, b, c) attempted a finer classification of El Niños, in

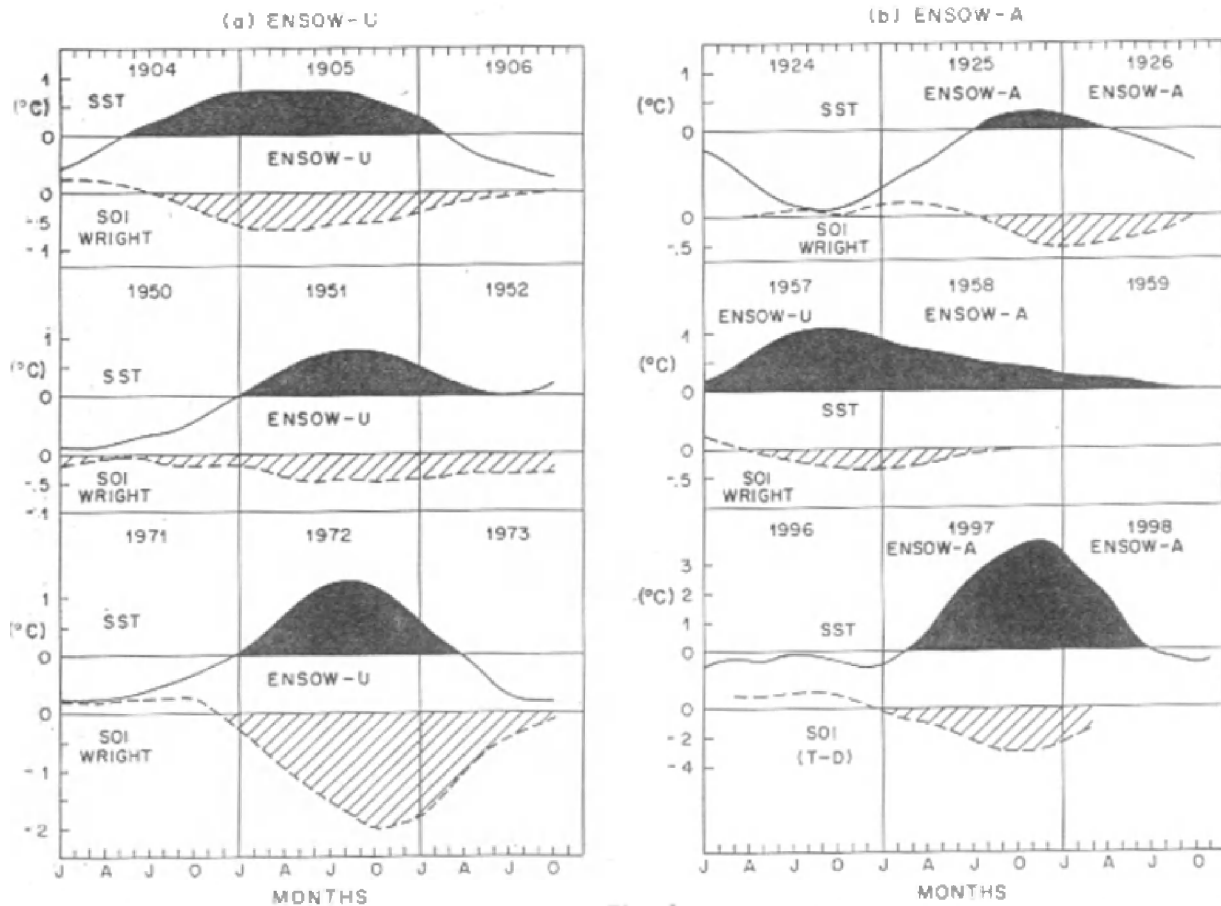


Fig. 1. Plots of 12-monthly running means of SST and SOI for selected El Niño events for 3 consecutive years, where the El Niño event is in the middle year. **(a)** Unambiguous ENSOW (ENSOW-U) 1905, 1951, 1972, for which SST maxima and SOI minima are in the middle of the calendar year, and **(b)** Ambiguous ENSOW (ENSOW-A) 1925, 1958, 1997, for which SST maxima and SOI minima are not in the middle of the calendar year.

which Unambiguous ENSOW type events (ENSOW-U, discussed later on) were found to be overwhelmingly associated with droughts in India, Sri Lanka, South-eastern Australia and some other parts of the globe.

2 El Niño characteristics and their implications for rainfall

2.1 Finer classification of El Niño events

The term ENSO is used nowadays for the general phenomenon of the El Niño/Southern Oscillation. However, in our classification here, its components EN, SO are used in their *literal* sense. Thus, every year was examined to check whether it had an El Niño EN (as listed in Quinn et al., 1978, 1987), and/or Southern Oscillation SOI minimum (SO) and/or warm (W) or cold (C) equatorial eastern Pacific sea surface temperature (SST) anomalies. For SO and SST, 12-month running means were used. Several years were ENSOW, i.e., El Niño (EN) existed near Peru coast, SOI had minima (SO), and Pacific SST was warmer (W).

These were subdivided into two groups viz., *Unambiguous ENSOW* (ENSOW-U) where El Niño existed near Peru coast and the SOI minima and SST maxima in the 12-month running means were in the middle of the calendar year (May–Aug.) and, *Ambiguous ENSOW* (ENSOW-A) where El Niño existed near Peru coast, but the SOI minima and SST maxima were in the early or later part of the calendar year, not in the middle. Besides these, there were other El Niño years of the type ENSO (El Niño existed, SOI minima existed, but SST in the Pacific, away from the Peru-Ecuador coast, was neither warmer nor colder, just normal), ENW, ENC, i.e., El Niño existed, SOI minima did not exist, but SST was warm W, or cold C (before or after the El Niño) or just EN (i.e., El Niño existed only near Peru coast, no SOI minima or central and eastern Pacific SST maxima or minima). Some other years did not have an El Niño and were of the types SOW (SOI minima existed and Pacific waters were warm, but no El Niño was reported by Quinn et al.), SOC, SO, W, and C, where the last category C contains all Anti-El Niños, i.e., La Niñas. The types ENC or SOC look contradictory. Actually, in these years, EN and/or SO and C did not occur simultaneously. EN or SO occurred in a part of the year and was

followed or preceded by a cold Pacific SST, i.e., C event. Years not falling into any of these categories were termed as *Non-events*. We have this classification ready for all years from 1871 onwards. Appendix 1 gives the classification for 1900 onwards.

Figure 1 shows the plots of the 12-monthly running averages of Pacific SST (roughly, the Niño 3.5 region) and the Southern Oscillation Index SOI for a few selected intervals of 3 consecutive years, the middle year being an El Niño year. In Fig. 1a, SOI minima and SST maxima (12 monthly running means) are in the middle of the calendar year of the El Niño year (middle part) and these are ENSOW-U (Unambiguous ENSOW) years. For such years, the AISM (All India Summer Monsoon) normalized rainfalls had negative deviations. Figure 1b shows similar plots for some years of the ENSOW-A (Ambiguous ENSOW) type. Here, the SOI minima and SST maxima of the El Niño year (middle part) are not in the middle of the calendar year. For such years, the AISM deviations generally indicated almost normal or sometimes even excess rainfall (see below). During 1871–1990, there were 47 El Niño years. Among these, 16 could be designated as Unambiguous ENSOW, 15 as Ambiguous ENSOW and 16 as Other EN.

2.2 Rainfall effects in All India Summer Monsoon

The various El Niño years and the corresponding normalized AISMR (All India Summer Monsoon Rainfall) anomalies were as in Table 1. For AISMR, the mean and the standard deviation (σ) were calculated and the value of every year was expressed as *normalized*, i.e., deviation from the mean, divided by the standard deviation. These normalized deviations AISMR(σ) were categorised as normal (no symbol) if these were within $\pm 0.5\sigma$; *f* (moderate floods) and *d* (moderate droughts) if within 0.5 to 1.0 σ and -0.5 to -1.0σ respectively; *F* (severe floods) if above 1.0 σ and *D* (severe droughts) if below -1.0σ .

As can be seen, the Unambiguous ENSOW are all associated with droughts (*d*, *D*), while the Ambiguous ENSOW are mostly associated with normal rainfall. Here, the years shown in *italics* are II year events (1958 in 1957–1958, etc.) and it seems that these II years do not show droughts. The Other EN are associated with droughts, floods and normal rainfall.

2.3 Rainfall effects in other regions

Overwhelming association of ENSOW-U with droughts or floods was observed for many other regions. The Tables will not be shown here. As a measure of the association, the fraction of positive rainfall deviations in any category of years could be considered. For example, for ENSOW-U category of 16 events, All India Summer Monsoon Rainfall had all negative deviations and the number of positive deviations was zero. Hence, the fraction of Positive/Total was 0.0. At other locations where there were floods, the fraction of positives would be near 1.0. Figure 2 shows the plot of

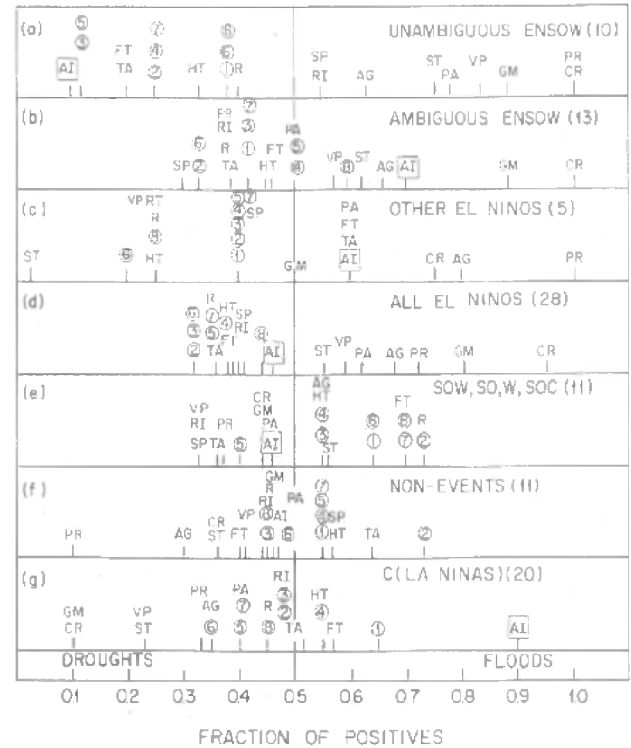


Fig. 2. Plot of positive fractions for various rainfalls for different types of events. The central vertical line is for a fraction of 0.5, i.e., equal number of positive and negative rainfall deviations, implying poor relationship. Extreme left, droughts; extreme right, floods.

these fractions of positive deviations. It refers to regions in South America, but similar plots are available for other parts of the globe. The vertical line in the middle is for a fraction 0.5, implying equal number of positive and negative rainfall deviations and hence, a *poor relationship*. Regions in the left extreme have very low fractions (0.0, 0.1, 0.2) of positive rainfall deviations, implying a good relationship with droughts. Regions in the right extreme have large fractions (0.8, 0.9, 1.0) of positive rainfall deviations, implying a good relationship with floods. Thus,

1. For Unambiguous ENSOW-U, a very good association is with droughts in India; Sri Lanka (southwest monsoon); Tasmania (Southeast Australia); Singapore, Brunei, Indonesia (Southeast Asia); Ceara, Rio grande do Norte, Paraiba, Pernambuco (Northeast Brazil). Also, there is good association with floods in Chile, Peru, Core Region, Gulf of Mexico and South Brazil.
2. For Ambiguous ENSOW-A and Other El Niños, fractions are near 0.5, implying poor relationships. But for Peru and Core regions, relationship with floods is good even for these events.
3. For other types of years not involving El Niños, relationship is poor.
4. For Non-events also, relationship is poor.

Table 1. All India Summer Monsoon Rainfall (AISMR) during El Niño years of three types, Unambiguous ENSOW (ENSOW-U), Ambiguous ENSOW (ENSOW-A), and Other EN.

Unambiguous		Ambiguous		Other EN	
ENSOW-U		ENSOW-A			
Year	AISMR(σ)	Year	AISMR(σ)	Year	AISMR(σ)
1877	−3.0 <i>D</i>	1878	+1.5 <i>F</i>	1871	−0.1
1888	−0.5 <i>d</i>	1914	+0.6 <i>f</i>	1873	−1.3 <i>D</i>
1896	−0.3	1919	+0.4	1874	+1.5 <i>F</i>
1899	−2.7 <i>D</i>	1923	−0.4	1880	−0.4
1902	−0.7 <i>d</i>	1925	−0.6 <i>d</i>	1884	+1.0 <i>F</i>
1905	−1.6 <i>D</i>	1926	+0.6 <i>f</i>	1887	+0.5 <i>f</i>
1911	−1.4 <i>D</i>	1931	+0.3	1889	+1.0 <i>f</i>
1918	−2.4 <i>D</i>	1940	−0.0	1891	−0.8 <i>d</i>
1930	−0.6 <i>d</i>	1948	+0.2	1897	+0.4
1941	−1.5 <i>D</i>	1953	+0.8 <i>f</i>	1900	+0.4
1951	−1.4 <i>D</i>	1958	+0.4	1907	−1.0 <i>d</i>
1957	−0.8 <i>d</i>	1963	+0.0	1912	−0.6 <i>d</i>
1965	−1.7 <i>D</i>	1969	−0.3	1917	+1.9 <i>F</i>
1972	−2.4 <i>D</i>	1976	+0.0	1932	−0.6 <i>d</i>
1982	−1.4 <i>D</i>	1983	+0.2	1939	−0.7 <i>d</i>
1987	−1.9 <i>D</i>			1943	+0.2

5. For C type events (La Niñas), a very good association is with floods in India; and droughts in Chile, Core region, Gulf of Mexico.

2.4 Can ENSOW-U be predicted?

If the relationship with ENSOW-U is so good, it could be used for predicting extreme rainfall events, if the ENSOW-U can be predicted with some antecedence. Unfortunately, this does not seem to be possible at present with any reasonable accuracy. Take the example of the 1997–1998 El Niño event. In the last two decades, many models have been developed to predict the evolution of El Niño as well as the Southern Oscillation. However, instead of the predicted turning warm only in late 1997, the El Niño started early (in March 1997). As remarked by Trenberth (1998), in digesting these predictions in real time, it was not possible to make a forecast until about April 1997 when the SST warming became clearly evident in the eastern tropical Pacific. Particularly deceptive was the forecast of the Cane-Zebiak model. This model had successfully forecast the 1986–1987 event (Cane et al., 1986; Zebiak and Cane, 1987). Hence their prediction of “normal by June 1997, warming thereafter” substantially inhibited the issuing of a forecast for March 1997.

In the Executive Summary for the La Niña Summit Report of a Workshop held in Boulder, Colorado from 15–17 July, 1998, Glantz (1998) mentions the key points about prediction of the El Niño 1997 as follows: (a) The performance in forecasting the onset of the 1997–1998 El Niño was largely *mediocre*. (b) Dynamical models as yet do not outperform the statistical ones, with respect to forecasting El Niño. (c) The communication between forecasters and users

still leaves something to be desired. It appears that neither really understands how the other thinks and what the other does or does not understand.

2.5 SST characteristics in the Pacific

The finer classification mentioned above is obtained by visual examination of the plots and no physical basis is offered, except perhaps that the circulation patterns affecting rainfall operate best when these maximise in the northern summer months. However, some relationship with SST patterns in the Pacific may be involved. Fu et al. (1986) identified two distinct patterns in Pacific SST, one operating in some El Niño years and the other in other El Niño years. Ward et al. (1994) separated years when AISM and Sahel had droughts or excess rainfalls, during 1904–1990. From these, years having El Niños or La Niñas were selected for analysis. The symbol FIT means years where rainfalls were as expected. Four categories were obtained:

- El Niño Type I (FIT years): El Niño years when India and Sahel had droughts (1904, 11, 13, 25, 41, 51, 57, 63, 66, 72, 76, 77, 82, 87).
- El Niño Type II (NOT FIT years): El Niño years when India and Sahel had excess rains (1912, 15, 18, 19, 23, 26, 30, 31, 32, 40, 47, 52, 53, 58, 65, 69, 83, 86, 90).
- La Niña Type I (FIT years): La Niña years when India and Sahel had excess rains (1906, 09, 16, 24, 50, 54, 64, 70, 74, 75, 88).

- La Niña Type II (NOT FIT years): La Niña years when India and Sahel had droughts (1908, 17, 20, 21, 28, 38, 39, 42, 43, 49, 55, 71, 73, 89).

Type I were the FIT years (rainfall as expected), while Type II were the NOT FIT years. The average SST longitude profiles for these four groups for JAS months are shown in Fig. 3. The dashed plot shows the average pattern (climatology) for 1951–1980. The full lines show the average SST for El Niño Type I (EN I) in the first plot, El Niño Type II (EN II) in the second plot, La Niña Type I (LN I) in the third plot and La Niña Type II (LN II) in the fourth plot, with the climatology (dashed line) superimposed on each plot. When the climatology is subtracted, the residuals (SST anomalies) are as shown in the lower part of Fig. 3. EN I has positive SST anomalies in the longitudes 130–180 W while EN II has positive SST anomalies near 170 E. LN I has positive SST anomalies near 150–160 E and large negative SST anomalies in 130–180 W, while LN II has small positive SST anomalies near 120 E and small negative anomalies in 130–180 W.

To test whether the 1997–1998 El Niño had a pattern like EN I or EN II, plots of SST anomalies were made at different longitudes 10° apart for JAS months for different latitude belts 4° apart (30 N–26 N, 26 N–22 N, ..., 2 N–2 S, ..., 22 S–26 S, 26 S–30 S). Substantial anomalies were seen in the 6 N–10 S belt, in a broad longitude around 120 W. Thus, the anomaly pattern resembled EN I pattern. As such, the 1997–1998 El Niño should be of the FIT type, suitable as a “tropic-wide oscillation of boreal summer rainfall” mentioned by Ward et al. (1994) and should have caused severe droughts in India, southeast Australia etc. However, this expectation was not fulfilled for the 1997–1998 event, leaving in doubt whether the longitude distribution in Pacific SST is a deciding factor.

3 The El Niño of 1997–1998

In March 1997, an El Niño developed and proved to be the strongest in the 20th century. Kane (1999a, b) has summarized its effects in different parts of the globe. For S. America, Southern Peru had droughts while eastern, northern Peru and Ecuador had floods. In Bolivia, parts adjacent to Brazil had floods while parts adjacent to Peru and Chile had droughts. In northeast Brazil, the observed rainfall for March–May 1997 (main rainy season) was 15–30% below normal. Thus, El Niño effects could be claimed. In 1998, there was a very severe drought in northeast Brazil. So again, El Niño effect could be claimed, though the El Niño was on the decline during March–May 1998. There was intense, prolonged drought in Amazonia and Roraima (both in northern Brazil). By 1998 end, northeast Brazil had some rainfall and came out of a prolonged severe drought. In South Brazil, Paraguay, Uruguay, Northeast Argentina and Chile, there were heavy rains in 1997. In Chile, deserts bloomed after 40 years. Thus, El Niño effect was fully visible. Northern Uruguay was wetter in April–July of 1998. September and October of 1998 were above normal in the southern part

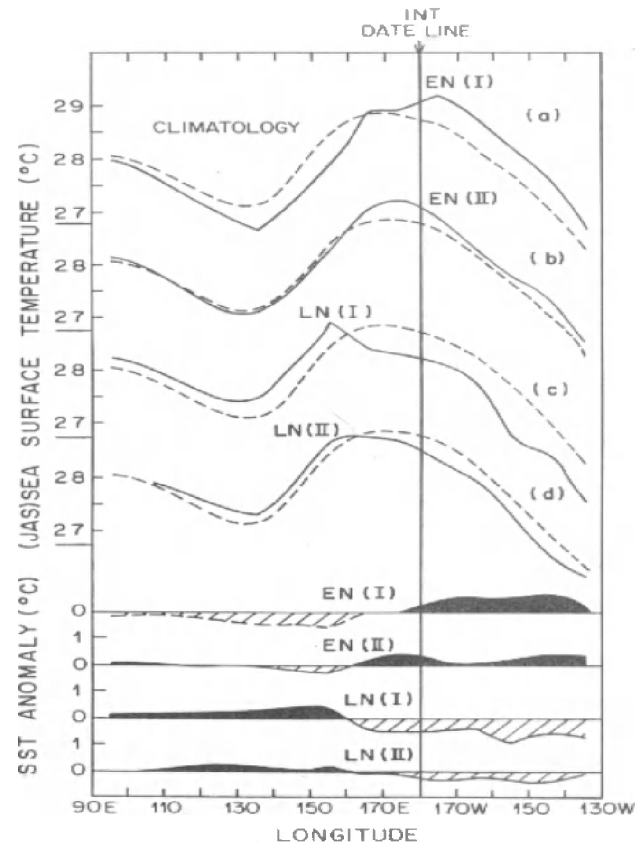


Fig. 3. Plot of July-August-September average SST (full lines) versus longitude in the Pacific for El Niño years of (a) the FIT type EN(I), which favoured droughts in Sahel and India, (b) NOT FIT type EN(II), which did not favour droughts, and for La Niña years of (c) the FIT type LN(I), which favoured excess rainfall in Sahel and India, (d) NOT FIT type LN(II), which did not favour excess rainfalls (Ward et al., 1994). The dashed line superposed on each plot is the climatology. Deviations (SST anomalies) from the climatology are plotted in the bottom part.

of Brazil. However, November and December were below normal.

4 Conclusions

For southern S. America, Kane (2002) examined data for various locations and concluded that whereas ENSOW-U events had very good association with floods, some ENSOW-A also had a good association. La Niñas were mostly associated with droughts but some were associated with floods. Some non-events (neither El Niños nor La Niñas) were also associated with extreme rainfalls. These facts render dubious the relationships of El Niño and La Niña with rainfalls. The Amazon river basin and its vicinity (15° N–20° S, 30°–80° W) indicated that the rainfalls were highly variable both from year to year and from region to region. Correlations with even nearby regions hardly exceeded 0.50, though correlations were better (up to 0.70) in the regions near the

Table A1.

Unambiguous ENSOW	Ambiguous ENSOW	Other El Niños	SOW etc.	Non- events	C La Niña
M 1902 RKP	M 1914 R	S 1912 ENSO	1904 SOW KP	1901	1903 K*P*
M 1905 R P	M 1919 II	M 1929 EN	1913 SOW KP	1915	1906 K*P*
S 1911 RKP	M 1923 RK	S 1932 EN	1944 SOW	1937	1908 K*P*
S 1918 I RKP	S 1925 I RKP	M 1939 EN	1977 SOW	1945	1909
M 1930 I RKP	S 1926 II	M 1943 EN	1979 SOW	1947	1910
S 1941 II R P	M 1931 II K*	M 1907 ENC	1920 W K*	1952	1916 K*P*
M 1951 RKP	S 1940 I P	S 1917 ENC	1968 W P	1966	1921
S 1957 I RKP	W 1948 P	S 1927 ENC	1986 W K	1978	1922
M 1965 RKP	M 1953 RK	S 1973 ENC	1959 SO	1980	1924 K*
S 1972 I RKP	S 1958 II	9 events	1974 SO	1981	1928 K*
S 1982 KP	W 1963 KP		1935 SOC	1984	1933
M 1987	W 1969 RKP		1936 SOC	1985	1934
12 events	M 1976 RKP		1946 SOC P*	1989	1938 K*
	S 1983 II P*		1949 SOC K*	1990	1942 K*
	14 events		14 events	14 events	1950
					1954 K*
					1955 P*
					1956
					1960
					1961
					1962
					1964 K*P*
					1967 P*
					1970 K*P*
					1971
					1975 K*
					1988
					27 events

eastern coast of Brazil. Moderate relationship with ENSO indices was obtained for the Amazon river basin and the regions to its north, and for NE Brazil, while moderate relationship with South Atlantic SST was obtained for NE Brazil and the region immediately to its west. All other relationships (with 30 mb wind, North Atlantic Oscillation Index, etc.) were obscure.

Appendix

Distribution of the years 1900–1990 in various categories. Symbols S (strong), M (Moderate), W (Weak) indicate the strengths of the El Niños involved. I and II indicate first and second years of double events (1957–1958 etc.). R, K, P indicate that these were selected as warm events by Rasmusson and Carpenter (1983), Kiladis and Diaz (1989) and Mooley and Paolino (1989). K* and P* indicate that these were selected as cold events by Kiladis and Diaz (1989) and Mooley and Paolino (1989).

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