II. The Principal Constituent of the Tides in the English and Irish Channels.

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1. A chart of co-tidal and co-range lines for the North Sea* was prepared at the Tidal Institute in the year 1923, and the methods then used have now been further developed and improved and applied to the English Channel, the Irish Sea, and their approaches.†

The methods used depend largely upon the known dynamical equations connecting the currents with the gradients of the elevations. If any assumption be made as to the values of the range of tide and the relative phases of the elevation and currents, then we can deduce from this information at one station, not only the directions of the co-tidal and co-range lines at that point, but also the degree of separation of the lines for any given unit of phase or range. This criterion, applied to speculative charts hitherto published, suffices at once either to verify or to condemn the charts. Again, without any assumptions at all, from the gradients at a number of stations on a line of small curvature, starting at a point at which the elevation is known, the elevations at all points along the line can be computed by simple methods of numerical integration.

For the North Sea there were available the results of harmonic analyses of tidal currents, observed over periods of a fortnight, at 15 stations, but for the English and Irish Channels much inferior data have had to be used, and most of this information was obtained many years ago. The Hydrographer ordered special observations to be taken across three lines at the entrances to the Channels, and the results were in approximate agreement with the older observations at such points where comparisons could be made. As most of the observations probably extended over only 24 hours, it became necessary to criticise the data, and to reject anomalies wherever such existed. Since the methods of integration tend to diminish the importance of individual errors the results of integration have fitted in with the coastal observations so well as to

^{*} PROUDMAN and Doodson, 'Phil. Trans.,' A, vol. 224, p. 185 (1923).

[†] The preparation of the new tidal charts was ordered by the Hydrographer of the Navy, and the cost of investigation and preparation defrayed by the Hydrographic Department of the Admiralty. The new tidal charts have been combined with the former North Sea Tidal Chart, and the combined chart is on sale to the general public as Admiralty Chart No. 301.

warrant the confident construction of co-tidal and co-range lines for these Channels. It will be found that the charts differ very considerably from the speculations previously published in that the latter are dominated by the idea of a tidal wave travelling more quickly in the middle of the Channel than at the sides, so that the charts show great curvature of the co-tidal lines inward to the Channel. This assumption is entirely wrong, and the new charts, on the contrary, show that many of the lines are actually curved (slightly) in the opposite direction.

2. Definitions of Tidal Elevations.

We shall deal only with the principal lunar constituent M_2 , and the tidal elevation, at any point, due to this constituent will be expressed as

$$\zeta = H \cos (\sigma t - g - 30^{\circ}) = H \cos (\sigma t - d) = \zeta_1 \cos \sigma t + \zeta_2 \sin \sigma t$$

where σ is the speed of the harmonic motion (i.e., $2\pi/\text{period}$), t is the time, and H, g are functions of position required for the co-range and co-tidal charts respectively. The origin of the time chosen for this work has been the time of high water (of M_2) at Dover; this choice was made because most of the data concerning tidal currents are referred to the time of high water at Dover. We define g to be the lag in phase of the elevation behind that of the corresponding equilibrium constituent on the meridian of Greenwich, and at Dover the value of g has been taken as 330°. If t is expressed in hours we have $\sigma = 28.984$ °.

3. Definitions of Tidal Currents.

As a general rule tidal currents are rotatory and the vector-diagram of velocity is an ellipse. The velocity-components in the directions of the major and minor axes of this ellipse are expressed respectively by

and
$$u_0 \cos \left(\sigma t - g' - 30^\circ\right)$$

$$v_0 \sin \left(\sigma t - g' - 30^\circ\right),$$

i.e., by $u_0 \cos(\sigma t - d')$ and $v_0 \sin(\sigma t - d')$. It is necessary to specify the direction (6) in which the maximum current is $+u_0$, and this angle will be measured from the east round by the north. The positive minor axis, by a convention, is taken as 90° in advance of the positive major axis, and thus v_0 can be either positive or negative: in the former case the currents rotate in the positive sense, mathematically.

Taking Cartesian axes Ox, Oy on the mean surface of the sea, to the east and to the north respectively, then we define the corresponding current components as

$$u = u_1 \cos \sigma t + u_2 \sin \sigma t,$$

$$v = v_1 \cos \sigma t + v_2 \sin \sigma t.$$

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4. Currents in the Open Sea.

The principal source of information of currents is the Admiralty publication, "Tides and Tidal Streams."* Details of streams, between 6 hours before to 6 hours after the time of high water at Dover, are given for a large number of stations, and these are tabulated, with values of current components for M_2 , and the elevation-gradients, in Table I. The notation S 263 means that the station is the third station on p. 26 of the publication referred to. Figs. 1 and 2 indicate the situations of the stations.

TABLE I.—Current Data.

Station.	Latitude.	Longitude.	Depth.	₹£ ₁ ,	v ₁ .	\mathcal{U}_2	₽3.	∂ζ₁/∂r.	∂ζ./∂y.	∂ζ,/∂x.	∂ζ₁/∂y.
S. 222 223 224 231 232 234 235 234 235 236 251 252 253 254 256 261 262 263 271 272 273 341 342 343 361 362 363 364 365 366 367 371 372 373 391	49 16N 49 52 49 02 49 40 49 15 49 02 49 25 49 25 49 45 50 00 49 35 49 49 49 30 49 29 49 49 49 49 49 30 49 29 49 49 49 50 15 50 11 49 50 49 35 49 35 49 35 49 45 50 50 15 50 11 50 50 54 50 55 50 54 50 57 50 54 51 03 51 09 51 13 51 16 49 45	8 11W 7 58 6 46 6 28 6 03 5 39 5 20 4 35 3 45 5 41 5 28 5 03 4 15 3 19 4 45 2 20 1 10 0 00 1 17 2 20 1 16 1 00 0 30 0 27E 1 00 1 16 1 27 1 22 1 28 1 36 1 29 6 35 8 1 30 8 1	fathoms 70 70 70 70 57 61 61 55 45 40 54 45 52 48 50 56 35 39 26 39 91 25 17 14 7 15 15 20 15 54 54	-15 -10 -25 -25 -30 -10 -51 0 -38 -30 -35 -42 -20 -21 -20 0 25 0 -5 45 -25 0 20 25 25 30 20 60 30 45 40 40 35 -15	5 10 - 5 - 25 20 - 17 0 - 10 - 5 20 - 6 10 - 10 - 15 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	30 25 5 0 10 40 32 - 50 - 67 10 15 54 15 - 15 27 - 80 - 125 - 80 - 70 - 175 - 80 - 105 - 90 - 105 - 80 - 105 - 30 - 40 - 50 - 10 - 1	25 5 40 25 20 25 10 25 18 30 - 15 12 - 15 - 10 - 35 - 10 - 35 - 10 - 35 - 20 - 35 - 35	- 36 - 25 - 12 - 5 - 40 - 34 - 59 - 72 - 89 - 18 - 4 - 79 - 9 - 11 - 55 - 117 - 179 - 123 - 150 - 111 - 193 - 167 - 150 - 128 - 92 - 15 - 5 - 5 - 12 - 15 - 17 -	- 19 4 - 29 - 8 6 - 26 48 19 69 20 2 23 - 20 44 31 45 - 7 - 26 - 150 - 117 - 14 48 - 3 - 55 - 74 - 84 - 127 - 196 - 145 - 180 - 317 - 278 - 4	4 - 9 9 - 8 - 20 10 - 69 - 9 - 67 - 32 - 23 - 47 - 45 - 19 - 29 - 45 - 68 - 25 - 43 - 288 125 - 72 - 33 115 - 49 - 62 - 64 100 122 - 75 - 99 125 - 7 - 5	- 29 - 15 - 8 - 48 - 60 - 57 - 62 - 18 - 9 - 72 - 5 - 41 - 113 - 125 - 96 - 109 - 64 - 112 - 158 - 121 - 7 - 19 - 18 - 16 - 12 - 31 - 12 - 143 - 202 - 17

* "Tides and Tidal Streams," H.M. Stationery Office, 1st Ed. (1909).

Table I.—(continued).

Station.	Latitude.	Longitude.	Depth.	u ₁ .	v_1 .	u.	v _g ,	$\partial \zeta_1/\partial x$.	$\partial \zeta_1/\partial y$.	$\partial \zeta_1/\partial x$.	∂ζ₂/∂y.
S. 392 393 401 402 403 421 422 491 493 501 502 541 542 543 551 572 1611 1612 1613 1614 1615 1621 1812 1813 A ₁ B ₂ C ₁ C ₃ A B C D E F G 1951 1952 1953 1954 1955 1956 1961 1962	\$\frac{49}{52N}\$ \$\frac{50}{50}\$ \$\frac{35}{60}\$ \$\frac{49}{56}\$ \$\frac{49}{50}\$ \$\frac{50}{49}\$ \$\frac{50}{50}\$ \$\frac{31}{50}\$ \$\frac{37\frac{1}{2}}{50}\$ \$\frac{42\frac{1}{2}}{50}\$ \$\frac{42\frac{1}{2}}{50}\$ \$\frac{44}{51}\$ \$\frac{46}{51}\$ \$\frac{30}{51}\$ \$\frac{31}{52}\$ \$\frac{51}{30}\$ \$\frac{31}{51}\$ \$\frac{30}{51}\$ \$\frac{31}{30}\$ \$\frac{51}{31}\$ \$\frac{30}{51}\$ \$\frac{31}{30}\$ \$\frac{51}{31}\$ \$\frac{30}{31}\$ \$\frac{51}{30}\$ \$\frac{30}{31}\$ \$\frac{31}{30}\$ \$\frac{30}{31}\$ \$\frac{51}{30}\$ \$\frac{30}{31}\$ \$\frac{31}{30}\$ \$\frac{30}{31}\$ \$\frac{30}{3	0 35W 6 26 04 6 06 12 5 52 48 2 29½ 2 20 2 34 2 25 0 54 1 00 10 18 8 11 8 6 39 4 25 5 52 1 04 11 00 10 18 8 41 11 8 6 39 5 52 5 36 5 54 6 38 7 5 23 8 27 10 30 6 04 5 54 5 54 8 57 10 4 8 57 10 6 8 57 10 6 10 6 10 6 10 6 10 6 10 6 10 6 10 6	fathoms 48 45 40 40 45 45 40 20 13 27 10 15 35 16 10 100 55 45 30 25 30 20 15 58 49 58 80 90 70 60 70 7	$\begin{array}{c} -25 \\ -15 \\ -15 \\ -30 \\ -30 \\ -20 \\ -45 \\ -20 \\ -25 \\ -20 \\ -30 \\ -12 \\ 10 \\ -15 \\ 0 \\ -30 \\ -25 \\ -10 \\ 5 \\ -25 \\ -20 \\ -35 \\ -20 \\ -40 \\ -35 \\ -20 \\ -40 \\ -35 \\ -25 $	$\begin{array}{c} 10 \\ -5 \\ -10 \\ -5 \\ -10 \\ -10 \\ -5 \\ -10 \\ -10 \\ -5 \\ -10 \\ -60 \\ -60 \\ -60 \\ -60 \\ -60 \\ -60 \\ -60 \\ -60 \\ -20 \\ -20 \\ -20 \\ -10 \\ -20 \\ -15 \\ -20 \\ -15 \\ -20 \\ -65 \\ -50 \\ -45 \\ -50 \\ -45 \\ -50 \\ -45 \\ -25 \\ 35 \\ \end{array}$	10 10 10 10 10 10 10 10 10 10 10 10 10 1	35 20 35 35 40 20 40 40 40 35 40 20 40 40 35 40 20 35 5 30 30 30 40 20 30 30 20 25 10 -110 -120 -100 -100 -100 -100 -100 -	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	$\begin{array}{c} -22 \\ -12 \\ -33 \\ -16 \\ -22 \\ -6 \\ -40 \\ 85 \\ 29 \\ 116 \\ 11 \\ 56 \\ -40 \\ 13 \\ 49 \\ -40 \\ 13 \\ 49 \\ -40 \\ 13 \\ -20 \\ 54 \\ 75 \\ 76 \\ 120 \\ 108 \\ 3 \\ 10 \\ 8 \\ -28 \\ -14 \\ -15 \\ -2 \\ 25 \\ 148 \\ -104 \\ 98 \\ \end{array}$	3 1 18 4 1 - 7 20 - 55 37 - 59 - 49 - 43 - 14 26 - 13 - 9 - 49 - 49 - 72 - 108 - 145 - 85 - 75 - 11 - 31 - 9 - 14 - 33 - 123 - 123 - 127 - 104 - 96 - 89 - 89	$\begin{array}{c} 1\\ -19\\ -22\\ -29\\ -4\\ -12\\ 15\\ 52\\ 66\\ -44\\ -71\\ 212\\ 187\\ 107\\ 126\\ 84\\ 59\\ 22\\ -37\\ -53\\ -37\\ -53\\ -37\\ -53\\ -39\\ -46\\ -30\\ -45\\ -39\\ -18\\ -25\\ -29\\ -10\\ -11\\ -53\\ -37\\ -53\\ -26\\ -30\\ -18\\ -29\\ -19\\ -11\\ -53\\ -37\\ -34\\ -29\\ -37\\ -34\\ -37\\ -34\\ -37\\ -37\\ -37\\ -37\\ -37\\ -37\\ -37\\ -37$

The Hydrographer undertook to obtain information of currents at a number of places on three sectional lines:—

A-From Tusker Rock to St. David's Head.

B-From C. Clear to Land's End.

C-From Land's End to Ushant.

Deductions from these tabulations are also tabulated in Table I. Certain unpublished data for stations in the approaches (A to G) are also given in this table. The Hydrographic Department provided corrected data for stations 1951 to 1956.

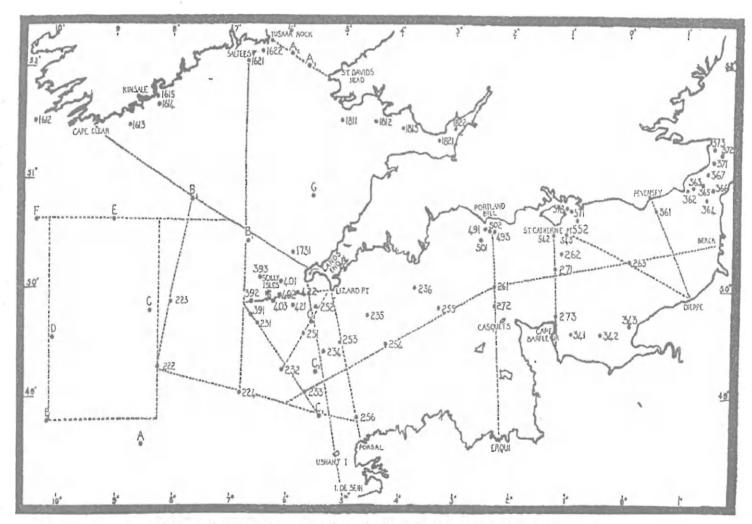


Fig. 1.—Current-stations and sectional lines in the English and Bristol Channels.

The velocity of current is recorded as for spring tides, and the appropriate reduction factors have been assumed to be the same as for the elevations at neighbouring constal stations (see § 9). The corresponding values of u, v (in directions E and N) were then computed, and u_1 , u_2 , v_1 , v_2 , were deduced as follows:

Since

$$u(t) = u = u_1 \cos \sigma t + u_1 \sin \sigma t$$

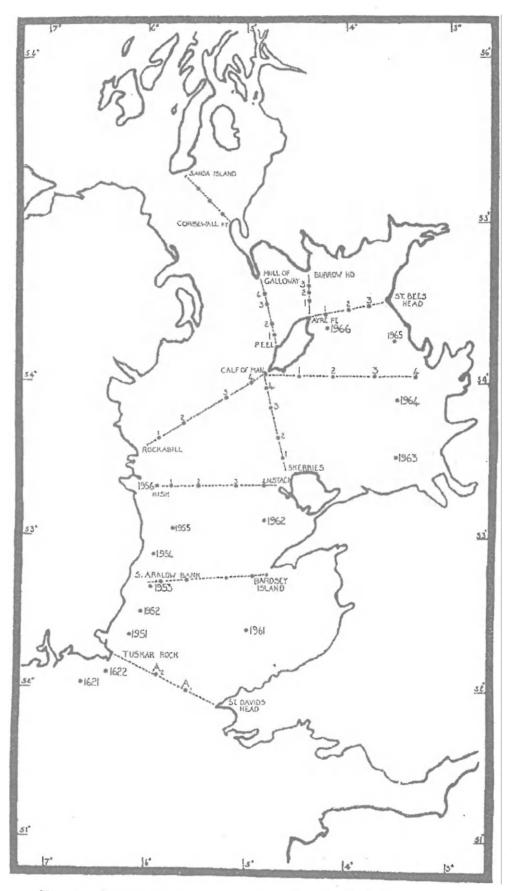


Fig. 2.—Current-stations and sectional lines in the Irish Channels.

then if t is taken at intervals of 1 hour from -6 to +6, we can eliminate u_2 by taking values of $\frac{1}{2}\{u(t) + u(-t)\}$, and this will be equal to $u_1 \cos \sigma t + a$ "constant," if there be any permanent current or M_4 . It is a simple matter to deduce from the maximum and minimum currents, at t = 0 and t = 6, the values of u_1 and the "constant." Similarly, from $\frac{1}{2}\{u(t) - u(-t)\}$ we obtain values of $u_2 \sin \sigma t$ and the maximum value at t = 3 gives u_2 . A simple graph enables a proper estimation, and an allowance for errors of observation, to be made. As a rule, no greater accuracy has been attempted than to give velocities to the nearest multiple of 5 cm. per second.

5. Dynamical Equations.

The relations between the currents and the elevations at and near a point are expressed in the form (PROUDMAN and DOODSON, loc. cit.):—

$$\frac{\partial \zeta_1}{\partial x} = -au_2 + bv_1 - cF_1$$

$$\frac{\partial \zeta_1}{\partial y} = -av_2 - bu_1 - cG_1$$

$$\frac{\partial \zeta_2}{\partial x} = au_1 + bv_2 - cF_2$$

$$\frac{\partial \zeta_2}{\partial y} = av_1 - bu_2 - cG_2$$

where x, y are measured in units of 10^7 cm., u, v are measured in centimetres per second and $a=1\cdot432$, $b=1\cdot487\sin\lambda$, $c=1\cdot019\div h/10000=56/f$, $\lambda=$ latitude, h= depth in centimetres, f= depth in fathoms, and F_1 , G_1 , F_2 , G_2 , are components of the frictional forces at the bottom of the sea. We define F_1 , F_2 to be the coefficients of $\cos \sigma t$, $\sin \sigma t$, respectively, in $0\cdot002u\sqrt{u^2+v^2}$ and G_1 , G_2 to be similar coefficients derived from $0\cdot002v\sqrt{u^2+v^2}$. These coefficients have generally been obtained by direct harmonic analysis of the synthesised expressions for $u\sqrt{u^2+v^2}$, $v\sqrt{u^2+v^2}$ at intervals of $\sigma t=0^\circ$, 30° , 60° , ..., 330° , but sufficiently accurate values, in view of the roughness of the data, can be obtained from the formulæ:

$$\begin{split} \mathbf{F_1} &= 0 \cdot 0017 u_1 \sqrt{u_0^2 + v_0^2}, \quad \mathbf{F_2} &= 0 \cdot 0017 u_2 \sqrt{u_0^2 + v_0^2} \\ \mathbf{G_1} &= 0 \cdot 0017 v_1 \sqrt{u_0^2 + v_0^2}, \quad \mathbf{G_2} &= 0 \cdot 0017 v_2 \sqrt{u_0^2 + v_0^2}. \end{split}$$

The dynamical equations are expressed in terms of the mean currents in the vertical line below the point, from top to bottom, but in practically all cases we only know the surface currents. Fortunately, it is known that the surface current has often been found to be very approximately equal to the mean current.

6. Rates of Change of Amplitude and Phase-lag.

From the equations $H^2 = \zeta_1^2 + \zeta_2^2$ and $\tan d = \zeta_2/\zeta_1$, using the special time origin, we immediately deduce the rates of change of H and of d in any direction (s) to be given by

$$\frac{\partial \mathbf{H}}{\partial s} = \cos d \, \frac{\partial \zeta_1}{\partial s} + \sin d \, \frac{\partial \zeta_2}{\partial s}, \qquad \mathbf{H} \, \frac{\partial d}{\partial s} = \cos d \, \frac{\partial \zeta_2}{\partial s} - \sin d \, \frac{\partial \zeta_1}{\partial s} \, \dots \, . \quad (6.1)$$

Let ψ , ψ' = angles made by co-tidal and co-range lines with Ox, respectively;

R, R' = positive values of H
$$\frac{\partial d}{\partial s}$$
, $\frac{\partial H}{\partial s}$ in directions $\psi + 90^{\circ}$, $\psi' + 90^{\circ}$, respectively.

For a co-tidal line g (and therefore d) is a constant, and for a co-range line H is constant; hence

$$\mathbb{R} \cos (\psi + 90^{\circ}) = \mathbb{H} \frac{\partial d}{\partial x} \qquad \mathbb{R} \sin (\psi + 90^{\circ}) = \mathbb{H} \frac{\partial d}{\partial y},
\mathbb{R}' \cos (\psi' + 90^{\circ}) = \frac{\partial \mathbb{H}}{\partial x}, \qquad \mathbb{R}' \sin (\psi' + 90^{\circ}) = \frac{\partial \mathbb{H}}{\partial y},
\mathbb{R}' \cos (\psi' + 90^{\circ}) = \frac{\partial \mathbb{H}}{\partial x}, \qquad \mathbb{R}' \sin (\psi' + 90^{\circ}) = \frac{\partial \mathbb{H}}{\partial y},$$

whence R, R', ψ , ψ' can be uniquely determined. By this convention the phases and amplitudes increase positively in the directions 90° in advance of the directions of the co-tidal and co-range lines respectively.

7. Deductions at a Single Station.

If the currents have been observed at a single station then the direction of the cotidal and co-range lines are functions of g only and not of g. Possible limits of g may be suggested by the coastal data, and generally it is sufficient to compute ψ , ψ' for two (or at most three) values at intervals of 30° in g. Approximate values of g can then be obtained and the distances between neighbouring co-tidal (and co-range) lines can thus be obtained. This method is very valuable, even if no other is available.

If integrations of $\partial \zeta/\partial s$ are possible from coast to coast then the values of H and g at stations on the line become known, but it frequently happens that the current-stations are not well placed for this purpose; the terminal points of the line may be among shoals or eddies, involving difficulties as to the proper terminal values of H and g, and these terminal values in any case may be subject to criticism; hence it is desirable to compute the directions of the co-tidal and co-range lines at each station, however the values of g may have been obtained (Table VIII).

8. Deductions from Coastal Currents.

Near the coasts the currents move backwards and forwards in one direction, generally parallel with the coast; hence $v_0 = 0$. If we take the axis of x to be along the direction of the current (θ) and take the origin of time so that the current is represented

by $-u_0 \sin \sigma t$, then, using zero suffixes to denote quantities referred to this orientation of the axes, we can write for equations (6.1), (6.2)

$$\frac{H}{bu_0} \frac{\partial g}{\partial x} = -\frac{a}{b} \sin \phi + e \cos \phi = \frac{R}{bu_0} \cos (\psi_0 + 90^\circ),$$

$$\frac{1}{bu_0} \frac{\partial H}{\partial x} = e \sin \phi + \frac{a}{b} \cos \phi = \frac{R'}{bu_0} \cos (\psi_0' + 90^\circ),$$

$$\frac{H}{bu_0} \frac{\partial g}{\partial y} = \cos \phi = \frac{R}{bu_0} \sin (\psi_0 + 90^\circ),$$

$$\frac{1}{bu_0} \frac{\partial H}{\partial y} = \sin \phi = \frac{R'}{bu_0} \sin (\psi_0' + 90^\circ),$$

where $\phi = d - d' - 90^{\circ}$ and $e = 0.0017cu_0/b$.

The computations required by these formulæ to give ψ_0 , ψ_0' , and thence ψ , ψ' by adding 0, the direction of the axis Ox_0 , are very simple. The value of g is obtained from the coastal data. The angle $d' + 90^\circ$ is obtained taking the time at which the current changes from the direction θ to the opposite direction, in hours after high water at Dover, and multiplying that number by 29.

The directions of the co-tidal lines and co-range lines for a large number of points along the coasts have been computed and utilised.

9. Coastal Elevations.

Harmonic constants for M₂ are only known at a few stations round the coasts, and the tidal data at coastal stations consist of:—

H.W.F.C. = the time-interval between the local common transit of sun and moon at syzygy and the succeeding high water at the place.

L.W.F.C. = the corresponding low water interval.

M.H.W.S. = Height of mean high-water springs, relatively to chart datum, which is approximately low-water springs.

M.H.W.N. = Ditto for mean high-water neaps.

M.L. = mean level.

For only a small fraction of the places do we find all the above quantities tabulated, and often only the first and third are given. The simplest procedure for the deduction of M_2 -constants is to compare values of the known harmonic constants g with the values of k = H.W.F.C. (in hours) \times 28·984, and this method is of value only if both g and H.W.F.C. are well established. It is necessary, of course, to avoid places like Portland (with double low-waters), Southampton (with double high-waters), and Havre (in a shallow estuary), and ultimately we chose Liverpool, Dover, and Brest as suitable places. Curiously enough, we find that at these places g - k is only \pm 1° and we have accordingly taken g = k throughout, for simplicity. The theoretical relation between H.W.F.C.

and g is not, in any case, very simple as it depends upon the relations of S_2 to M_2 , and other constituents such as M_4 , MS_4 , ..., all have influence.

With regard to the coastal elevations H, greater difficulties occur because chart-datum does not necessarily mean low-water springs. Thus, at Penzance (very near Newlyn, for which H = 5.6 feet) we have S = M.H.W.S. = 18.5 feet, N = M.H.W.N. = 14.8 feet, L = M.L. = 10.6 feet, and we see at once that chart datum is about 2.5 feet below the level of low-water springs, whence values of S, N, L, referred to M.L.W.S. as datum, are 16.0 feet, 12.3 feet, 8.1 feet, respectively; of course, in this reduction we have taken the mean level at about half the value of S. The ratios of H to S, N, L (with $L = \frac{1}{2}S$) at Newlyn thus become 0.35, 0.46, 0.71 respectively; the latter value is, of course, the ratio of H to S - L if $L = \frac{1}{2}S$.

In general, wherever S and L are tabulated we obtain H from the difference, using an appropriate factor.

An alternative method is to compare H with $\frac{1}{2}S + \frac{1}{2}N - L$, since the latter quantity is approximately itself equal to M_2 , modified by shallow-water constituents. In practice we found that there is nothing to be gained by this method. With proper reductions, we get the following values:—

					H	/(S-L).	H/S.	H/N.	N/S.	$H/(\frac{1}{2}S + \frac{1}{2}N - L).$
Liverpoo	1		i	4	*	0.74	0.37	0:47	0.78	0.95
Dover	4		a	4	6	0.76	0.38	0.49	0.78	0.97
Brest	4	4			4	0.68	0.34	$0 \cdot 46$	0.74	0.92

These values agree very well with deductions from the known values of M₂ and S₂. We have used interpolated values at intermediate places.

In Table II we give values of H and g, deduced from the Admiralty Tide Tables as explained above. Two columns are given for H: in the first column the values are

Table II.—Values of H and g for Coastal Stations.

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]	н.	Ħ.	9.	H.	日.	g.		В,	H.	g.
St. Mary's 1 Tresco 1		172 244 — — — — — — — — —	130 124 127 123 136 130 131 133 143 147 151 157 157	Salcombe 154 Start Point 156 Dartmouth 152 Torquay 147 Exmouth — Lyme Regis 122 Bridport 127 Portland Bill — Portland* 64 Ncedles — Portsmouth 140 Ryde — Bembridge Point — Chichester —	122 - 95 78? - 146 150 161	165 165 182 174 187 184 175 191 194 282 334 329 329 334	Ventnor Nab Tower Selsea Bill Littlehampton Shoreham Brighton Newhaven Eastbourne Hastings Rye Bay Dungeness Dover* Deal Ramsgate*	138 	187 240 277 186	321 330 341 329 326 326 325 316 329 328 331 327 341

TABLE II—(continued).

(b)

	H.	H.	g.		Ħ.	H.	g.		H.	H.	g.
Calais Wissant Boulogne Le Touquet Berck St. Valery-en-Somme Cayeux Le Tréport Dieppe St. Valery-en-Caux Fecamp Havre* Honfieur	239 253 300	345? 342? 335?	340 338 329 326	Ilos des Ehbiens Braye	188 190 229 280 335 	430? 440? 	228 218 206 192 191 181 181 183 182 178 180 196	Portrieux Paimpol Ile de Brehat La Croix Light Lizardreiux Heaux de Bréhat Plougrescan Treguier Ploumaneal Ile Lovet Ile de Bas L'Abervrach Porsal	350 355 325 315 330 305 286 304 269 260 243 250 231		168 172 174 170 170 169 156 158 156 142 139 132
Trouville	230 232	_	264 275	St. Helier	281 356 329		192 186 181	Ushant	215 210 202		105 114 107
Courseulles	212 220	_	258 266	Les Ecrehoux Les Minquiers	341 356		189 185	Brest*	207 204	_	108 104
St. Vaast	206		258 252		415 370	_	181 174	Douarnenez	193		104 103

(c)

	Н.	H.	g.		H.	H.	g.	П.	Н.	g.
Crookhaven Skull Cape Clear Baltimore Castletownsend Clonakilty Bay Courtmacshessy Kinsale Queenstown* Cork Ballycottin	106 	93 - 122 - - 143	- 1	Bridge New Ross	132 134 134 148 139	144 99? 99? 63? 53?	158 154 177 176 164 166		53! 41? 38? 41? 43? 40! 97? 134 148	217 231

TABLE II.—(continued).

(d)

	H.	H.	g.		H.	Н.	g.		H.	H.	g.
Lynmouth Porlock Minehead Watchet Bridgewater Bar Weston Walton Bay Portishead Bristol	225 223 —	244 251 231 242 247 290 298 323 346 365 382 - 434 - 347 308	123 130 137 136 151 152 167 152 173 165 175 183 189 190 198 200 204 208 212 201 231	Chepstow Newport Cardiff Flatholm Barry Nashpoint Portheawl Port Talbot Swansea Worms Head Llanelly Burry Port Ferryside Carmarthen Bar Tenby Caldy Road Stackpole Quay Pembroke Dock St. Anne's Head Skomar Smalls Lighthouse Ramsey Sound	428 409 390 385 306 300 264 283 262 279 257 218 208 215	413 		Fishguard Newport Cordigan New Quay Aberystwith Aberdovey Sarn-y-bach Barmouth Sarn Badrig Port Madoc Bar Pwllheli St. Tudwall Road Bardsey Port Dinlleyn Holyhead Amlwch Llanddwyn Island Carnaryon Trwyn du	152 148 161. 161. 153 150 146 131 126 173	132 130 130 - 154? 143? - 220	202 203 203 217 220 227 222 225 217 220 227 224 222 251 295 297 268 275 300

(e)

I	H. E	I. g .	I	H.	Н.	g.		H.	H.	g.
Londonderry Warren Point Coleraine Port Rush* Skerries Ballycastle Bay Red Bay Carnlough Maiden Rocks Larne Sealasaig Glengarrisdale	88	72 183 - 186 - 190 19 181 31 186 - 305 663! 326 314 12 153 17 153	Loch Crinan Feulin Ferry West Loch Tarbert Port Ellen Gigha Sound Mull of Cantyre Sanda Island Cambeltown Lamlash Brodick Bay Garrock Head Skipnese	68 74	35 45? 57 40 45? 90 91 105 102 108 93	146 139 136 73 145 68 306 338 340 342 343 343 343 343	Ardrishaig Burnt Isles Loch Striven Head Rothsay Bay Greenock* Port Glasgow Largs Mill Port Ardrossan Troon Ayre Stranraer	112 134 115 - 102 101 97	112 68? — — — — 115? 115	349 343 345 346 346 344 344 344 344 339

TABLE II.—(continued).

(*f*)

	H.	H.	g.		H.	H.	g.		H.	H.	g.
Point of Air Hilbre Island* Liverpool* N.W. Light Formby Point Freston Ribble Light Wyre Light Lancaster Glasson Dock Morecambe Barrow Docks Tarn Point Whitehaven Harrington Workington Maryport Silloth	298 305 300 310 308 310 267 262	296 302 221 	316 319 326 319 307 329 314 324 326 326 326 326 326 326 328 332 338	Port Carlisle Douglas* Ramsey Castletown Port St. Mary Calf Sound Peel Ayre Point Annan Foot Southerness Abhey Head Kirkeudbright Wigton Garliestown Port William Mull of Galloway Port Patrick Belfast	226 215 226 226 226 	210 	352 326 324 323 323 327 322 324 350 344 323 323 323 328 324 326 324 307	Donaghadee South Rock Killard Point Ardglass Newcastle Cranfield Point Soldiers Point Pile Light Boyne River Entrance Balbriggan Skerries Rogerstown Malahide Howth Dublin (Poolbeg Light) ,, (North Wall)*	147 149 162 156 158 149	174 	325 318 316 319 321 319 330 316 319 310 317 327 327 323 325 326

obtained from (S — L) and in the second column from the average values of H/S, H/N provided that the values of N/S are approximately correct. If there is any doubt about the values they are marked (?), and all values obtained from S alone are so marked. For places marked * actual harmonic constants at the place are given.

10. Western Approaches to the English and Irish Channels.

From the beginning, great difficulty was experienced in integrating from Land's End to Ushant, and it was not until a first approximation to the co-tidal lines had been made that it became possible to criticise the coastal data. These criticisms are discussed in the following account of the basic work done in this region. It was possible to avoid Land's End by integrating from the Scilly Isles to the Lizard; from a suitable point on this line, integrations were made on a line running south to Ushant and the He de Sein. Other sectional lines provided cross checks, and the processes of fitting together the various pieces of the puzzle were rather elaborate. In the end a satisfactory network of points was completed, sufficing to describe the tides over the whole region.

We shall denote by s the distance in units of 10^7 cm. along a sectional line. Gradients of ζ_1 and ζ_2 along the line are derived from the components to the East and to the North, given in Table I. We denote by letters P_1, P_2, \ldots , various special points (not current-stations) along the lines. We thus get the lines, distances, gradients, etc., as in Tables III to VII; the values of ζ_1 and ζ_2 , equal to H cos $(g+30^\circ)$, H sin $(g+30^\circ)$ respectively, are deductions except where they are obtained from the coastal data in Table II.

Table III.—Sectional Lines at Entrance to English Channel.

Line.	Point.	S.	∂ζ ₁ /∂s.	∂ζ₂/∂s.	ζ,,	ζ,.	H.	g.	Remarks.
I	P,	0.46	5 	- 3 -7 -24 -	-152 -153 -154 -160 -162 -161 -141	73 74 75 74 69 61 48	168 170 171 176 176 172 149	124 124 124 125 127 129 131	Junction line IV. From Table II. From Table II.
II	Land's End P P P Ushant Iie de Molène Ile de Sein .		- 3 - 22 - 30 	-15 32 52	-160 -161 -160 -165 -176 -176 -170 -123	61 60 57 66 101 120 123 134	171 172 170 178 203 212 210 180	129 129 130 128 120 116 114 103	Cp. Table II. From line I. Cp. Table II. From Table II. From Table II.
III	Lizard 8 252 C _s	0.25	- 6 - 9 - 6 23	8 26 40 58	-141 -141 -142 -143 -143	48 51 60 66 83	149 150 154 157 165	131 130 127 125 120	From Table II.
17	P ₁		-3 -35 -35 -44	12 I 27 28 21	-152 -152 -153 -155 -155 -158 -165	73 75 79 83 88 93	168 169 172 176 180 188	124 124 123 122 121 121	From line I. See line III.
V	Lizard S 253 S 256 Porsal	1.34	-48 -24	69 55	-141 -165 -195 -200	48 48 97 115	149 172 215 2 31	131 134 124 120	From Table II. From Table II.

TABLE IV.—Sectional Lines West of the English and Bristol Channels.

						0			
Line.	Point,	8.	∂ζ1/∂s.	∂ζ./∂s.	ζ1.	ζ2.	Н.	g.	Remarks.
AI	S 224 P ₁ B ₁ P ₈ P ₇ S 1621 Saltees	0.86 1.54 2.13 2.71 3.30	-34 -15 4 -1 41	17 -10? -34 - 54	-140 -153 -157 -156 -153 -143 -139	77 74 50 14 -25 -40 -35	159 170 165 157 154 148 144	121 124 132 145 169 166 164	See line IX. From lines I, IV
VII	C. Clear	0·20 1·52 2·30	-27 -24 -25 -11	-26 -26 -2 5 15	- 94 -107 -134 -157 -163 -164	64 56 48 50 52 57	112 121 142 165 171 173	117 122 130 132 132 131	Estimated. From line VI. Cp. II and Tuble II.

TABLE IV—(continued).

Line.	Point.	8.	$\partial \zeta_1/\partial s$.	∂ζ₂/∂8.	ζ ₁ ,	ζ_2 .	H.	g.	Remarks.
VIII	B, S 223 S 222	0.00 1.09 1.76	- 7 1 30	25 20 32	134 139 125	48 72 88	142 145 152	130 123 115	From line VII. See line IX.
IX	S 256 P ₆	0.00 0.38 0.53 1.55 2.60	64. 57 7 36	10 25 -14 -12	-195 -176 -168 -144 -125	97 101 103 88 70	215 203 196 168 143	124 120 119 119 121	From line V. From line II. See line IV. See line VI. See line VIII.

TABLE V.—Sectional Lines Across the English Channel.

Line.	Point.	δ,	∂ζ1/∂s.	∂ζ₂/∂s.	ζ ₁ .	ζ2.	H.	g.	Remarks.
X	S 493 S 261 S 272 Casquets	0.00 0.42 0.69 0.88 1.02 1.07 1.15 1.24 1.27 1.61 1.75 2.06	- 85 - 45 - 56	- 52 - 94 - 96 	- 72 - 98 -117 -135 -165 -180 -209 -229 -230 -296 -296 -339	- 67 -100 -122 -140 -159 -193 -188 -198 -198 -178 -178 -178	98 140 169 195 — — — — 370	193 196 196 196 	Table II.
XI	S 271 S 273 Barfleur St. Vaast	0.00 0.21 0.41 0.65 0.89 1.00 1.11	- 51 - 15 150?	-212 -141 -1091 -	70 56 46 39 37 39 64	- 32 - 70 -101 -137 -172 -185 -196	76 89 111 141 176 189 206	305 291 265 256 252 252 258	See text. Table II. Table II.
XII	S 543	0.00 0.28 0.56 0.83 1.11 1.39 1.67	143 	- 9 - 8? - 16	118 156 190 222 253 283 310	- 23 - 31 - 39 - 45 - 48 - 50 - 54	119 159 194 226 257 286 314	319 319 318 319 319 320 320	See text.
XIII	Pevensey	0.00 0.08 0.33 0.59 0.85 1.11	45	- 53 - 90	259 263 273 285 300 310	- 23 - 26 - 32 - 31 - 10 - 54	260 264 275 287 303 314	325 324 323 323 322 320	See text. Table II.

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TABLE VI.—Sectional Lines down the Centre of the English Channel.

Line.	Point.	δ.	$\partial \zeta_1/\partial s$.	$\partial \zeta_1/\partial s$.	ζ ₁ ,	ζ	H.	g.	Remarks.
XIV	S 255	$\begin{array}{c} -2 \cdot 68 \\ -2 \cdot 36 \\ -2 \cdot 10 \\ -1 \cdot 50 \\ -1 \cdot 11 \\ -0 \cdot 72 \\ -0 \cdot 36 \end{array}$	- 49 - 29 - 58 ?	- 32 1 -54 -120 2 - 44 -16	-154 -158 -169 -175 -188 -185 -170 -138 - 98	90 88 80 64 - 25 - 75 - 99 -105 -100	179 180 186 186 190 200 196 174 140	120 121 125 129 158 172 180 187 196	Line IX. Line IV. Line II. Line V.
ΧA	S 261 S 271	0·42 0·83 1·26 1·68 2·27	115 	- 19 - 60 - 67 	- 98 - 26 46 136 222 285 304	-100 -109 -101 - 72 - 45 - 34 - 27	140 112 111, 154 226 287 305	196 227 265 302 319 323 325	Line XI. Line XII. Line XIII. See Table II.

TABLE VII -Sectional Lines in Irish Sea.

Line.	Point,	δ.	∂ζ1/∂s.	∂ζ₂/∂s.	ζ ₁ .	ζ2.	H.	g.	ф.	ψ'.
1	Skerries	0.00 0.09 0.24 0.49 0.64 0.73	58 29 19 -12	129 84 66 99	173 179 190 201 205 203	-100 - 86 - 68 - 46 - 34 - 25	200 199 201 206 207 205	300 304 310 317 320 323	46 40 37 9	316 310 306 282
2	Rockabill	0·00 0·09 0·29 0·63 0·83 0·92	21 18 59 72	-42 3 7 29	149 152 160 176 191 203	- 25 - 29 - 32 - 32 - 29 - 25	150 154 162 178 192 205	320 319 319 320 321 323	103 45 35 18	18 354 310 291
3	Peel	0.40	9 - 5 -31 -56	46 57 95 127	176	21 	178 170	323 330	19 7 357 352	295 283 273 267
4	Ayre Point	0·09 0·15 0·21	-17 37 69	137 141 178	203 225	- 21 - 0	205	323	347 6 12	264 283 289

TABLE VII—(continued).

Line.	Point.	s.	∂5./∂s.	∂ζ₂/∂ε.	ζ,.	ζ2.	H.	g.	ψ.	ψ'.
5	Ayre Point	4			203	- 25	205	323		_
	2	0.26	96 97	$-11 \\ -5$	215 234	20 28	217	323 323	324 328	215 250
	St. Bees Head		 	6	248 257	- 28 - 27	250	323 324	341	260
6	Calf of Man	0.00			204	- 25	205	323	_	
	Languess	0.22	172	-35	220 237	- 28 - 31	230 238	323 322	19	295
	$\begin{bmatrix} 2 & . & . & . & . & . & . & . \\ 3 & . & . & . & . & . & . & . \end{bmatrix}$	0.44	98 80	-23 15	256 280	- 37 - 39	257 281	322 322	353	295 270
	Walney Island	1.01	93 —	50	299 309	$-30 \\ -24$	300 310	324 326	333	248

TABLE VIII.—Values of ψ and ψ' at Current Stations.

Station.	ψ.	ψ'.	Station.	ψ.	ψ'.	Station.	ψ.	ψ.	Station.	ψ.	ψ'.
222	338	273	361	325	270	543	330	275	A,	306	198
223	299	267	362	263	165	551	356	293	41.2	302	193
224	4	314	363	287	184	552	335	271	B,	329	60
231	318	324	364	285	183	571	27	300	B _a	311	47
232	318	225	365	289	166	572	347	251	$\mathbb{C}_{\mathbf{i}}$	308	217
233	332	280	366	264	176	1611	54	303	G,	301	176
234	303	219	367	248	156	1612	137	72	C,	262	122
235	189	1.05	371	277	177	1613	290	187	A	353	286
236	239	135	372	221	166	1614	292	188	В	288	135
251	286	193	373	182	160	1615	341	266	C	338	264
252	216	72	391	335	252	1621	218	131	D	325	301
253	315	240	392	12	352	1622	232	137	E	314	245
254	39	336	393	356	227	1731	352	294	F	331	233
255	283	187	401	53	347	1811	40	315	G	302	209
256	310	220	402	339	295	1812	293	224			
261	257	152	403	349	295	1813	300	241			
262	303	204	421	301	310	1951	346	233			
263	320	248	422	304	74	1952	334	218	1	Į	
271	292	201	491	123	42	1953	342	225	1]	
272	250	151	493	292	180	1954	9	286			
273	298	118	501	261	135	1955	18	332	1	1	1
341	246	149	502	354	250	1956	75	11			Į
342	279	172	541	300	230	1961	305	203		1	
343	314	197	542	328	262	1962	4	262			ļ

The gradients at points P_s to P_s are obtained by interpolation between pairs of stations lying close to the line.

The integrations along line I are very simple, and quite definite agreement with terminal values is easily obtained. The range is a maximum about S 422, corresponding

to the rapid increase of H round the coast near Land's End. Though P_2 is comparatively near to Land's End, both ζ_1 and ζ_2 are considerably less than the apparent local values at Land's End, as given in Table II.

The values of ζ_1 and ζ_2 along line II are obtained graphically; the curves were drawn to fit the gradients and the terminal data, with such slight modifications as were suggested by line IX. Note the differences between the values of H and g deduced for Ushant and those given in Table II.

No great importance is attached to lines III and IV as the observations do not appear to be very good, but both lines agree at S 232, and the values of H and g at C_1 are in quite good accordance with those deduced from line IX.

Two very important lines are those running from C. Clear to Land's End and northwards from S 224 to Saltees, on the Irish coast. The latter yields the unimpeachable result (indicated also by the coastal currents) that the co-tidal line through Saltees passes westwards, turns south, and later runs east—that is, the co-tidal line has a considerable curvature outward towards the Atlantic Ocean. At C. Clear we ignore the Tide Tables, for the coastal currents indicate that at C. Clear the value of g should be about equalto that for Skull and rather greater than for Crookhaven, while the value of H should be a little greater than at these other points. AnIestimate of the currents at points P_s and P_s has been made from stations S 1611, S 1612, S 1613, S 1614 and from stations S 1731, S 401, S 402, S 403, respectively. Lines VIII and IX yield results at S 222, a station well out at sea. There are certain small discordances between two derived values of H and g at one and the same station, but these are of the order of uncertainty in the coastal data. On the whole, the results of the integrations may be considered as very satisfactory.

11. The English Channel.

A number of stations lie on a line running approximately down the centre of the English Channel, but it is desirable first of all to use such cross-sectional lines as can possibly be obtained. The usual coastal difficulties are encountered, but information can be obtained along four of these lines, Nos. X to XIII.

A line due north (true) through the Casquets ends at station S 493, just east of Portland Bill, and a number of current-stations are found in this neighbourhood. By interpolating between current-stations data are obtainable for integration along the line from Portland Bill to S 493, with increments 0 and -5 in ζ_1 and ζ_2 respectively. Hence we deduce from Table II values of ζ_1 and ζ_2 at S 493 as given, with some measure of uncertainty, because of the peculiarities of the tides between Portland and Southampton. South of the Casquets, the line passes between places on or near the Channel Islands, and by interpolation between these stations a number of values of ζ_1 and ζ_2 have been deduced. Plotting gradients as usual indicates rapid changes in ζ_1 and ζ_2 , principally affecting H rather than g, near the Casquets. The value of H deduced from the Casquets is greater than would be indicated even by the incomplete data of

the Tide Tables, but an independent justification for this conclusion can be obtained by a rough integration westwards from Braye. The currents are rather strong but we obtain, apart from friction, $\partial \zeta_1/\partial s = 0.5$ U, $\partial \zeta_2/\partial s = 0.2$ U, where U is the maximum current, which we can take to be about 150 cm. per second (3 knots). With $\delta s = 0.13$ we find the values of ζ_1 , ζ_2 at Casquets respectively less than those at Braye by 9 and 4, whence, from Table II (b), we obtain H = 196, g = 195 at Casquets; this agrees very well with line X.

For sectional lines XI and XII it is necessary to deduce terminal values near the Isle of Wight. At Nab Tower we have $\zeta_1 = 138$, $\zeta_2 = 0$, and along the line to S 543 (s = 0.15) we compute gradients of ζ_1 , ζ_2 at S 543 to be 118, 202, whence $\zeta_1 = 119$, $\zeta_2 = -29$ at S 543; another estimate may be obtained from the line Ventnor to S 543, gradients 143, -9, and s = 0.10, whence $\zeta_1 = 117$, $\zeta_2 = -17$; the average values yield $\zeta_1 = 118$, $\zeta_2 = -23$ at S 543. Also, along the line due east from S 542 to S 543 we have average gradients 215, 59, and with s = 0.15 we deduce increments of ζ_1 , ζ_2 to be 42 and 9 respectively; whence $\zeta_1 = 76$, $\zeta_2 = -32$ at S 542, corresponding to H = 82, g = 307; the results of sectional line XI indicate the slightly smaller values of H = 76, H = 305, at S 542.

Sectional line XI is drawn almost due south to C. Barfleur, but we have deviated it slightly to pass through Barfleur and St. Vaast, for the terminal data are very indefinite because of the strong currents off C. Barfleur, and the conditions can only be satisfied by a rapid change in ζ_1 , near C. Barfleur.

For sectional lines XII and XIII we use Dieppe as a terminal point, and the coastal currents off the straight coast near Dieppe are of value. By considering the rates of increase of ζ_1 and ζ_2 along the coast a mean current was deduced from the dynamical equations, with a maximum velocity of 80 cm. per second and turning from an easterly direction to a westerly one at about 10 minutes before H.W.D.; this agrees fairly well with the data in "Tides and Tidal Streams," if we consider the stretch of coast between Treport and Valery-en-Caux.

Line XIII terminates at Pevensey between Eastbourne and Hastings, and interpolation gives H=260, g=325 at Pevensey; we ignore the data for Hastings, for the coastal currents indicate that g progressively increases from Eastbourne to Dover.

Finally, two lines down the centre of the Channel have been used as a check upon the results of the sectional lines across the Channel. Line XIV is drawn through S 233, S 254, S 255, to S 261, and line XV is drawn from S 261 through S 271, S 263 to Berck. After slight modifications of the first approximations yielded by the cross-lines, the values of ζ_1 and ζ_2 along XIV and XV are in almost perfect agreement with the gradients indicated by the currents.

12. St. George's Channel: Tuskar Rock to St. David's Head.

At a number of lightships in this Channel the current data have been tabulated at every hour from 6 hours before to 6 hours after high water at Dover, but for cross-

sectional integrations the data are not so accurate; the directions of linear currents, and the maximum velocities only being given, the currents are stated to be slack throughout the region between $\frac{1}{4}$ hour before until $\frac{3}{4}$ hour after high water at Dover. On the "first sectional line,"* from Tuskar Rock to St. David's Head, four stations are given, which we shall denote by a_1 to a_4 , and the information has been supplemented by special observations at stations A_1 and A_2 . These later observations are generally consistent with the former if we take the earlier limit of $\frac{1}{4}$ hour before H.W.D. for the time of turning. We then deduce results as follows:—

	Point.							S.	$\partial \zeta_1/\partial s$.	∂ζ ₂ /∂s.					
Tuska	J.		¥	ø		,		•	*	•			0.00	-121	— 50
a_4	L	,	4	F	b	4					·	ŧ	0.09	- 21	-127
A_2		٠		*	L					A		4	0.18	12	- 88
a_3	•	٠	٠		à	4		*	•	s#		٠	0.24	_ 9	117
A_1		4	٠	,	٠		4						0.43	— 21	- 93
a_2	4		4					٠	ća .	*	•		0.45	- 32	- 90
a_1	7	4	,			4				4	*	å	0.58	- 29	162
St. Da	ıvi	ď	3 E	I ea	nd					4		1	0.67	- 16	- 75

The gradients given for St. David's Head are, of course, due to coastal currents, data from Bais Bank; the currents near Tuskar turn about 2 hours before high water Dover, hence the somewhat exceptional values of the gradients there. The difficulties in dealing with this line arise from the poor coastal observations at or near the terminal points. From Waterford, on the South Coast of Ireland to Kingstown, on the East Coast, only the values of M.H.W.S. and M.H.W.N. are given, and most of these seem to be known only to the nearest foot or 0.5 foot, while the times are given to the nearest hour or half-hour in some instances. On the Welsh side, there is no really trustworthy information (except possibly at Fishguard) between New Quay and Skomar. At St. David's Head, for instance, H.W.F.C. must be between 06:00 (Ramsey Sound) and 06:58 (Fishguard), with M.H.W.S. lying between 17.0 feet and 12.4 feet.

Important results arise from the consideration of the directions of the co-tidal lines and co-range lines. The currents at the stations a_1 , a_2 , a_3 , a_4 are so similar in direction and magnitude that we can take average values; similarly for A_1 and A_2 ; whence we deduce:—

g.	Stations.	ψ.	ψ'.	R.	R'.
۵		0	0		
180	A	281	178	96	120
	α	279			
210	A	313	205	112	103
	a	310			

The sectional line is at an angle of about 330°.

^{* &}quot;Tides and Tidal Streams," p. 186.

The numerical values of $\partial \zeta_1/\partial s$, $\partial \zeta_2/\partial s$ are such that we can at once assert that ζ_1 and ζ_2 change, for all practical purposes, linearly with s, and their increments for Tuskar to St. David's Head will be -14 and -66 respectively. Let the values at the centre of the line be H_1 and H_2 . Then we have to find by trial such values of H_1 and H_2 that they will give values of H and g, on the coast, reasonably in conformity with the Admiralty Tide Tables, and such that the direction of the co-tidal line and co-range line through the mid point are correct. We obtain results as follows:—

	Mid P	oint.		Tus	kar,	St. David's Head.		
H.	y.	H_1 .	H ₂ .	II.	g.	H.	g.	
100	$ \begin{cases} 180 \\ 210 \end{cases} $ $ \begin{cases} 180 \\ 210 \end{cases} $	- 87 50 130 75	- 50 - 87 - 75 -130	82 69 129 118	162 201 168 205	125 132 174 182	191 214 188 213	

Variation in H makes only little difference in the resulting values of g. It is clear that the co-tidal line through the centre of the sectional line must strike the Irish coast to the north of Tuskar, and the Welsh coast to the south of St. David's Head; and since the sectional line is along $\theta = 330^{\circ}$, then ϕ is a little, and only a little, less than 330°. The latter argument would point to $g = 210^{\circ}$ approximately, at the centre of the line, but this appears to be undoubtedly too high for the coastal data; the highest value we can allow for the centre of the line is about 195°. The co-range lines definitely indicate H = 110, very nearly, at the centre of the sectional line, whence we deduce:—

The directions of the co-tidal lines are not altogether satisfactory by this compromise, but examination of a few observations at A_1 and A_8 which were made near the bottom shows that the currents at the bottom are a little to the left of those at the top and turn a little later, and the corrections due to each tend to make ϕ larger and more in conformity with the run of the co-tidal lines.

It has already been stated that on the Irish coast the tidal data do not appear, at first sight, to be worthy of much confidence, but the values deduced for Tuskar are in some agreement, as it happens, with the value of Table II (c). On the Welsh coast, we have ample evidence that the values of H for Smalls Lighthouse and Ramsey Sound in Table II (d) are much too high, for the coastal currents show beyond all question that the range of tide steadily decreases from Swansea to Fishguard, and a first approximation to the co-range lines shows that the range is a minimum near Fishguard. It appears, therefore, that the values deduced for St. David's Head are reasonably correct.

13. St. George's Channel: S. Arklow to Bardsey.

The four current-stations on this line yield the values of gradients as follows, ψ and ψ' being computed for $g=240^{\circ},\ d=270^{\circ}:$

Point.	s.	$\partial \zeta_1/\partial s$.	$\partial \zeta_2/\partial s$.	ψ.	ψ'.	H.	g
S. Arklow	0.00		_			35	228
	0.09	26	-150	354	252	48	235
	0.28	38	-128	351	251	73	241
	0.53	1.7	-117	357	260	106	244
	0.72	-28	-112	15	282	125	243
Bardsey	0.81	and the second first				135	242

The values of H, g are deduced with the help of the following considerations. If H_1 , H_2 , are the values of ζ_1 and ζ_2 at the centre of the line then the terminal values are $H_1 - 12$, $H_2 + 55$ at S. Arklow and H_1 , $H_2 - 46$ at Bardsey; and since the gradients of H in this instance are those of ζ_2 we should get an increment of 101 in H from Arklow to Bardsey. The coastal values of H, g, at S. Arklow and Bardsey are given in Tables II (c), II (d), as 40, 231° and 131, 222° respectively. By taking various values of H_1 and H_2 we deduce that the most appropriate terminal values are 35, 228° and 135, 242° respectively. On the Irish side these values fit quite well with the general data; on the Welsh side g = 222° is very unlikely and the revised value of 242° is in much better accordance with the coastal values as a whole.

14. St. George's Channel: Kish to Holyhead.

On this sectional line, we similarly deduce :-

		4	4						
Point.	S.	$\partial \zeta_1/\partial s$.	$\partial \zeta_2/\partial s$.	ζ_1 .	ζը.	H.	g.	ψ.	ψ'.
Kish	0.00			140	- 17	140	323	PROPER	m n
	0.09	5	- 82	140	24	141	320	63	348
	0.28	9	-100	142	- 42	140	313	55	341
	0.53	37	- 98	145	- 68	1.52	304	34	318
	0.72	23	-123	148	— 89	170	298	33	310
North Stack .	0.81	of Plane Institute III		150	100	180	295	deservation	

The values for Kish are obtained from the general run of the co-tidal and co-range lines from Kingstown. There is an excellent agreement between coastal elevations and gradients.

15. Irish Sea.

A number of sectional lines have been dealt with in this region, and the results are given in Table VII. Apart from small amendations of coastal data no difficulties have been experienced in deducing the elevations from coast to coast.

16. North Channel.

An amphidromic point has been located in this channel. Its position was located in the first place by assuming the values of ζ_1 and ζ_2 to be on two planes, whence the zero lines were deduced, of which the intersection gave the required point. By studying the directions of the co-tidal lines for assumed values of g it was possible to satisfy all the conditions.

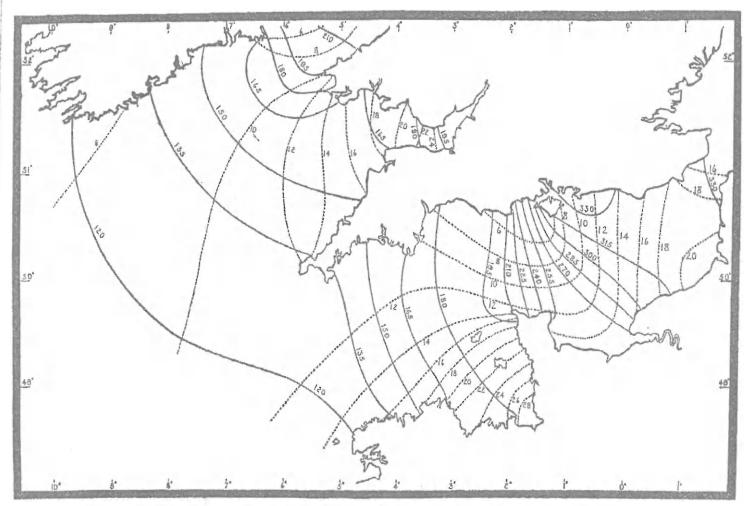
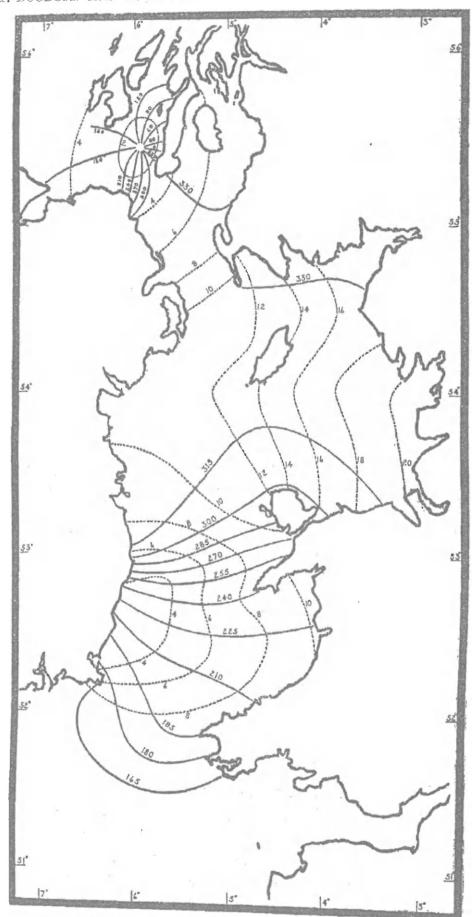


Fig. 3.—Co-tidal and co-range lines in the English Channel. (Reproduced from Admiralty Chart No. 301 by permission of the Controller of H.M. Stationery Office and the Hydrographer of the Navy.)

17. The Final Charts.

Having obtained all the data along the sectional lines and having drawn in all the appropriate directions of the co-tidal and co-range lines, it was a simple matter to construct the charts, though it is perhaps needful to say that a large number of considerations have had to be borne in mind during the process, and that it is impossible to discuss all the details. In the Straits of Dover the co-tidal line for 330° differs slightly from the line given on the chart for the North Sea, owing to improved treatment of the data. Co-tidal lines are represented by full lines, and co-range lines by broken lines; the phases are given in degrees and the amplitudes or semi-ranges in feet.



Fro. 4.—Co-tidal and co-range lines in the Irish Channel. (Reproduced from Admiralty Chart No. 301, by permission of the Controller of H.M. Stationery Office and the Hydrographer of the Navy.)

Summary.

Details are given of the method of construction of two new tidal charts for the English and Irish Channels, prepared for the Hydrographic Department of the Admiralty, and supplementing a similar chart for the North Sea.

The use of tidal currents for the calculation of gradients of the surface at a number of stations along a sectional line, together with the known coastal data, by graphical and numerical integration, enables the surface elevation to be computed from coast to coast. By a network of such lines the errors of observation are smoothed out and the resulting charts conform to the observational data and to the dynamical equations of motion. Previous charts have not satisfied the dynamical conditions. The new charts show that popular conceptions of the propagation of tides are largely erroneous, as the co-tidal lines do not bend inwards to the channel; that is, the time of high water is not accelerated in the centre of the channel, relatively to the times on the coasts. At certain places the new charts indicate errors of as much as 3 hours in the times of high water as shown by older charts.

