6.5 New data on heavy mineral distribution in the Sothern Kara Sea surface sediments

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Introduction

Among the traditional methods of marine geological investigations the study of heavy minerals is one of the main procedures helpful for the identification of source areas and transport pathways of terrigenous matter. The last 10 years of investigations of the Kara Sea modern sedimentation system led to the development of a detailed scheme of facies zonality for the Late Holocene sediments in this region (Levitan et al. 1996: Levitan 2001, 2002). According to this concept, there are four main facies subzones (FIIA -FIID) in the Ob-Yenisei (FII) zone, which are also divided into local parts. In particular. the zone of river sediments FIIA consists of the area of bottom erosion FIIA/1 and subzone FIIA/2 of winnowing of fine- and concentrating of coarse-size fractions due to strong currents. Zone of fresh and sea water mixing FIIB containes two subunits: FIIB/1, the area of hydrodynamic barrier with domination of coarse fractions, and FIIB/2, where fine material prevails. The transit zone FIIC is divided into five parts: FIIC/1 - area of bedrock outcrops, FIIC/2 - subzone of Quaternary relict sediments, FIIC/3 - area of recent sandy reworked sediments, and two units in paleoriver valleys (channels): FIIC/4, filled by sediments and FIIC/5 - empty valleys with bedrock outcrops due to strong currents. FIID, the facies zone of accumulation in the northern troughs, consists of FIID/1 - the area of northern and north-western Ob-Yenisei shoal slope characterized by high sedimentation rate, and FIID/2 - the subzone of low sedimentation rate.

Facts and Methods

Surface (0-1 cm) sediments from 41 geological stations (Fig. 6.17) sampled by means of box corer during the SIRRO R/V "Akademik Boris Petrov" (2001) cruise were investigated in fraction 63-125 μ m. The samples were separated by bromoform (density 2.87 g/cm³) for heavy and light subfractions. Heavy grains were mounted to permanent slides with meltmount (n=1,68), identified and counted using a polarisation microscope. Unfortunately, I5 examined samples contain just a few heavy grains; these results can only be useful for mapping and conclusions as supplementary information. It is necessary to note that we used only data with 200 and more grains in each sample and mineral groups (clinopyroxenes with aegirine-augite, amphiboles, etc.) for mapping and description.

Results

The results of mineralogical analyses are shown in Table 6.7. Among the heavy minerals in most of studied samples clinopyroxenes, epidote group and "black ore" minerals dominate (almost everywhere more than 15 %). The garnet and amphiboles content also exceeds 10 % in some of the investigated samples. The concentration of other minerals and mineral groups does not reach 10 %.

Clinopyroxenes.. The maximum amount of clinopyroxenes (Fig. 6.17) was found in the sediments of river valleys, in the eastern and northwestern parts of the investigated area. In the south of the region the content of these grains increases northward from the FIIA/1 (16.2 %) to FIIC/3 (39.6 %) subzone along the Yenisei transect. Very high amount of clinopyroxenes (31.7 %) was also discovered in the sediments of station BP01-72 in the Ob estuary. The surface sediments of FIID/1 and FIID/2 facies subzones are also characterized by high and medium amount of clinopyroxenes (19,4 – 28.6 %). Distribution of grains of this group in the sediments of FIIC/3 and FIIC/4 areas varies and, probably, rather depends on the general region peculiarities (geological structure) than reflects any local facies reguliarities. In particular, one of the largest content of clinopyroxenes (35,3 %) in sediments from the FIIC/4 zone (station BP01-28) is formed by well-rounded semi-transparent "old" grains never met in other studied samples here. The local decrease of these minerals is found in the central and midwestern parts of the study area (also subzones FIIC/3 and FIIC/4).

Epidote group minerals. The amount of epidote group minerals varies from 12,3 % in the Yenisei sediments to 37,0 % in the mid-east of the study area. Minimum values of epidote content were described in the sediments of the Yenisei transect, eastern part of FIIC/4 subzone and at the north-eastern area. Generally, the sediments of the Ob estuary as well as those from FIIC and FIID facies zones, are enriched by epidote (Fig. 6.18).

Black ore minerals. Black ore minerals definitely dominate in the south-east of the area (Fig. 6.19). These minerals prevail in the bottom sediments along the Yenisei transect and even exceed 50 % in the sample from station BP01-17. However, the sediments taken at station BP01-24 (at the northern end of this transect) contain very low quantity of black ore minerals – 6.9 %. This can be explained by the intensive accumulation of sand-size grains earlier, along the hydrodynamical barrier zone FIIB/1 located southward of this station. The Ob sediments have a low amount of this group minerals (maximum 13,3 %). Two other maxima of black ores are found in samples from stations BP01-68 (FIIC/3) and BP01-45 (FIIC/4). Here grains were redeposited from the Mesozoic rocks. Black ores content in the sediments of the northern part of the investigated region (FIIC and FIID zones) increases northward varying from 10 to almost 29 %. There is one exception with very low amount of these grains in the reworked and redeposited sediments of station BP01-55 (FIIC/3 subzone), also explained by the composition of underlying rocks.

Amphibole minerals. Relative content of amphiboles varies from 1,4 to 14,4 % (Fig. 6.20). Hornblende strongly dominates over the other amphibole minerals and composes 89 – 100 % of amphiboles. The amount of amphiboles is decreasing northward and exceeds 10 % at the northwest of subzone FIID/2. The maximum content of these grains (12.4 %) was found in surface sediments from station BP01-24 (described earlier)

located just near the coastal source. The Ob sediments are also enriched by amphiboles supplied by the river from the West Siberian Plain.

Garnet. The amount of garnet in bottom sediments does not reach 15 %. There are no strong variations of its distribution. In general, minimum values of garnet content are located in the nortwest of the shoal slope (zone FIID) and at the end of the Yenisei transect (zone FIIB) to the north of the hydrodynamic barrier area. The enrichment of sediments by garnet grains is increasing northward. On the whole, the garnet distribution depends on the influence of local sources, sediments composed by redeposited terrigenous grains containing higher amount of these minerals.

Titanous minerals. The minerals of this group are represented by sphene and rutile grains. The amount of Ti minerals varies from 1,1 to 8,7 %. The areas of sediments enriched by sphene and rutile are just close to their local sources: Yamal Peninsula, Taimyr coastal zone and at the north of the region with Mesozoic terrigenous underlying rocks.

Fe oxides and rock fragments. The amount of Fe oxides is about 3 % in average with the maximum value in the Yenisei estuary bottom sediments (6,7 %). The rock fragments content does not exceed 2 % in general, although there are some local anomalies with high amount of these grains in samples from paleo-valleys of FIIC/4 subzone and at the southwest of the shoal slope (FIID/1). It is necessary to note that the samples from the stations that we used just as the supplementary information (with insufficient heavy subfraction amount) are strongly enriched by rock fragments and Fe oxides at the Ob transect, both Ob and Yenisei barrier zones (FIIB/1), the areas of dominant redeposition (FIIC/3) and in the upper shoal slope (FIID/1).

Clinopyroxene/epidote ratio. Clinopyroxene/epidote ratio is an important indicator of the riverine terrigenous matter in the sediments of the Ob-Yenisei facies zone (Levitan et al. 1996). In accordance with the geological difference of the source provinces drained by both large rivers, the Yenisei bottom sediments have a higher value of this parameter than the Ob samples (Table 6.8). In general, this ratio is decreasing northward from the estuaries to the facies subzones IIC and IID (Fig. 6.21).

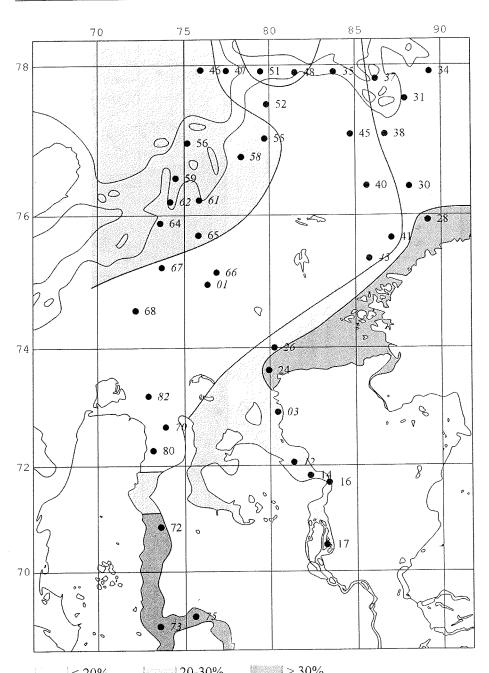
Discussion and conclusions

The distribution of heavy minerals in surface sediments have been studied according to the facies zonality of the eastern Kara Sea. Generally, a few main groups of heavy minerals dominate in the fine sand fraction of sediments: clinopyroxenes, black ores, epidotes, garnet, titanous minerals, Fe oxides and rock fragments. The distribution of black ores and the last four listed grain groups shows that they are located very close to the sources. The increasing of their amount (especially, Fe oxides and rock fragments) may be used as an indicator of local genesis of studied terrigenous grains. The variations of epidote content in sediments mostly depends on the presence of other mineral groups even along the Ob transect. The sources of clinopyroxenes are the Yenisei river and Taimyr Peninsula. The clinopyroxenes/epidote ratio distribution using as the indicator of Yenisei supply in the sediments exhibits a sublatitude zonality well coincided with the main facies units.

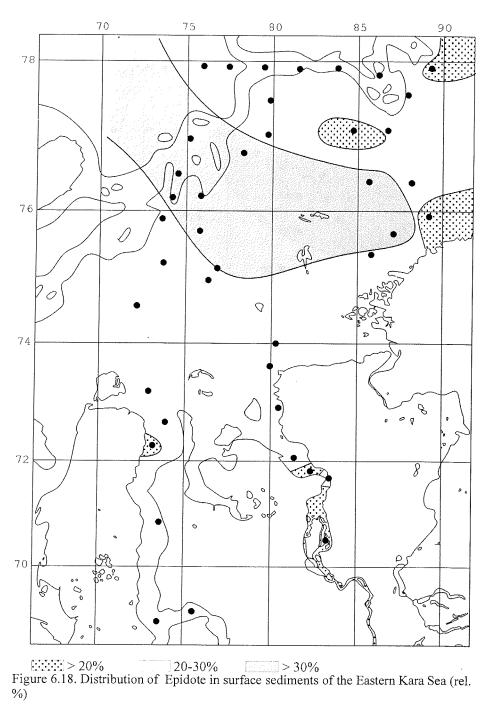
The mineral composition of FIIA/2 subzone samples (Table 6.8) in the south of estuaries demonstrates the dominance of local sources (the river coasts are located just close to studied sediments) for the formation of modern deposits. The hydrodynamic barrier zone FIIB/1 is characterized by strong accumulation of main mineral groups as from far continental sources (clinopyroxenes and epidote) so from the local ones (black ores, Fe hydroxides and rock fragments). But the distribution of amphiboles demonstrate that these minerals probably can overcome the barrier due to their high buyoancy comparably with other heavy grains. The samples from the FIIB/2 unit contain just a few heavy grains represented mostly by rock fragments. The composition of FIIC/3 unit sediments demonstrates the strong influence of local sources: underlying bedrocks and old Quaternary terrigenous sediments formed by riverine matter. A definite role in this relation can play also the supply of sediment matter from numerous small islands in the zone. In turn, FIIC/3 sediments are partly eroded and transported to the FIIC/4 and FIID subzones. The sediments of the paleo-valleys (FIIC/4) and shoal slope (FIID/1) contain material both from far sources (especially, Yenisei run-off, and probably redeposited from the FIIC/3 subzone area) and local sources. Increasing of clinopyroxenes, Fe hydroxides, rock fragments and ClPx/Ep ratio in FIIC/4 sediments in comparison with FIIC/3 sediments (Table 6.8) proves that channels are the pathways of riverine sediments northward the mixing zone. The role of local sources increasing for the FIID/2 area sediments - it follows from the high amount of garnet in the east and amphiboles in the west of this subzone.

Acknowledgment

We would like to express our gratitude to the analysts V.P. Kazakova and A.N. Rudakova for performing the separation of sand fraction. This is grant RFBR # 02-05-64017.



< 20%
20-30%
> 30%
Figure 6.17. Distribution of Clinopyroxenes in surface sediments of the Eastern Kara Sea (rel. %). Samples from the stations with numbers in italic contained less than 100 grains



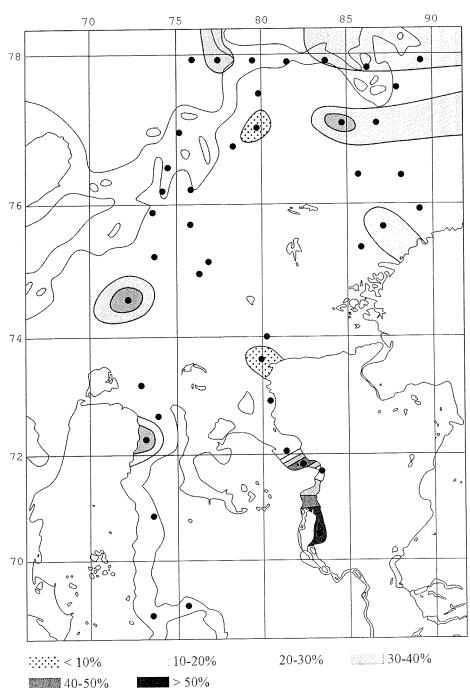
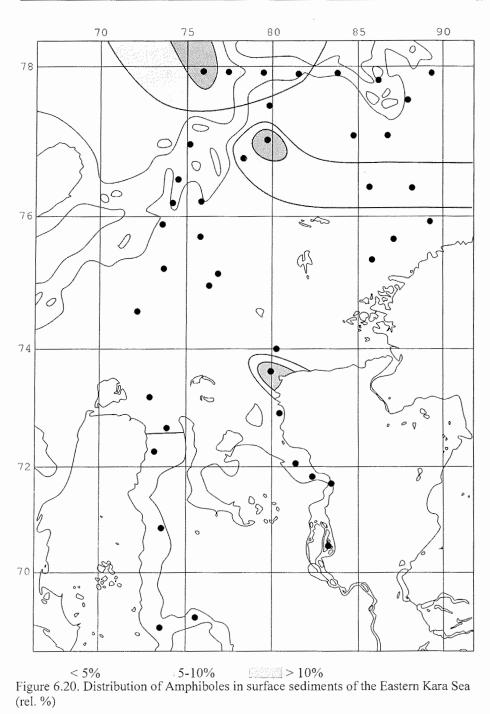
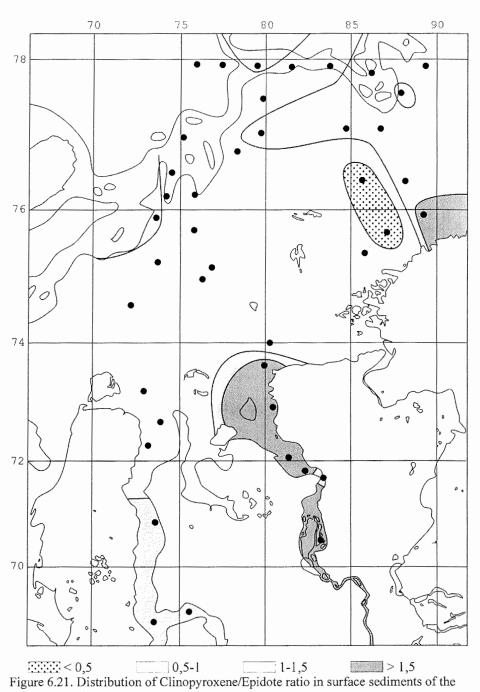


Figure 6.19. Distribution of Black ore minerals in surface sediments of the Eastern Kara Sea (rel. %)





Eastern Kara Sea

Table 6.7: Heavy minerals in the Eastern Kara Sea surface sediments (rel.%)

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Ċġ	grain	counted	0.4	0.4	, c)) (ر,ں د , ہ	٥, 4	0,0	0,0	0,0	0,0	0,4	0,0	1,1	8,0	0,0	0.0	0,0	0,0	0,4	0,4	0,0	0,4	0,0	0,0	0.0	0,0	0,0
Staur		grains c		0.0	0.7	, c),)	4,0	0,4	0,0	0,0	0,0	0,0	0,0	0,0	1,1	0,0	0,8	0,4	0,0	0,4	0,0	0,0	0,0	1,1	0,0	0.0	0,0	0,4
Act-	Trem	> 200	0,0	0,0	0.0	, c	0,0	0,0	0,4	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,4	0,0	0,0	0,4	8,0	0,0	0,0	0,0	0,0	0.0	0,0	0,0
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Brown	HBI		8,0	0,0	0,7	· -	t, -	·,'.	7,7	8,0	6,0	8,0	0,4	6,0	1,1	0,5	0,0	0,4	1,3	0,0	6,0	3,0	0,0	0,0	0,0	0,0	0,7	0,0	8,0
HBI			2,4	2,4	0,7	_	2,1	J, 4	4,0 1,0	3,7	/,'	3,2	3,1	3,7	5,6	8,0	10,0	4,9	5,8	7,3	2,6	10,2	3,5	4,8	5,7	2,8	3,7	5,6	t.
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B	ore		45,6	23,7	50,6	6.9	11.0	11,7	1/,/	19,4	2 x x x	7,07	21,6	13,9	2.1.2 Ci	33,3	13,9	21,1	19,7	16,4	15,7	8.9	15,1	17,8	10,7	17,0	31,2	13,3	30,4
$\mathbb{E}_{\mathbf{p}}$			12,3	24,1	12,5	9 66	18.7	, , ,	7,5	4,07	7,01	6,07	20,5	37,0	30,9	13,1	26,4	26,8	27,4	26,7	25.7	23,1	33,6	30,4	21,8	32.4	20,1	25,7	8,61
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Monaz	0,4	0,0	0,0	0,4	0,8	0,4	0,4	0,0	8,0	0,5	0,4	0,0	0,4	0,4	0,0	0,0	0,0	0,0	0,0	0,4	0,4	0,4	0,7	0,0	1,2
Brook.	0,0	0,0	0,0	0,4	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0	0,0
Cass	0,0	0,0	0,0	0,0	0,0	1,2	0,4	0,4	8,0	0,0	0,7	0,5	0,0	8,0	0,4	0,0	0,4	0,4	0,0	0,0	0,0	0,4	1,0	0,4	1,2
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ls Zr Chr Spi grains counted	0,8	0,0	0,5	0,7	1,2	1,7	1,7	0,4	1,2	0,0	1,9	2,7	0,4	3,4	6,0	6,0	0,0	0,0	0,0	6,0	1,1	1,6	6,7	0,4	4,7
Arfveds > 200 gr	0,0	0,0	0,0	0,0	0,0	0,4	0,4	0,0	0,0	0,5	0,0	0,0	0,4	8,0	0,0	0,0	0,0	0,4	0,8	0,0	0,0	0,4	0,0	0,0	0,0
Aeg- Aug	0,4	0,4	2,3	1,4	8,0	8,0	0,0	2,0	8,0	0,5	0,7	8,0	1,7	0,4	0,4	6,0	0,0	2,7	1,2	6,0	1,1	2,0	0,3	2,4	0,4
Andal.	0,0	0,0	0,0	0,0	0,4	0,0	1,3	0,4	0,0	0,0	0,0	0,0	0,0	0,4	0,0	0,4	0,0	0,4	0,4	0,0	0,0	0,0	0,0	0,0	0,0
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Sill	0,0	0,4	0,0	1,1	0,0	0,0	0,4	0,4	0,0	0,0	0,0	0,3	0,0	0,4	0,0	0,4	0,0	0,4	0,4	0,0	0,4	0,0	0,0	0,4	0,0
ō	0,0	0,4	0,5	1,1	2,0	8,0	0,4	0,4	2,3	6,0	0,0	1,4	6,0	8,0	0,4	0,4	0,4	1,1	8,0	0,4	0,4	0,4	0,0	0,0	0,0
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Brown	TOTA		2,5) (0,0	0,0	0	5,5	13,7	2,1		۱,۱	0,0	0.0	0.0) (0,0	0,0		, (0,0
HBI			4,9	0	, ,	13,5	3,9	7 1	, c	۲,۶	7,4	8), (1,9	4,6	8	i v	٠,٠	5,7	6.3	, ,	7,0
Garn			6,2	× 7	, ,	٥, ′	2,0	2.1	í c	0,0	1,	5 4	- c	1,9	0,0	5.6	0,0	٠, ١	2,9	5.2	` c	0,0
BI))		6,2	0 0	2,0	0,0	4,9	0.0	0 0	0,0	1,1	0.0	, (0,0	1,5	0.0	, C	oʻ,	1,4	0.0	` C	0,0
Ер			12,3	21.7	20.00	7,07	8,6	37.5	, ,	2,0	6,4	6.7		٥,٠	35,4	30.6	18.0	, ,	4,3	8,3	18.0	10,0
CIPx			18,5	30,4	20.0	5, 6	21,6	41.7	36.8	0,0	16,0	12.9	13.5	ر, د د د د	26,2	11,1	15,6	, <u>,</u>	1/,1	10,4	277	,,,
Station BP01		,	-	٣	12	,	97	37	43	2 1	28	61	Ş	7 \	99	<i>L</i> 9	70	· {	?	75	83	3

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Pyrite		<	0,0	0,0	0.0	0,0	0,0	0,0	0 0	5 0	0,0	0.0		0,0	0,0	0.0	0,0	0,0	0.0	0,0
Monaz			0,0	0,0	0.0	0.0	0,0	0,0	0.0		0,0	0.0		0,0	0,0	0	0,0	0,0	0 0	0,0
Brook.			0,0	0,0	0,0	0.0	, ,	0,0	0.0	` c	0,0	0,0	00	, ,	0,0	0	0,0	0,0	0.0	0,0
Cass		0) (0,'	0,0	0.0	, ,	0,0	0,0	_	1 , 1	0,0	0.0	0,0	0,0	0.0	, (0,0	0.0	0,0
Musc		0	, ,	2,0	0,0	0.0		0,0	0,0	0	, ,	0,0	0.0	, ,	0,0	0.0		0,0	0.0	0,0
Voulk. glass		0.0	0,0	, ,	0,0	0,0	0.0	0,0	0,0	00) ((0,0	0.0		,	0,0		0,0	0,-	0,0
Rock fragm.		0.0	0.0	5 1	27,5	0,0	0.0	2,5	14,9	3.2	1,5	13,4	0,0	16.7	, ,	17,2	37.1	1,10	6,3	8,6
Chrom Spinel	nted	0.0	0.0	, c	0,0	0,0	0.0	0,0	0,0	0.0		0,0	0,0	0 0) (0,0	00	,	0,0	0,0
Zr	ins counted	0,0	0.0	, (0,0	0,0	0.0		0,0	=	0	0,0	0,0	0.0	5 6	0,0	0 0	, ,	2,1	0,0
Arfveds	< 100 gra	0,0	0.0		0,0	0,0	0.0	. 0	0,0	0,0	0.0	0,0	0,0	0.0	, ,	0,0	0.0	,	0,1	0,0
Aeg. Aug		8,7	6.7	` -	0,1	4,7	2,6	7 3	ţ,	5,4	3 8	0,0	1,5	0.0	` -	٥, ١	1.4		7,1	3,3
Andal.		0,0	0,0	. 0	0,0	0,0	0,0	0.0	,,	0,0	0.0	, ,	0,0	0,0		0,0	0.0	, ,	0,0	0,0
Chlorite		8,7	6,7	0	0,0	7,4	7,9	2.1	1,1	8,6	0 0) · ·	۲,5	8,3	1.6	0,1	7,1	, ,	4,7	9,9
Ē		0,0	13,3	0	, ,	0,0	0,0	-	, ,	1,1	0.0	, (0,0	0,0	0	0,0	0,0		0,1	1,6
Sill	(0,0	0,0	0 0	0,0	0,0	0,0	0.0	2 0	0,0	0.0		0,0	0,0	0	, ,	0,0	0	0,0	0,0
ō	0	0,0	0,0	0.0	, ,	0,0	0,0	1.1		1,1	0.0		0,0	0,0	0	,	2,3	-	0,1	0,0
Enst- Hyp	t	8,7	0,0	2.9	, ,	2,5	10,5	4,3	, (7,7	0.0	0.0	7,7	8,3	73) i	2,9		1,7	9,9
Station BP01	,	n ;	12	26	3.7	5 5	43	28	6.1	10	62	99	00	29	70	> (//3	75	0 6	78

Table 6.8: Heavy mineral content in different lithofacies (rel.%) Lithofacies CIPx Garnet Black ores Amphib. Ti min. Fe hydrox. Rock frag. CIPx/Ep Εp A2 (Yen.) 22,9 0 n=1 12,5 1,5 50,6 1,5 1,1 3 1,8 B1 (Yen.) n=2 21,4 12,3 3,2 23,7 2,4 6,7 1,2 min 2,4 1,1 max 27,7 24,4 5,9 45,6 3,2 3,2 6,7 1,2 1,7 6,7 mean 24,6 18,2 4,6 34,6 2,8 2,8 1,2 1,4 B1 (Ob) n=2 min 16,2 19,8 4 13,3 5,1 2,8 2 0,4 0,8 max 31,7 25,7 6,7 30,4 5,6 7,5 2 3,6 1,2 24 22,8 5,4 5,2 2 mean 21,8 5,4 2 1 СЗ n=4 12,3 20,1 6,1 6,8 3,7 4,7 2,8 0 0,4 min max 24,6 32,4 11,9 31,2 14,4 6,2 6,1 3,8 1,1 18,6 26,6 mean 9 19 6,4 5,1 4,1 1,5 0,7 C4 n=6 13,1 13.1 0 3,2 6,9 1,4 2,8 3 0,4 min max 39,6 37 13,7 33,3 12,4 8,7 8,3 6,1 1,9 mean 23,8 23,6 8,8 16,6 5,8 5,8 4,9 2,9 1,1 D1 n=8 19,4 min 16,2 3,9 10,7 3,1 3 2,3 0,4 0,7 max 28,6 33,6 14 28,8 7,3 6,3 4,6 7,1 1,4 23,3 26,2 9,1 18,5 4,9 3,6 mean 5,2 1,9 0,9 D2 n=2 19,6 26,4 3,5 13,9 3 0,7 min 6,4 3,4 1,7

max

mean

23,4

21,5

26,8

26,6

6,8

5,2

21,1

17,5

10,4

8,4

4,8

3,9

4,8

4,1

1,9

1,8

0,8

0,8