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A permanent mud pum; ing installation as an alternative for local maintenance dredging

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KEYWORDS: mud flow, mud pumping, stratified mud, equilibrium profile, air jets.

ABSTRACT: A mud capture installation, provided with a fixed pump may be a substitute for conventional local maintenance dredging. If the mud deposits are stratified however the mud layer must be homogenized prior to the pumping. This can successfully be done using air jets. In the paper a theoretical background is provided for understanding the different behaviour of homogeneous and stratified mud.

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A PERMANENT MUD PUMPING INSTALLATION AS AN ALTERNATIVE FOR LOCAL MAINTENANCE DREDGING

SUMMARY To avoid difficult and expensive maintenance dredging at a harbour or dock entrance, in front of a navigation lock or under a landing-stage, it was suggested to build a mud capture installation, provided with a fixed pump. The idea is basically to remove the mud by (cheap) pumping, while it is still fluid, instead of (expensive) dredging of the consolidated mud.

A mud pumping installation may be very effective, provided that the mud layer is homogeneous, and the mud density is moderate. Stratified mud layers however, cannot be removed by pumping with the same efficiency. Therefore, layered mud deposits have to be homogenized before pumping e.g. using air jets, arranged as an air mattress, placed underneath the suction heads of the pumps.

RESUME Afin d'éviter des travaux de dragage locale, pour maintenir les profondeurs dans des ports, dans des chenaux d'accès des écluses, sous des pontons ou le long des quais, on a suggèré de construire une "trappe de vase", munie d'une pompe de vase fixe. Essentiellement l'idée consiste à enlever la vase lorsqu'elle est encore fluide, par pompage, au lieu de l'enlever après consolidation par un dragage coûteux.

On peut démontrer que le pompage de vase peut être très efficace, pourvu que la couche soit homogène et pas trop dense. Par contre le pompage de couches de vase stratifiées s'avère beaucoup plus difficile. Ces couches doivent d'abord être homogénisées par exemple à l'aide de jets d'air, montés sur un matelas, installé en dessous des bouches d'aspiration des pompes.

#### INTRODUCTION

Since the early sixties siltation has been a source of trouble for many Port Authorites. Indeed, by deepening and widening of the approach channel the cost of maintenance dredging increases exponentially. A similar and specific problem arises at the harbour of Zeebrugge (Belgium) and at several sites in Antwerp (Belgium).

In many cases the ship channel is considerably deeper than the surroundings, which may cause a steady fluid mud flow from the sea or the estuary towards the deeper parts of the harbour.

On the other hand in harbours, along quay walls, under pontoons and in small marinas, the mud which enters with the flood, settles down at slack water and forms a mud layer of increasing thickness, which hinders the commercial exploitation of the

To avoid the difficult and expensive maintenance dredging at a harbour or dock entrance; in front of a navigation lock or under a landing-stage, it was suggested to build a mud capture installation, provided with a fixed pump. The idea is basically to remove the mud by (cheap) pumping, while it is still fluid, instead of (expensive) dredging of the consolidated mud.

# LABORATORY TESTS

To study the feasibility of such a device, and to determine the design parameters, extensive Laboratory tests have been carried out at the Hydraulics Laboratory of the University at Leuven. In particular, a number of pumping tests were carried out in a 12 x 0.5 x 0.5  $m^3$  and a 2.4 x 2.4 x 1.2  $m^3$ flume to study the dynamic behaviour of North Sea

mud under pumping conditions.

It follows from the tests that a homogeneous mud layer can be pumped without difficulties up to densities of 1.150 and even 1.180 kg/m3 (The static shear stress then being of the order of I to  $3 \text{ N/m}^2$ ). The equilibrium slope of the water-mud interface being between 1/30 and 1/50. At the same densities and static shear stresses, stratified mud layers, built up by the precipitation of mud at slack water, behave in a very different way. The equilibrium slope, in the vicinity of the pump is much larger (1/1 !), which means that the "influence radius" of a pump can be very limited. This phenomenon was observed during a prototype test and was reproduced in the Laboratory.

In the subsequent sections a theoretical explanation is provided to explain why layered mud is much more difficult to remove by pumping than a layer of homogeneous mud.

### BASIC EQUATIONS

Assuming that the mud behaves as a viscous fluid, and that the mud flow can be assumed to be a steady and uniform density current in a prismatic channel, the simplified eq. of motion is (Kranenburg, Graf) :

$$U_2 \frac{\partial U_2}{\partial x} - g' S_0 + \frac{\tau_2 - \tau_1}{(\rho_2) S_2} = 0$$
 (1)

in which (Fig. 1):

U2 : the mean velocity of the fluid mud

a2 : the thickness of the mud layer

$$S_0$$
: the channel slope  
 $g' = g \left[ \frac{\langle p_2 \rangle - \rho_w}{\langle p_2 \rangle} \right]$  (reduced bouyancy)

 $\rho_w$ : the density of the water

(p2): the mean density of the mud

 $\tau_1$ : the shear stress at the water mud interface  $\tau_2$ : the shear stress at the channel bottom Taking into account the continuity equation, and since  $\tau_1$  = 0, unless a massive sliding of the mud layer occurs, Eq. 1 reduces to

$$\tau_2 = g'(\rho_2) a_2 S_0$$
 (2)

Since:  $\rho_2 = \rho_\omega + \alpha X$ 

in which X : the solids concentration (kg/m3)

$$\alpha = \frac{\rho_d - \rho_w}{\rho_d}$$

 $\rho_{\bf d}$  : the density of the solid mud particles Eq. 2 may be written as

$$\tau_2 = \alpha g a_2 \langle \chi \rangle S_0 \tag{3}$$

Assuming that X increases linearly over the depth of the mud layer (Fig. 1)

$$\langle \chi \rangle = \chi_0 + \frac{1}{2} m a_2 \tag{4}$$

in which  $X_0$ : the solids concentration at the mudwater interface

 $$\rm{m}$$  ; the concentration gradient For a steady uniform mud flow the eq. of motion thus reduces to

$$\tau_2 = \alpha g a_2 (X_0 + \frac{1}{2} m a_2) S_0$$
 (5)

This relation between S, a 2 and X is schematically represented in Fig. 2, and for specific values of the parameters in Fig. 3.

From these diagrams, several interesting conclusions may be drawn:

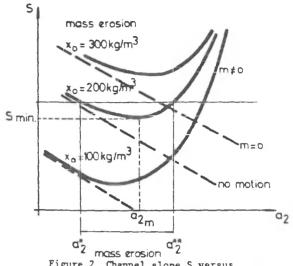


Figure 2 Channel slope S versus

Thickness of the sliding mud
layer a<sub>2</sub> for different values

of X<sub>0</sub> and m and fixed values of

A and n.

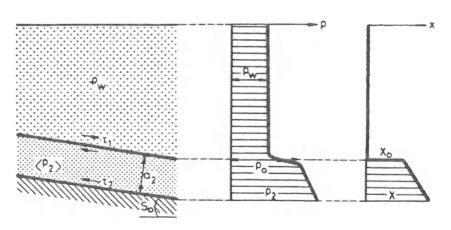


Figure ! Illustration of the notations

#### 4 NECESSARY CONDITION FOR MASS EROSION

Eq. 5 represents the equilibrium condition at the start of, and during the mass erosion of the mud layer: the shear stress induced by gravity is then balanced by frictional resistance.

Since however mud behaves approximately as a pseudoplastic (or Bingham) fluid, the static shear stress must be overcome to initiate a sliding of the mud layer. The static (or yield shear stress ("rigidity") is generally expressed as (Migniot)

$$\tau = A X^n$$

in which A and n are physical characteristics of the mud.

To obtain sliding of a layer of thickness a2

$$\alpha g a_2 (X_0 + \frac{1}{2} m a_2) S_0 > \tau_0$$

$$\tau_0 = A (X_0 + m a_2)^n$$
(6)

- 1. A mud layer, in which the density increases linearly with depth, can only slide under the influence of gravity, if the channel slope equals or exceeds S in . The value of S in depends on the density of the mud layer. To obtain mass erosion at S in the layer should have a thickness of a 2m. If the mud layer is thicker, only a layer of thickness a will move. Thinner layers however cannot alide, since the shear stress due to gravity is not sufficient to overcome the static shear stress in the mud.
- On channel slopes S > S<sub>min</sub>, mud layers can slide down, provided that their thickness, a<sub>2</sub> is in excess of a<sub>2</sub>. Of mud layers thicker than a<sub>2</sub>, only a layer of thickness a<sub>2</sub> would slide.

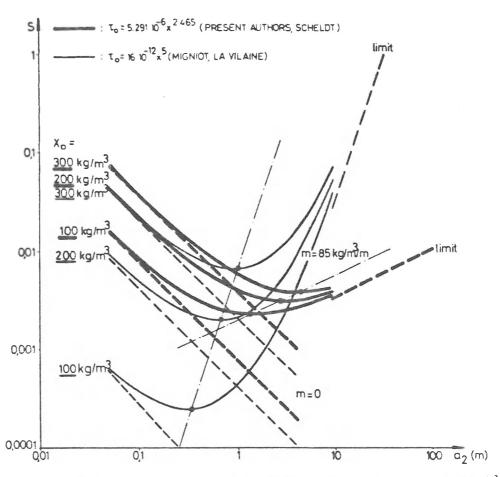
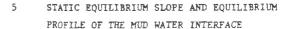
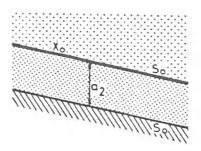


Figure 3 Channel slope versus thickness of the sliding mud layer, for m=0 and 85 kg/m³/m, and two different muds ( $\alpha=0.625$ )

- 3. The minimum slope, under which no mass erosion occurs, increases with the mud density  $X_0$  and decreases with increasing thickness  $a_2$  (provided that  $a_2 < a_{2m}$ ). As a consequence, the slope needed to set on a global slide in a hydraulic model, may be considerably larger than in the prototype (scale effect).
- 4. If the density of the mud layer is constant (m = 0), no S<sub>min</sub> exists. Neither does a maximum thickness a<sub>2</sub>. For a mud of given density, a minimum thickness to obtain mass erosion exists for each channel slope.



It follows from the previous paragraph that for mud of a given density in a prismatic channel at a given slope, a minimum thickness  $\mathbf{a}_2$  of the layer exists under which no (uniform) motion is possible. It is the value  $\mathbf{a}_2$  at which the static shear  $\tau_0$  is in equilibrium with the shear stress induced by gravity. It is suggested here, that mud can only be pumped provided that the slope of the water-mud interface is in excess of the static equilibrium slope, determined by the thickness of the layer.



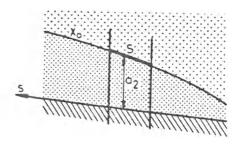


Figure 4 Comparison of uniform flow with gradually varied flow

Assuming that the theory for uniform flow holds approximately true for gradually varied flow, the static equilibrium profile of the mud water interface can be obtained by integrating the relation between S and  $\mathbf{a}_2$ : from Eq. 6 it follows:

$$\frac{da_2}{ds} = S = \frac{\tau_0}{\tau_2} = \frac{A (X_0 + m a_2)^n}{\alpha g a_2 (X_0 + \frac{1}{2} m a_2)}$$
(7)

Fig. 5 shows some equilibrium profiles for différent values of  $\boldsymbol{X}_{\underline{\alpha}}$  .

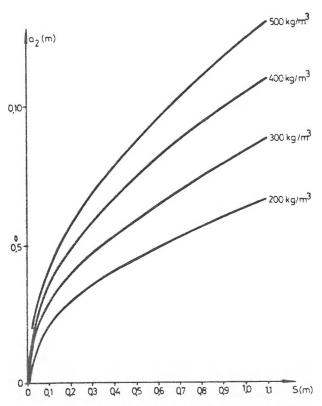


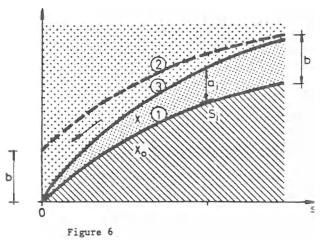
Figure 5 Static equilibrium profiles

0,8

0,6

## 6 EVOLUTION OF THE EQUILIBRIUM PROFILES IN THE CASE OF LAYERED MUD DEPOSITS

To explain the strong difference between the behaviour of homogeneous and stratified mud under pumping conditions, the following theory is put forward.



After the first pumping in O of homogeneous mud of solids concentration X<sub>O</sub>, the mud water interface is formed by the static equilibrium profile of the mud X<sub>O</sub> (① in Fig. 6). Suppose that, due to tidal effects a layer of thickness b and density X settles out of a suspension: a new mud water interface is formed (② in Fig. 6). Due to a renewed pumping action the mud X flows towards the pumping point under the influence of gravity (the mud X cannot be pumped since it is situated under its static equilibrium profile). At each distance s from O the static equilibrium profile of the X mud, is in fact the "bottom" for the X mud flow, having a slope S<sub>1</sub>. On this slope X mud can flow downward under the effect of gravity, provided that the thickness of the X mud layer exceeds the minimum value a<sub>1</sub>, determined according to Eq. 6. Integrating this equation

LAYER	X kg/m <sup>3</sup> Xi	LAYER	X kg/m <sup>3</sup>	SUBSIDENSE FACTOR
0	300	0	300	0,85 - 0,75 - 0,65
1 9	100	1 9	100	0,65 - 0,55 - 0,45

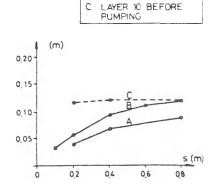
(m)

0,20

0,15

0 10

0.05



A: AFTER 5 CYCLES
B: AFTER 10 CYCLES

Figure 7 a)

04

(m)

0,20

0.15

0.10

0 05

Figure 7 b)

Figure 7 c)

Comparison between calculated mud water interfaces both without taking into account the subsidence effect a) and with the subsidence effect b), and experimental values c). The subsidence is simulated during the first thee cycles for each layer; for layer 0 with subsidence factors 0.85, 0.75 and 0.65 respectively, and for layers 1 to 9 with factors 0.65, 0.55 and 0.45 respectively.

0,4

0,6

s (m)

0.8

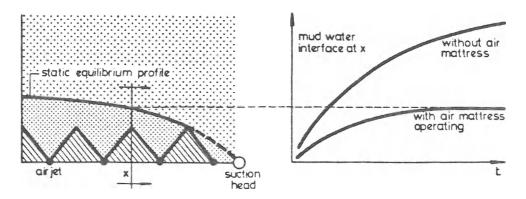


Figure 8 Schematic representations of the evolution of the mud water interface under the effect of pumping and homogenizing with air jets

starting at 0 gives the new equilibrium profile (③ in Fig. 6), under which the mud cannot be removed by pumping. The whole procedure is repeated during the successive sedimentation and pumping cycles. It thus becomes clear that the mud water interface builds up, regardless the action of the pump. Carrying out the simulation of the described phenomenon on a computer, the effect of the subsidence of the successive layers may be taken into account (Fig. 7).

#### 7 CONCLUSION

It became evident from experiments both in the laboratory and in situ, and can be explained by theory, that homogeneous and layered mud behave in a quite different way when pumped. In the particular case of layered mud, after each pumping cycle, always a layer of mud remains, which cannot be removed by pumping: the bottom builds up regardless the pumping. Homogeneous mud however can be pumped until the equilibrium profile is reached.

As a consequence, to remove layered mud deposits by pumping, the mud must be homogenized before pumping. This could be done by water jetting, or by using an air bubble mattress, located on the fixed bottom underneath the suction heads of the pumps. Because of its moderate energy consumption an air bubble mattress was prefered and its effect tested both in the laboratory and in situ. It was found that using a bubble mattress, combined with a fixed pump installation the mud level can be maintained constant, thus substituting expensive conventional dredging (Fig. 8).

At Leuven, laboratory tests are actually carried out, to optimize the disposition of the air jets, the air discharge, the aeration frequency, the pumping frequency and the mud pumping capacity. The system is covered by international pattent rights by S.B.B.M. (Belgium).

#### 8 ACKNOWLEDGMENT

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