LIVIAN

Edouard POTY¹ & Luc HANCE²

(4 figures)

- 1. Paléontologie animale et humaine, Département de Géologie, Université de Liège, Bât. B18, Allée du 6 Août, Sart Tilman, 4000-Liège, Belgium; E-mail: e.poty@ulg. ac. be.
- 2. Carmeuse Coordination Center, Bd de Lauzelle, 65, 1348 Louvain-la-Neuve, Belgium; E-mail: Luc.hance@skynet.be.

ABSTRACT. The Livian Substage is about 3-4 Ma long and was defined as corresponding to the Cf5 Foraminifer Zone by Conil *et al.*, 1977 (= MFZ12 Zone of Devuyst & Hance, in Poty *et al.*, in press). Its base is taken at the base of the "Banc d'or de Bachant", a bentonite capping the Moliniacian Neffe Formation (Fm). The top of the substage corresponds to the top of the Bay-Bonnet Member (Mbr) of the Grands-Malades Fm, below the appearance of the first foraminifers marking the base of the Warnantian Substage, at the base of the Thon-Samson Mbr (Bonne River Fm). The substage comprises the Lives Fm, including the Haut-le-Wastia, Corphalies and Awirs Mbrs, and the Grands-Malades Fm including the Seilles-Maizeret and Bay-Bonnet Mbrs. It is mainly composed of shallowing-upward parasequences which were deposited by aggradation on the shelf as suggested by their lateral regularity. The Livian Substage comprises the third-order sequence 7 (from the base of the Lives Fm to the top of the Seilles-Maizeret Mbrs) and the TST of the sequence 8 (Bay-Bonnet Mbr) of Hance *et al.* (2001). The Livian corresponds to the RC5γ Coral Subzone and the RC6 Zone. The Livian can be correlated with the British Holkerian Substage.

KEYWORDS: Livian, Middle Viséan, Carboniferous, Belgium, biostratigraphy, lithostratigraphy, palaeogeography.

Reference to this paper: Poty, E. & Hance, L., 2006. Livian. *In* Dejonghe, L., ed., Current status of chronostratigraphic units named from Belgium and adjacent areas, *Geologica Belgica*, Brussels, 9/1-2: 133-138.

1. Name

Livian (English), Liviaan (Dutch), Livium (German), Livien (French).

2. Age

From 335 to 331 Ma (i.e. 4 Ma long) in a time scale with minimum ages and from 337.5 to 334.5 (i.e. 3 Ma long) in a time scale with maximum ages (Menning *et al.*, 2001).

3. Authors

Conil, R., Groessens, E. & Pirlet, H., 1977. Nouvelle charte stratigraphique du Dinantien type de la Belgique. Annales de la Société géologique du Nord, XCVI: 363-371.

The Substage corresponds to the Cf5 Foraminifer Zone of Conil et al. (1977), which is characterized by the appearance of Koskinotextularia and Pojarkovella nibelis (Conil et al., 1977; Paproth et al., 1983; Conil et al., 1991). The base of the Substage is taken at the base of the « Banc d'or de Bachant », an argillaceous bed (bentonite L1 of Delcambre, 1989) separating the Lives Formation (Fm)

from the underlying Neffe Fm. Its top is below the entry of the foraminifers of the Cf6 Zone of Conil *et al.* (1977) at the base of the Thon-Samson Member (Mbr) (base of the Warnantian Substage).

4. Historical type area

Cliff and quarry situated at Lives, on the right bank of the Meuse valley, downstream from the Beez motorway bridge (Fig. 1). The base of the stage is exposed near the

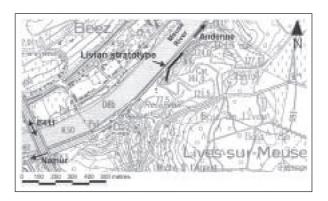


Figure 1. Location of the Livian stratotype.

foot of the cliff (Fig. 2) along the highway N90 (Hance, 1979).

Geological map n° 144 (Namur-Champion, 1901), under revision by Delcambre & Pingot (in preparation).

5. Description

In the stratotype, the underlying Neffe Fm is dolomitized and not well exposed. Only the lower part of the substage is visible, corresponding to the Lives Fm, including the Haut-le-Wastia Mbr (32 m thick), the Corphalie Mbr (19.6 m thick) and the lower part of the Awirs Mbr (about 20 m thick). See 7 for lithological description. The Livian type area was part of the Namur sedimentation area (NSA; Hance *et al.*, 2001).



Figure 2. Overview of the Livian stratotype.

6. Historical background

The Viséan was previously considered as a series divided into three stages (Conil *et al.*, 1977). From base to top: Moliniacian (lower Viséan), Livian (middle Viséan) and Warnantian (upper Viséan). It has recently been redefined as a stage by the IUGS Subcommission on Carboniferous Stratigraphy (Heckel, 2004), and Livian consequently became a substage.

7. Lithology

The substage is about 136 m thick and includes the Lives Fm and the Grands Malades Fm.

The Lives Fm is composed of well-bedded, pale grey to dark limestones, often rich in rugose corals and productids, organized in shallowing upward parasequences. Its base corresponds to a 10-30 cm thick argillaceous bed of volcanosedimentary origin (« Banc d'or de Bachant »;

bentonite L1 of Delcambre, 1989), locally pedogenized. It comprises three members, from its base to the top: The Haut-le-Wastia Mbr (« $V2b\alpha$ » of Conil et al., 1967), composed of 30 m of thick-bedded, pale grey to grey limestones in parasequences in which stromatolites and lime mudstones are dominant, and brecciated in places due to the dissolution of evaporitic layers.

The Corphalies Mbr (« $V2b\beta$ », ibid.), composed of thick-bedded, dark, bioclastic limestones overlain by thin-bedded, dark, lime mudstones, and forming a single, 18 m-thick, parasequence ("sequence number 0"). Some

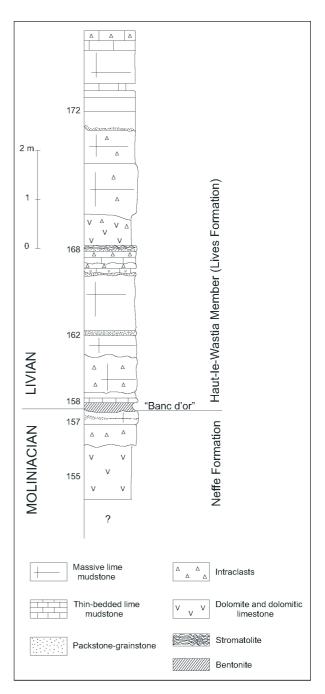


Figure 3. Log of the base of the Livian in the stratotype (from Hance, 1979)

LIVIAN 135

metric to decametric microbe-bryozoan reefs are locally present (Lauwers, 1992; Chevalier & Aretz, 2005). A bentonite of volcanic origin (bentonite L3 of of Delcambre, 1989), 5-30 cm thick, is present in the upper part of the parasequence and is recorded everywhere in the basin.

The Awirs Mbr (« V2by- ε », ibid.), 35 m thick mainly

The Awirs Mbr (« V2b γ – ϵ », ibid.), 35 m thick, mainly composed of dark grey limestones in which bioclastic packstones are dominant, and locally rich in cherts.

The Grands Malades Fm comprises the Seilles Mbr (eastern part of the NSA) that passes laterally to the Maizeret Mbr, and the overlying Bay-Bonnet Mbr.

The Seilles Mbr is 35 to 40 m thick and is composed of thick-bedded, pale limestones, organized in shallowing upward parasequences, dominated by bioclastic limestones in which brachiopods and rugose corals are common, and capped by algal boundstones and mudstones.

The Maizeret Mbr is 35 m thick and is composed of thick-bedded, pale limestones (mainly algal boundstones and mudstones-wackestones), dolomites, breccias, and thick beds of coarse or prismatic calcite (pseudomorphs of evaporites); breccias can dominate.

The Bay-Bonnet Mbr is composed of 8 m of thick-bedded, pale, laminated stromatolitic limestones (boundstones), with numerous serpulids.

The top of the substage corresponds to the top of the Bay-Bonnet Mbr, the base of the overlying Thon-Samson Mbr (Bonne River Fm), ie. a bed of crinoidal limestone, corresponds to the base of the Warnantian Substage. In the Boulonnais (Poty, 1994), the substage comprises the

In the Boulonnais (Poty, 1994), the substage comprises the upper part of the Haut-Banc Limestone (corresponding to the Haut-le-Wastia Mbr and to part of the Corphalie Mbr up to the bentonite L3; the lower part of the Haut-Banc Limestone corresponding to the Terwagne and Neffe Fm), the *Siphonodendron martini* Dolomite (from the bentonite L3 of the Corphalie Mbr to the top of the Awirs Mbr), the Lunel Limestone (= Seilles Mbr) and the Napoleon Limestone (= Bay-Bonnet Mbr).

In the Campine Basin (north of the Brabant Massif), Livian deposits seem to be generally missing. The base of the substage can be correlated with the base of the Clifton Down Limestone of the Avon section (Bristol) in Southwest England (Hance *et al.*, 2002), and with the base of the Holkerian British Substage in North Wales (Great britain).

8. Sedimentology and palaeogeography

The Lives Fm corresponds to the LST (Haut-le-Wastia Mbr) and to the TST (Corphalies and Awirs Mbrs), and the Seilles – Maizeret Mbrs to the HST (and to the FSST?), sensu Plint & Nummedal (2000) of the third-order sequence 7 of Hance *et al.* (2001). The Bay-Bonnet Mbr corresponds to the TST of sequence 8. The Livian is composed of shallowing-upward parasequences (about 33. 1 to 18 m-thick, can be recognized in sequence 7), comprising typically from the upper part of the Haut-

le-Wastia Mbr: bioclastic pack-grainstones rich in corals, brachiopods, foraminifers and calcareous algae in their lower part and mudstones and stromatolitic boundstones in their upper part.

The lithology and the lithostratigraphy of the substage is relatively constant, suggesting deposition by aggradation on the shelf areas that were covering the whole Namur-Dinant Basin, from eastern Belgium to northern France (Boulonnais and Avesnois).

During the Livian however, more open marine facies developed progressively to the north, while restricted facies and evaporites developed in the southern part of the basin. It is mainly the case in the Haut-le-Wastia Mbr and Maizeret Mbr, giving rise respectively to the "Petite Brèche Viséenne" and to the "Grande Brèche Viséenne" by solution collapse. This inversion in the shape of the basin could be related to an early phase of Variscan shortening.

9. Palaeontology

9.1. Foraminifers

The Livian Substage is covered by the MFZ12 Zone of Devuyst & Hance (in Poty et al., in press) which marks a drastic change in the composition of foraminifer associations (= Cf5 Zone of Conil et al., 1977, 1991). The guide Pojarkovella nibelis enters at very short distance above the base of the Lives Fm in the southernmost part of the DSA. In the other sedimentations areas of the Namur-Dinant Basin, the base of the Lives Fm is almost devoid of plurilocular foraminifers and P. nibelis enters high in the Haut-le-Wastia Mbr (Poty et al., 2002). In the Livian stratotype, the first occurrence data of P. nibelis is 14.3 m above the base of the substage. In addition to this index taxon, Koskinotextularia is a useful auxiliary guide entering at a similar level, low in the zone. Other typical elements include Palaeotextularia gr. consobrina, Lituotubella sp., Rhodesinella sp., Omphalotis minima, abundant Eostaffella spp. and Endothyra spp., Nodosarchaediscus spp. and Archaediscus spp. Associations are more diversified in the upper part of the zone (Grands Malades Fm), where large forms are common.

9.2. Rugose corals

The Livian is covered by the RC5γ Subzone and the RC6 Zone (Poty, 1985; Conil et al., 1991). The RC5 γ Subzone is characterized by Axophyllum vaughani, Siphonodendron irregulare, Caninophyllum archiaci and Clisiophyllum garwoodi with the appearance in the first open-marine bioclatic limestones in the Haut-le-Wastia Mbr, close to the base of the Substage,.

The base of the Corphalie Mbr is characterized by a horizon which contains the first *Lithostrotion araneum*, the oldest representative of the genus *Lithostrotion*. It

marks the base of the RC6 Zone (Lithostrotion araneum interval Zone, Poty, 1985). This horizon can be traced as far as the Avon Gorge section (Bristol, England), where it occurs at the base of the Lower Clifton Down Limestone ("Lithostrotion basaltiforme band" of Vaughan, 1906). Lithostrotion vorticale, Siphonodendron sociale and Siphonophyllia siblyi appear in the Livian (Zone RC6). Note that the most common coral species in the Livian are the long-ranging Siphonodendron martini and S. irregulare. The top of the RC6 Zone corresponds to the top of the Substage (Poty et al., in press).

9.3. Other fossils

Brachiopods (mainly productids and Composita) are common, but a revision is needed for a better evaluation of their biostratigraphic potential. Goniatites are very uncommon and without stratigraphic value to define the substage. Facies are unsuitable for conodonts and most common taxa belong to the genera Cavusgnathus and Taphrognathus (Webster & Groessens, 1991).

10. Chronostratigraphy

Fig. 4 gives the Foraminiferal and Coral Zonations used in Western Europe. More details about the chronostratigraphy and the stratigraphical correlations outside Western Europe are given in Poty *et al.* (in press).

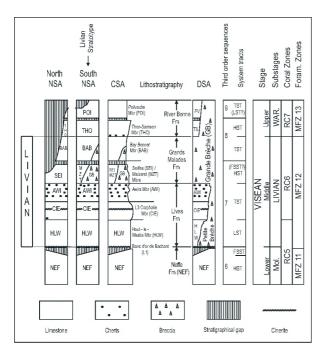


Figure 4. Litho-, bio- and sequence stratigraphy of the Livian Substage in the Namur (NSA), Condroz (CSA) and Dinant (DSA) sedimentation areas

11. Geochronology

No radiometric dates are available for the Livian in Belgium. The time scale used is based on datations obtained from other Carboniferous basins in Europe and in Australia (Menning *et al.*, 2001).

12. Structural setting

Livian deposits of the Namur-Dinant Basin are part of the Variscan orogenic cyclus. They are distributed on both sides of the Midi-Eifel fault. Deposits in the Ardenne Allochton and in the southern part of the Brabant Parautochton are largely folded. Those situated in the northern part of the Namur Synclinorium are gently dipping to the south.

13. Reference sections in Belgium

Numerous outcrops and quarries expose the base of the substage along the Meuse valley between Namur and Liège. Good reference sections for large parts of the substage including its base are the Awirs and the Engihoul quarries at Flémalle (NSA), the Leffe quarry at Dinant and the Haut-le-Wastia quarry at Anhée (DSA).

14. Main contributions

Gerards & Michot (1963); Pirlet (1963); Conil, Groessens & Pirlet (1977); Devuyst, Hance & Poty (2005); Hance (1979); Paproth *et al.* (1983); Conil *et al.* (1991); Lauwers (1992); Hance, Poty & Devuyst (2001, 2002); Poty *et al.* (2002); Chevalier & Aretz (2005).

15. Remarks

The base of the Livian Substage, defined at the base of a lithological marker bed (« Banc d'Or de Bachant ») and not coinciding with the entry of the palaeontological guides in the stratotype is not satisfactory. It limits the use of the Livian Substage in its original definition to the areas which were covered by the "Banc d'Or" bentonite, i.e., the Namur-Dinant Basin. That bentonite suffered pedogenis in many places and in the stratotype, and indicates a stratigraphical gap (erosion and/or break in sedimentation) between the third-order sequences 6 and 7 of Hance *et al.* (2001). The search for a lectostratotype using the stratigraphical incoming of a palaeontological guide as boundary criterion is in progress. The problem is difficult as generally the Moliniacian-Livian boundary interval displays very shallow-water facies with an

impoverished fossil association. It is not certain yet that the guide *Pojarkovella nibelis* is not entering in the late Moliniacian, in more open marine settings or that the Moliniacian guides persist into the early Livian. That coexistence between *P. nibelis* and Moliniacian guides has been found in South China (Devuyst *et al.*, 2003), in the Mendips (SW England, Austin *et al.*, 1973) and in the Visé sedimentation area in Belgium (Conil, in Kimpe *et al.*, 1978, p. 58).

16. Acknowledgements

F.X. Devuyst (Trinity College, Dublin) and M. Aretz (Universität zu Köln) kindly corrected the English and commented on an early version of this paper.

References

AUSTIN, R., CONIL, R. & RHODES, F.H.T., 1973. Recognition of the Tournaisian-Viséan boundary in North America and Britain. *Annales de la Société Géologique de Belgique*, 96 (1): 165-188.

CHEVALIER, E. & ARETZ, M., 2005. A microbebryozoan reef from the Middle Visean of the Namur Syncline (Engihoul quarry). *Geologica Belgica*, 8/1-2: 109-119.

CONIL, R., GROESSENS, E., LALOUX, M., POTY, E. & TOURNEUR, F., 1991: Carboniferous Guide Foraminifera, Corals and Conodonts in the Franco-Belgian and Campine Basins: their potential for widespread correlations. *Courier Forschungsinstitut Senckenberg*, 130: 15–31.

CONIL, R., GROESSENS, E. & PIRLET, H., 1977. Nouvelle charte stratigraphique du Dinantien type de la Belgique. *Annales de la Société géologique du Nord*, XCVI: 363-371.

CONIL, R., PIRLET, H. & LYS, M., 1967. Echelle biostratigraphique du Dinantien de la Belgique. Service Géologique de Belgique, Professional Paper, 13: 1-56.

DELCAMBRE, B.,1989. Marqueurs téphrostratigraphiques au passage des calcaires de Neffe vers ceux de Lives. *Bulletin de la Société belge de Géologie*, 98, 2: 163-170.

DEVUYST, F.X, HANCE, L., HOU, H., WU, X., TIAN, S., COEN, M. & SEVASTOPULO, G. 2003. A proposed Global Stratotype Section and Point for the base of the Viséan Stage (Carboniferous): the Pengchong section, Guangxi, South China. *Episodes*, 26 (2): 105-115.

DEVUYST, F.X., HANCE, L. & POTY, E., 2005. The Dinantian of Southern Belgium revisited: sedimentary history and biostratigraphy. A guidebook of key sections. I.U.G.S. Subcommission on Carboniferous Stratigraphy field trip, May 2005: 1-79.

GERARDS, J. & MICHOT, P., 1963. Le Viséen moyen, partie supérieure: V2b. In MICHOT, P, GÉRARDS, J., MONTY, C. & PIRLET, H., Excursion G – 1° partie: Sédimentologie des formations viséennes du Synclinorium de Namur, dans la vallée de la Meuse. 6° Congrès International de Sédimentologie Belgique & Pays-Bas: 10-12.

HANCE, L., 1979. Description de deux nouvelles coupes dans le Moliniacien (V1-V2a) des environs de Namur. Service Géologique de Belgique, Professional Paper, 165: 1-14.

HANCE, L., POTY, E. & DEVUYST, F.X., 2001. Stratigraphie séquentielle du Dinantien type (Belgique) et corrélation avec le Nord de la France (Boulonnais, Avesnois). *Bulletin de la Société Géologique de France*, 172, 4: 411-426.

HANCE, L., POTY, E. & DEVUYST, F.X., 2002. Sequence stratigraphy of the Belgian Lower Carboniferous – tentative correlation with the British Isles. *In* HILLS, L.V., HENDERSON, C.M. & BAMBER, W. (eds.), Carboniferous and Permian of the World. *Canadian Society of Petroleum Geologists Memoirs*, 19: 41-51.

HECKEL, P.H., 2004. Chairman's column. Newsletter on Carboniferous Stratigraphy, IUGS Subcommission on Carboniferous Stratigraphy, 22: 1-3.

KIMPE, W.F.M., BLESS, M.J.M., BOUCKAERT, J., CONIL, R., GROESSENS, E., MEESSEN, J.P.M.Th., POTY, E., STREEL, M., THOREZ, J., VANGUESTAINE, M., 1978. Paleozoic deposits east of the Brabant massif in Belgium and Netherlands. *Mededelingen Rijks Geologische dienst (The Netherlands)*, 30-2: 37-103.

LAUWERS, A.S., 1992. Growth and diagenesis of cryptalgal-bryozoan buildups within a mid-Viséan (Dinantian) cyclic sequence, Belgium. *Annales de la Société géologique de Belgique*, 115: 187-213.

MENNING, M., WEYER, D., DROZDZEWSKI, G. & WENDT, I., 2001. More radiometric ages for the Carboniferous time scale. Newsletter on Carboniferous Stratigraphy, IUGS Subcommission on Carboniferous Stratigraphy, 19: 16-18.

PAPROTH, E., CONIL, R., BLESS, M. J. M., BOONEN, P., BOUCKAERT, J., CARPENTIER, N., COEN, M., DELCAMBRE, B., DEPRIJCK, CH., DEUZON, S., DREESEN, R., GROESSENS, E., HANCE, L., HENNEBERT, M., HIBO, D., HAHN, G., HAHN, R., HISLAIRE, O., KASIG, W., LALOUX, M., LAUWERS, A., LEES, A., LYS, M., OP DE BEEK, K., OVERLAU, P., PIRLET, H., POTY, E., RAMSBOTTOM, W., STREEL, M., SWENNEN, R., THOREZ, J., VANGUESTAINE, M., VAN STEENWINKEL, M. & VIESLET, J. L. 1983. Bio- and lithostratigraphic subdivisions of the Dinantian in Belgium, a review. *Annales de la Société Géologique de Belgique*, 106: 185-239.

PIRLET, H., 1963. La sédimentation rythmique de la partie inférieure du V3a dans le bassin de Namur, les relations entre le Dinantien et le Namurien de Namêche à Moha. *Annales de la Société géologique de Belgique*, 86: 461-468.

PLINT, A.G. & NUMMEDAL, D., 2000. The falling stage system tract: recognition and importance in sequence stratigraphic analysis. In: Sedimentary responses to forced regressions. *Geological Society of London, Special Publications*, 172: 1–17.

POTY, E., 1985: A rugose coral biozonation for the Dinantian of Belgium as a basis for a coral biozonation of the Dinantian of Eurasia. Compte-rendu X Congrès International de stratigraphie et de Géologie du Carbonifère, Madrid 1983, 4: 29-31.

POTY, E., 1994. Nouvelles précisions sur les corrélations stratigraphiques du Dinantien du Boulonnais et de la Belgique : application de la biozonation corallienne. *Comptes Rendus de l'Académie des Sciences, Paris*, 319, s. II: 467-473.

POTY, E., DEVUYST, F.X. & HANCE, L., in press. Upper Devonian and Mississippian foraminiferal and rugose coral zonation of Belgium and Northern France: a tool for Eurasian correlations. *Geological Magazine*. POTY, E., HANCE, L., LEES, A. & HENNEBERT, M., 2002. Dinantian lithostratigraphic units (Belgium);

In BULTYNCK, P. & DEJONGHE, L., eds., Guide to a revised lithostratigraphic scale of Belgium. *Geologica Belgica*, 4/1-2: 69-94.

VAUGHAN, A., 1906. The Carboniferous Limestone series (Avonian) of the Avon Gorge. *Proceedings of the Bristol Naturalist's Society*, 4th series, 1: 74-168.

WEBSTER, G.D., GROESSENS, E. 1991. Conodont subdivisions of the Lower Carboniferous. *In* BRENCK-LE, P.L., MANGER, W.L. (eds). Intercontinental Correlation and Division of the Carboniferous System. *Courier Forschunginstitut Senckenberg*, 130: 31-40.

Manuscript received on 14.05.2005 and accepted for publication on 18.10.2005.