

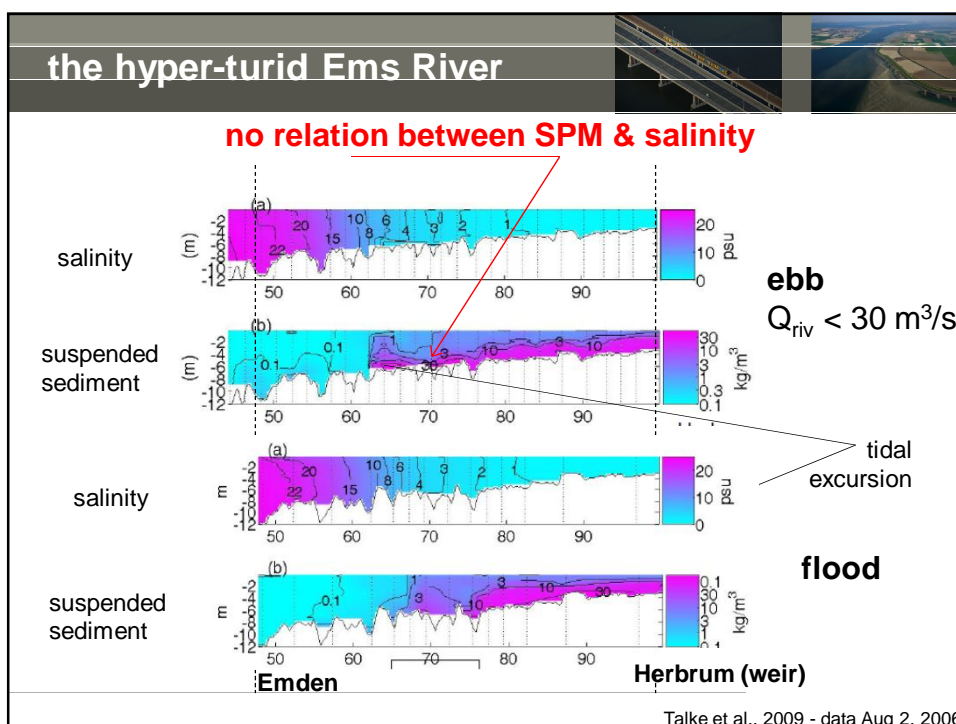


net transport of fine sediments in a narrow, converging estuary – the Sea Scheldt

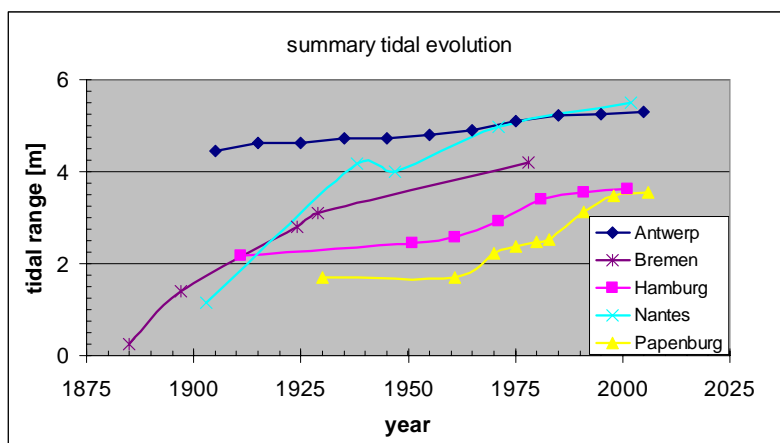
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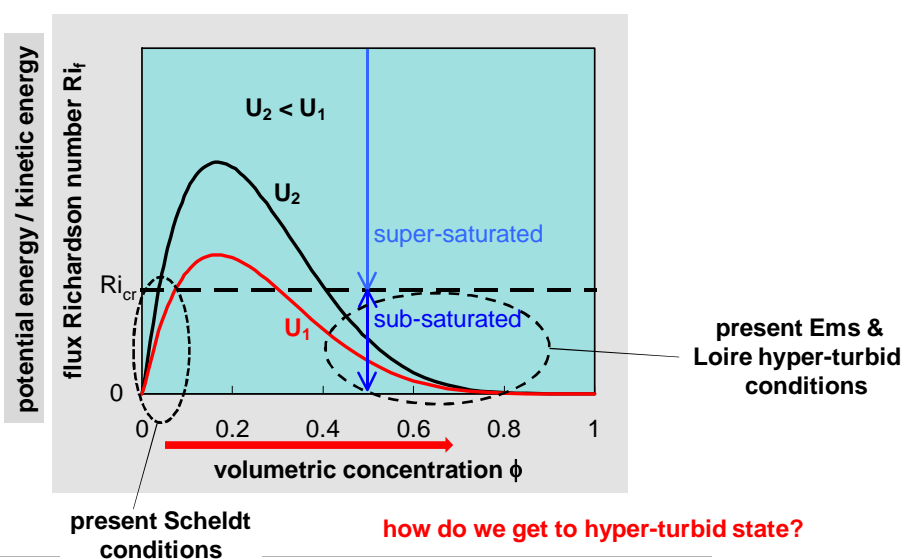


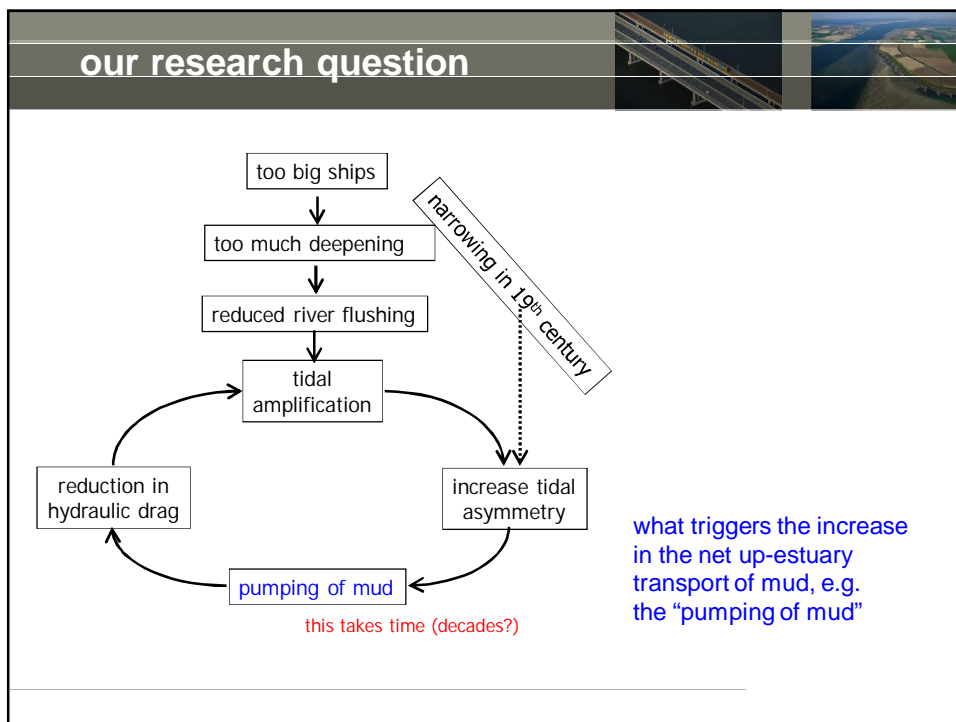
historical evolution of the tidal range



tidal amplification in Loire (Nantes) and Ems (Papenburg) can be explained for **50%** by deepening and narrowing and for **50%** by sediment-induced reduction in **hydraulic drag**

our research question – the stability diagram





net sediment transport processes

net sediment balance		SPM over water column	water-bed exchange
$\Delta\text{SPM}/\Delta t =$	import by estuarine circulation	2DV or 3D	not relevant
	import (or export) by tidal asymmetry	see next slide	necessary
	export by river flushing	depth-averaged	not relevant
	source by bed/bank erosion	depth-averaged	-
	sink by sedimentation	depth-averaged	-
	transport by tidal advection = gross effect		

2DV or 3D: vertical SPM gradients required for net transport

vertical gradients are affected by sediment-induced stratification

net transport by tidal asymmetry

net transport by tidal asymmetry			SPM over water column	water-bed exchange
alluvial bed (hyper-turbid estuary)	peak flow asymmetry	$T \propto U^4$	depth-averaged	necessary
	ebb/flood asymmetry (settling lag)	$(du/dt)_T$	depth-averaged	necessary
starved bed (“normal” estuary)	internal asymmetry (vertical mixing)	$\epsilon_z \propto U^2$	2DV or 3D	necessary
	slack water asymmetry (scour lag)	$(du/dt)_{u=0}$	depth-averaged	necessary

we start here, thus we look for balance between:

- estuarine circulation
- river-induced flushing
- tidal asymmetry in vertical mixing and slack water

asymmetry for starved bed conditions

for starved-bed systems: slack-water asymmetry & asymmetry in mixing

velocity U [m/s]

time t [hrs]

— $U^4 = 0.2 U^2$

vertical mixing ϵ_z [m^2/s]

time t [hr]

— mixing $\sim U^2$

flood = 1.26 x ebb

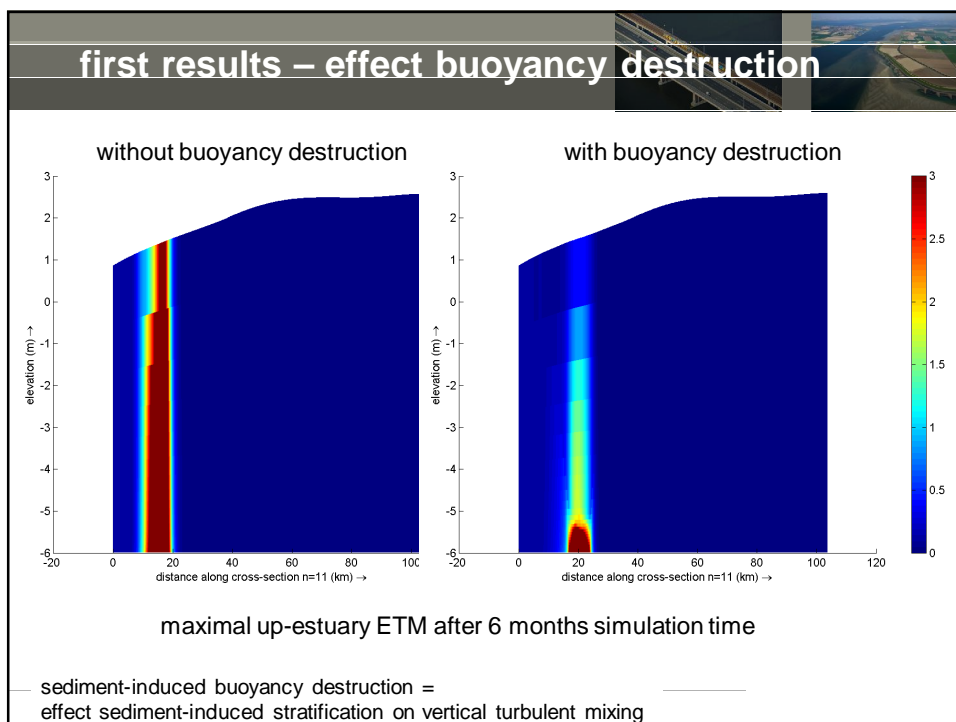
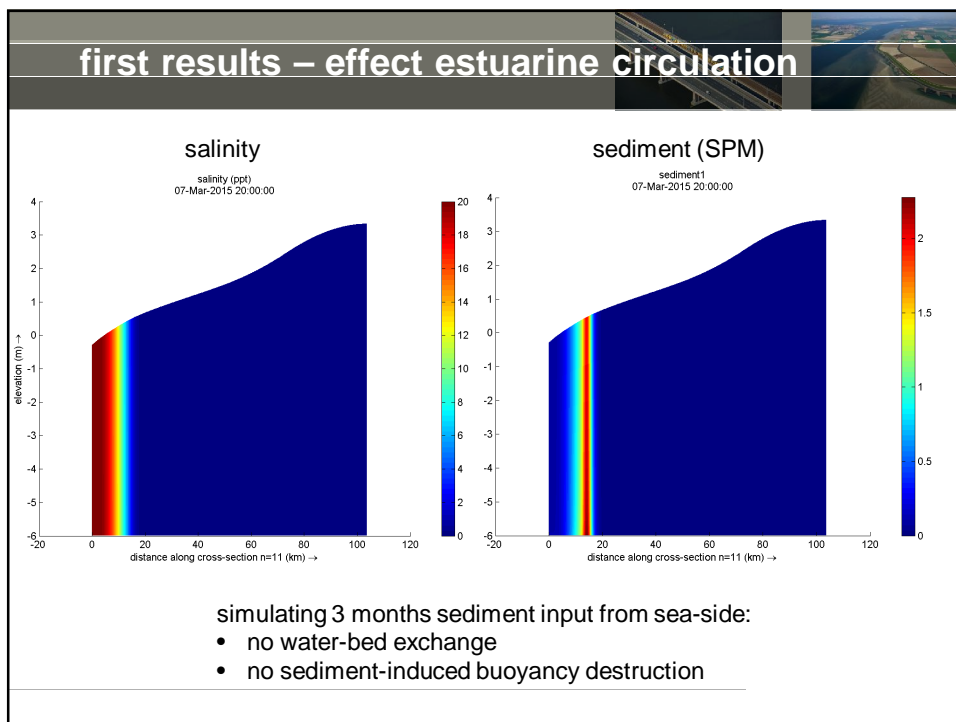
our approach

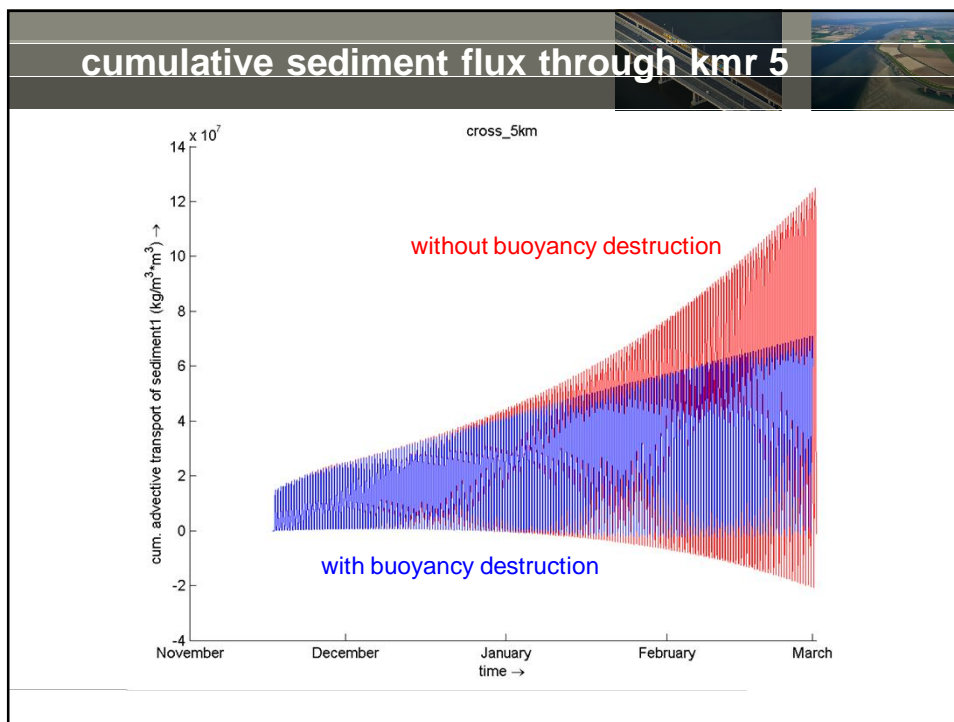
- idealized, converging estuary with constant depth
- dimensions, tidal forcing and river flow from Sea Scheldt river
- no intertidal area, hence **flood-dominant**
- implemented in Delft3D
- switch on/off various processes one by one (or combinations)
- vary boundary conditions (Q_{riv} and external M_4 amplitude & phase)
- vary internal processes (equation of state, erosion/deposition)

first results – development of asymmetry

- profound development of asymmetry in peak velocity, hence vertical mixing: induces **up-estuary transport**
- $t_{LWS} < t_{HWS}$, hence **down-estuary transport**, but reducing in up-estuary direction

mouth $x = 0$		maximum salinity intrusion		$x = 40$ km	
LWS	HWS	LWS	HWS	LWS	HWS
45 min	24 min	33 min	24 min	24 min	21 min





discussion and conclusions

- we used Delft3D for studying net transport of fine sediments in schematized estuary, resembling Sea Scheldt,
- this approach seems promising, but many more analyses are required,
- the first results indicate:
 - ETM can be produced by estuarine circulation alone,
 - sediment-induced buoyancy destruction keeps fines closer to the bed, reducing horizontal transports,
 - the realistic effect of sediment-induced buoyancy destruction is large, hence cannot be omitted in sediment transport studies,
 - tidal asymmetry without water-bed exchange does not induce net transport of fine sediment,
 - these numerical results are in line with theory, but have never been quantified.